Dynamique de processus océaniques de méso- et de subméso-échelle à partir de simulations numériques et de données *in situ*

Marion Kersalé Soutenance de Thèse – 15 Octobre 2013

Sous la présidence de : Ivan DEKEYSER

Comité : Gilles REVERDIN

Bruno BLANKE

Annalisa GRIFFA

Directeurs de thèse : Anne PETRENKO

Andrea DOGLIOLI











Dynamics of meso- and submesoscale oceanic processes from numerical simulations and *in situ* data

Marion Kersalé PhD defense – October 15, 2013

Under the presidence of : Ivan DEKEYSER

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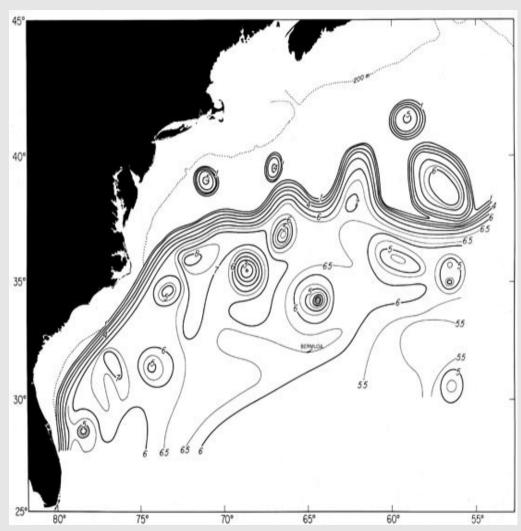




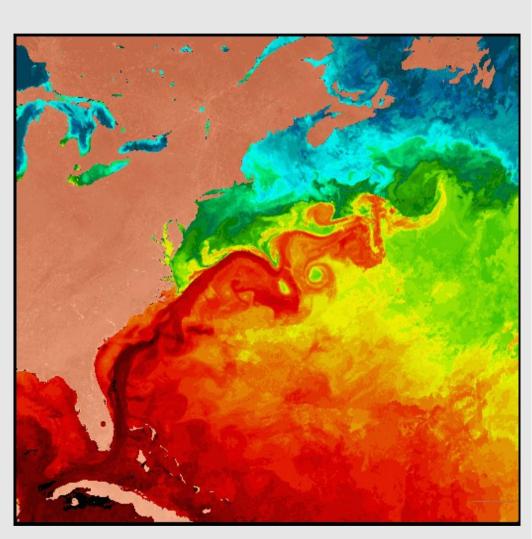


Introduction - Mesoscale variability

Discovering Gulf stream variability from space



Path of the Gulf Stream from hydrographic and satellite data (1975)



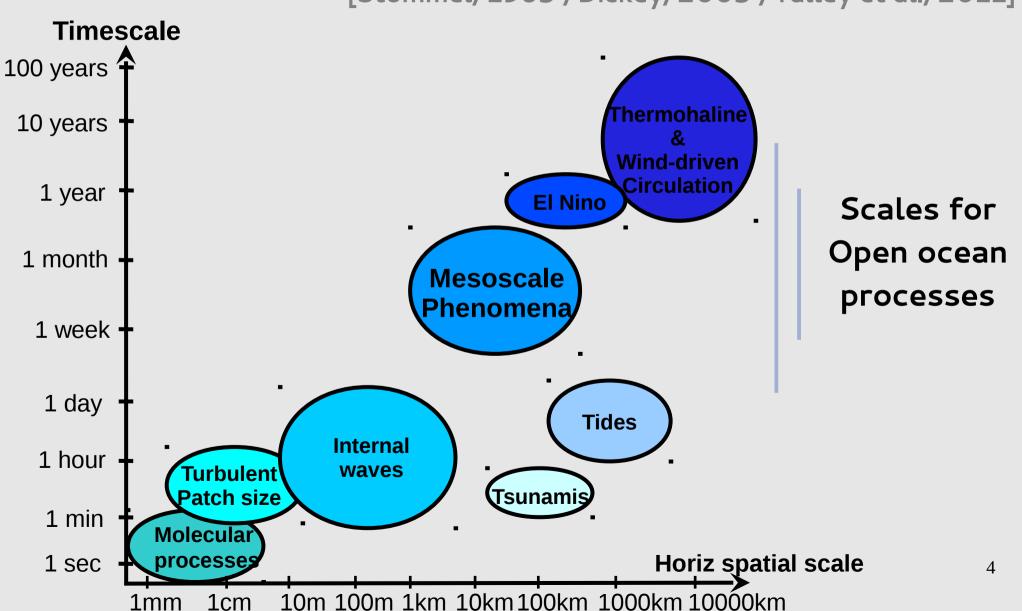
Thermal infrared image of the Gulf Stream ©NASA ³

Introduction - Time/space scale diagram



Numerous physical processes over a large range of scales

[Stommel, 1963; Dickey, 2003; Talley et al., 2011]



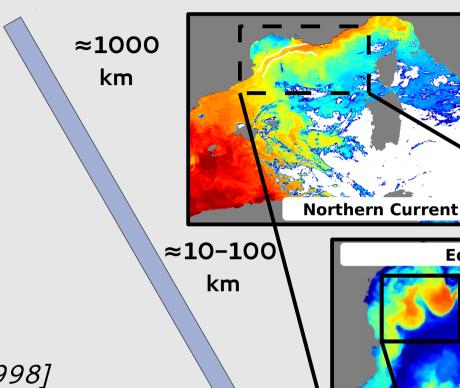
Introduction — Coastal processes scales

Smaller scales than the open ocean characterization

Large scale flow

Mesoscale

Scales



Geostrophic equilibrium ~O(R_d) [Chelton et al., 1998]

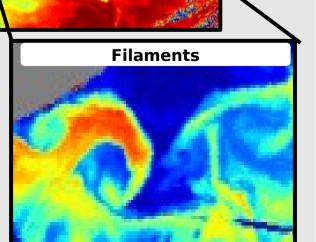
≈1-10

km

Submesoscale

Scales < R_d [Capet et al., 2008]

Numerous and energetic features => Transfer of energy



Sea Surface

Temperature

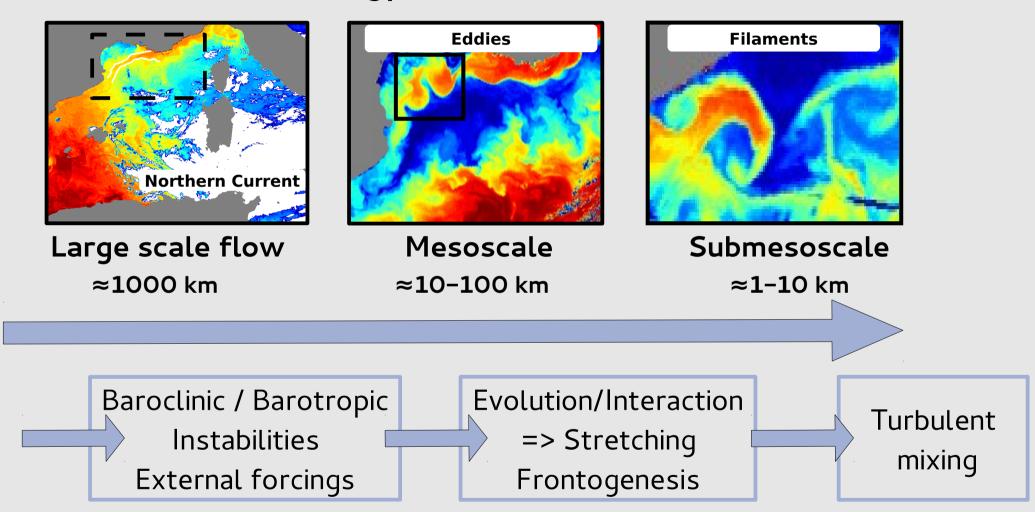
Images (SST)

©NOAA

Eddies

Introduction - Transfer of energy

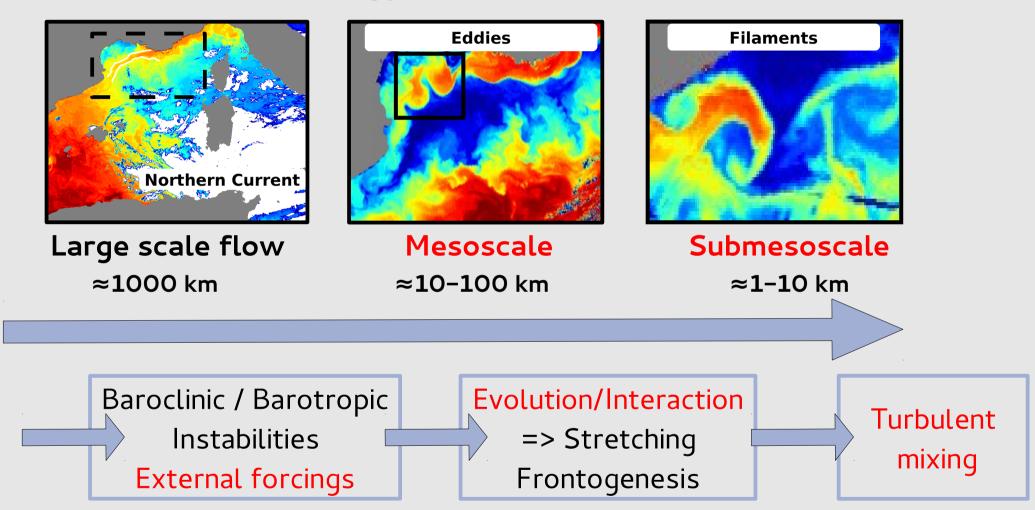
Forward transfer of energy



Processes involve in the forward transfer of energy

Introduction - Transfer of energy

Forward transfer of energy



Processes involve in the forward transfer of energy

Introduction - Deep ocean vs coastal





Hawaiian archipelago

- Deep ocean eddies
- Relative importance of the external forcings on the **generation** of mesoscale eddies

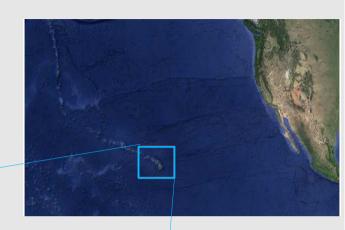
Gulf of Lion

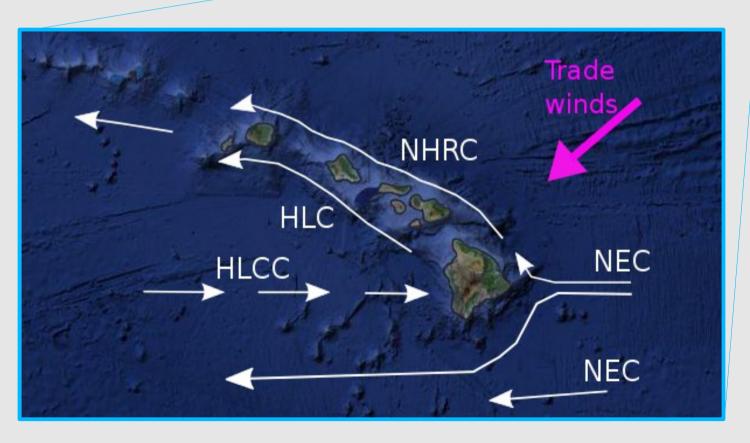
- Coastal eddies
- Process of generation
- Impact of meso- and submesoscale features on exchanges and turbulent mixing

Introduction - Area of study I

Hawaiian archipelago

- Complex circulation
- Effects of the archipelago topographic forcing on both the NEC and the trade winds
- The wake is responsible for the formation of the HLCC

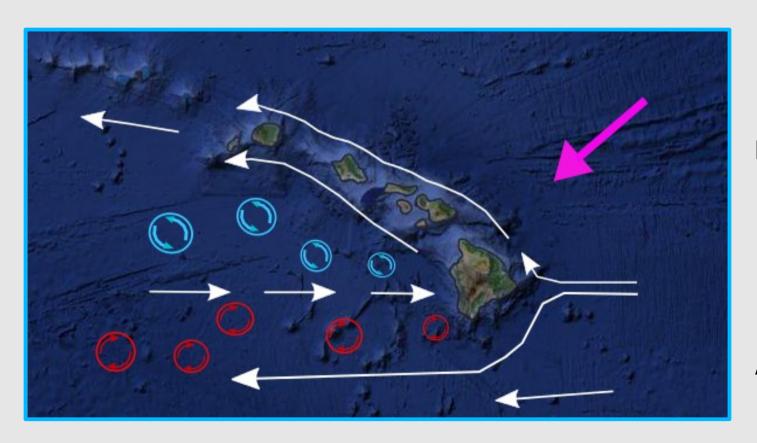




Introduction - Area of study I

Mesoscale activity in the lee of the islands

- Cyclonic/Anticylonic eddies North/South the HLCC [Lumpkin, 1998]
- Classical mechanism of formation of eddies in the lee of an obstacle
- The wind stress variations drive divergent and convergent Ekman transports in the upper layer of the ocean [Patzert, 1969; Chavanne, et al., 2002]



Northern Hemisphere



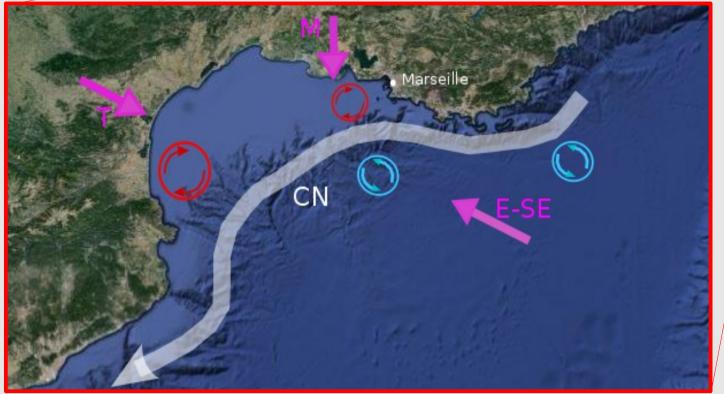


Introduction - Area of study II

Gulf of Lion (GoL)

- Hydrodynamics complex and highly variable [Millot, 1990]
- Influenced by the NC effective dynamical barrier
- Recurrent generation of mesoscale eddies [Petrenko, 2003; Hu et al., 2009; Allou et al., 2010; Schaeffer et al., 2011]

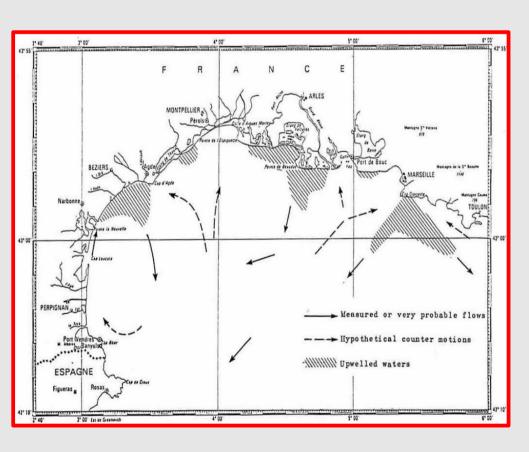


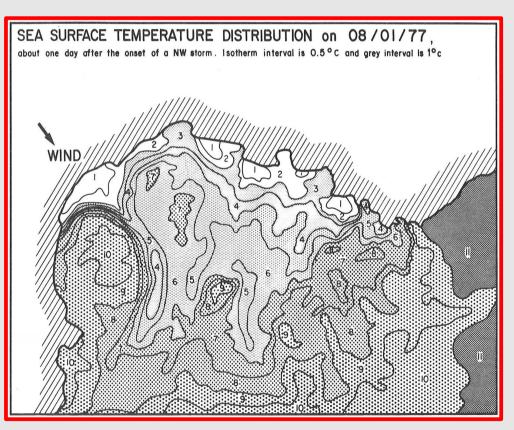


Introduction - Area of study II

Mesoscale activity in the western part of the GoL

- First observation of an anticyclonic circulation
- LATEX (<u>LAgrangian Transport Experiment</u>) Project Understand influence of coastal mesoscale eddies





Scheme of the circulation using mooring data [Millot, 1979]

Sea Surface Temperature Images (SST) [Millot, 1982]

Plan

Part I – Tools for study mesoscale features

Numerical codes
In situ experiments

Strategies

Part II - Characteristics of mesoscale eddies

Eddy center detection

Horizontal characteristics

Vertical characteristics

Part III – Forcings and generation

Forcing sensitivity – Wind sensitivity

Process of generation

Generation of submesoscale structures

Impact on coast-offshore exchanges

Part IV – Study of turbulent mixing

In situ experiment

Calculation of diffusion coefficients

Lagrangian experiment

Conclusion & perspectives

Mesoscale activity

Submesoscale

Turbulence

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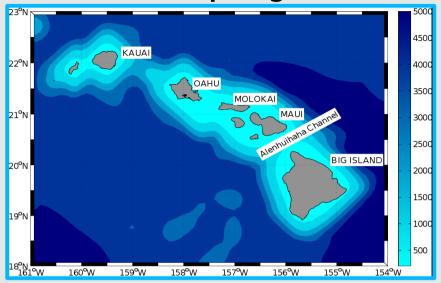
Submesoscale

Turbulence

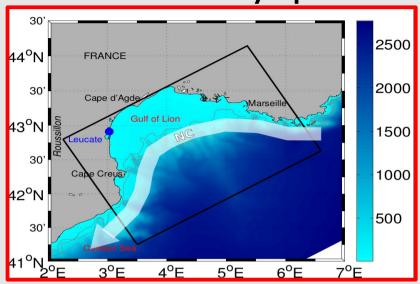
14

Part I - Numerical codes

Hawaiian archipelago - ROMS



Gulf of Lion - Symphonie

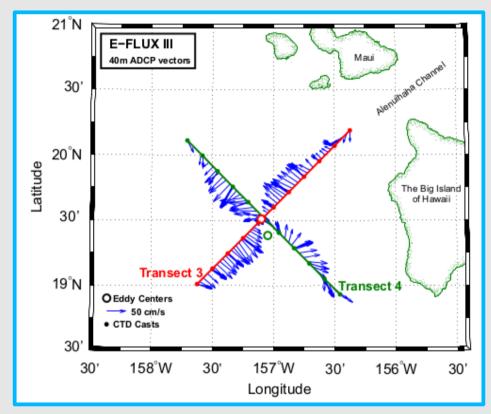


Model domains with realistic bathymetry [m]

	ROMS	Symphonie
Туре	Climatology — 10 years	Realistic — Year 2009
Resolution	1/10° (~10km)	1km
Horizontal grid	Arak	rawa C
Vertical grid	30 Sigma	40 sigma-z hybrid
Bathymetry	ETOPO	EPSHOM
Closure Scheme	Large et al. [1994]	Gaspar et al. [1990]
Atmospheric forcings	COADS/QuikSCAT	Météo-France Aladin
Boundaries	Radiation BC - WOA	Radiation BC - MFS
Initialization	WOA	Restart Hu et al. [2009]

Part I – *In situ* experiments

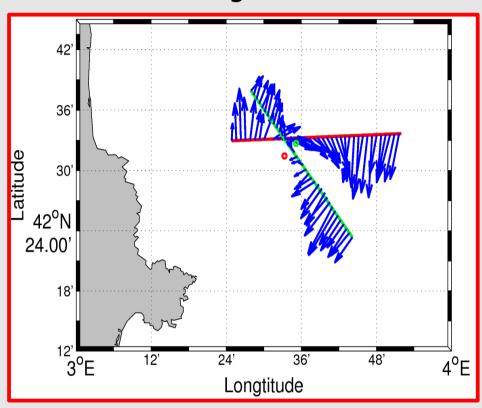
Hawaiian archipelago – E-Flux III March 10–27, 2005



ADCP current vectors at 40m depth [Nencioli et al., 2008]

 Sampling of a cyclonic mesoscale eddy: Opal [Dickey et al., 2008]

Gulf of Lion — Latex09 campaign August 24–28, 2009



ADCP current vectors at 15m depth

 Sampling of an anticyclonic mesoscale eddy: Latex-09 eddy

Part I - Strategies

In situ data

Opal

characteristics



Hawaiian archipelago

Climato. forcings
Wind Topo Currents

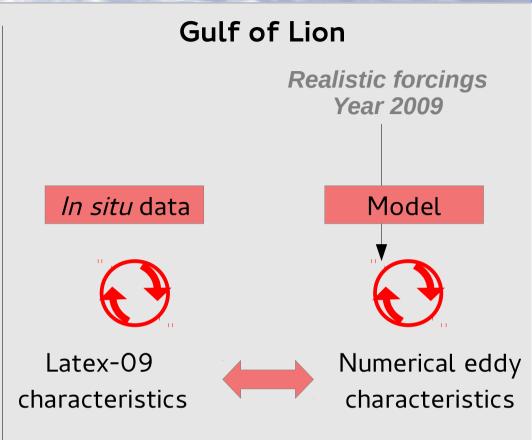
Model

Numerical eddies

characteristics

Validate the model and find the best set up

What is the **relative importance**of the forcings on the **generation of mesoscale eddies?**



Find a similar numerical eddy at the same period

What is the **generation process** of the Latex-09 eddy?

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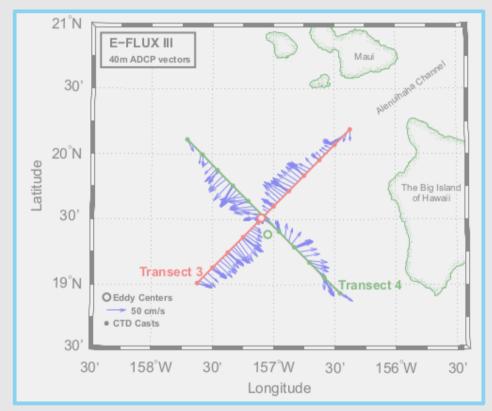
Mesoscale activity

Submesoscale

Turbulence

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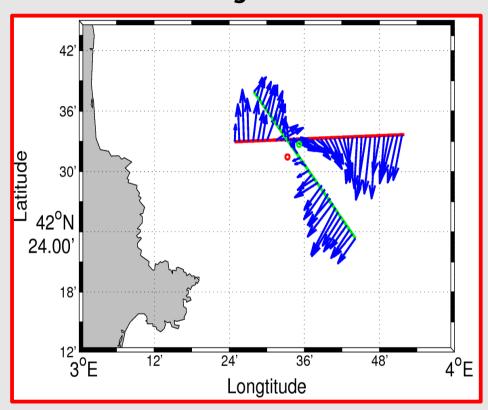
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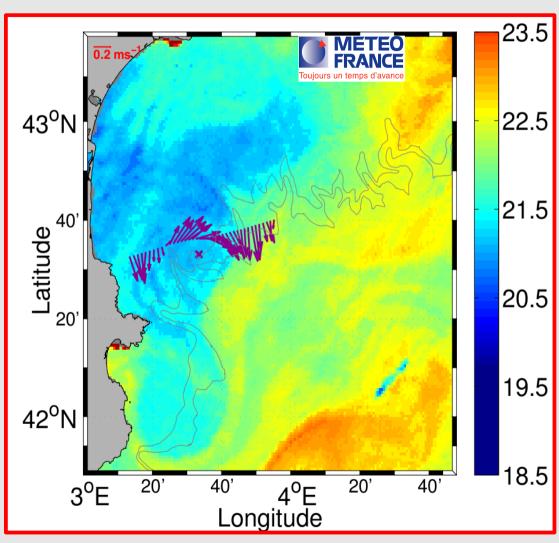
ADCP current vectors at 15m depth

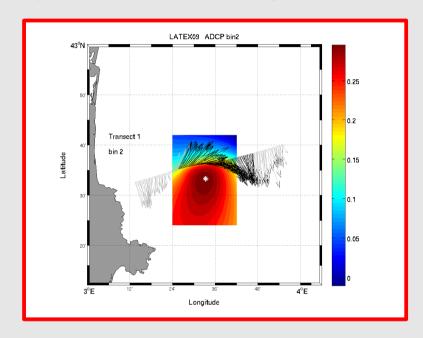
- Sampling of an anticyclonic mesoscale eddy: Latex-09 eddy
- Ships, surface drifters, and satellite sensors

Eddy center detection

Transect 1

Latex09 campaign
ADCP 15m depth - SST (°C) August 28





Center of the eddy:

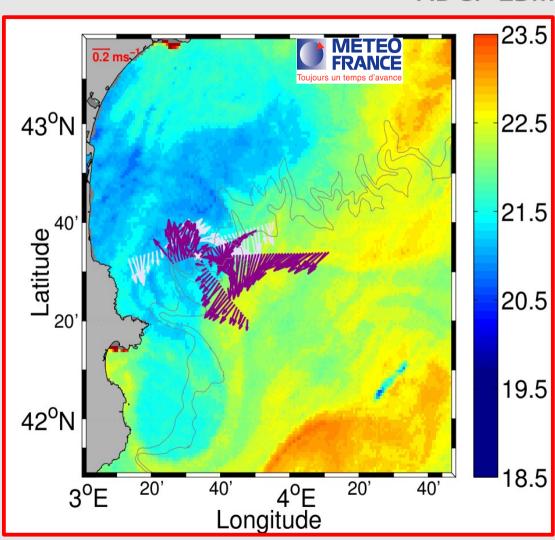
Grid point for which the mean tangential velocity computed from the nearest ADCP records is maximum
[Nencioli et al., 2008]



Horizontal Characteristics

Transect 1-2-3-4

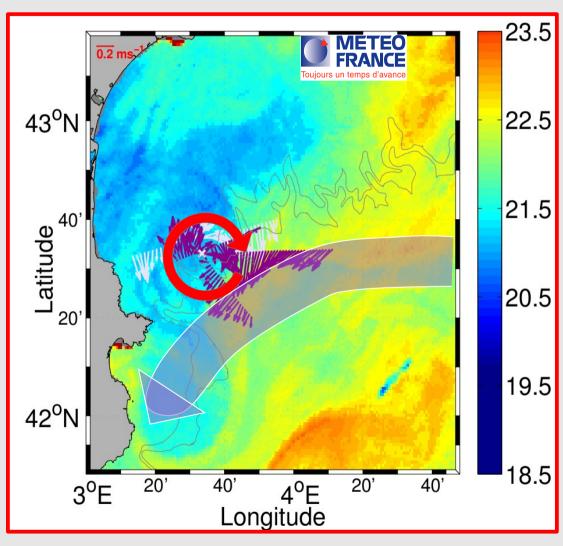
Latex09 campaign
ADCP 15m depth - SST (°C) August 28



Horizontal Characteristics

Transect 1-2-3-4

Latex09 campaign
ADCP 15m depth - SST (°C) August 28



Anticyclonic circulation

 $V_{max} \sim 0.4 \text{ m.s}^{-1}$ T ~ 3 days

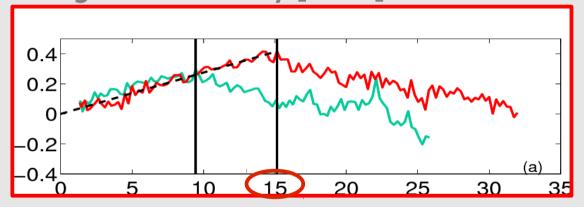
Center: 3°34'E – 42°33'N

Presence of the NC

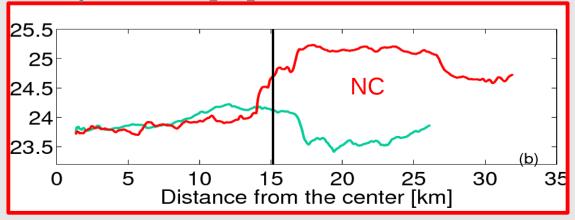
Horizontal Characteristics Transect 3

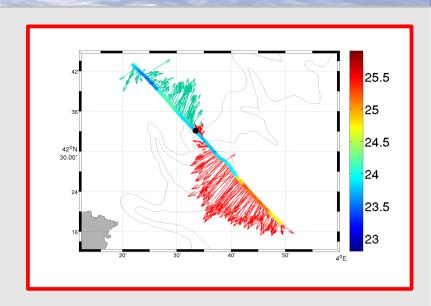
Tangential component decomposition with respect to the position of the center [Nencioli et al., 2008]

Tangential velocity [m s⁻¹]



Temperature [°C]



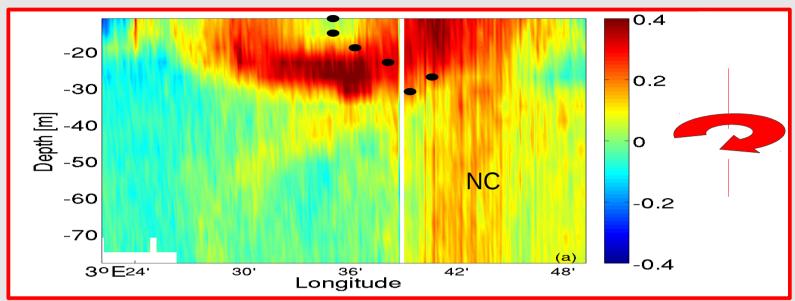


Latex09 campaign
ADCP 15m depth - SST (°C)

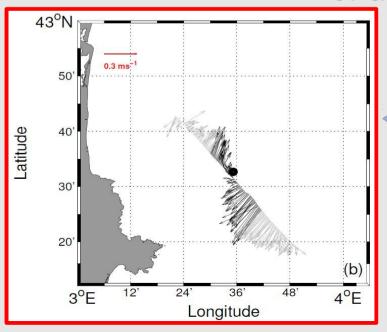
$$\mathbf{D_{eddy}} = \mathbf{\bar{D}} \pm \sqrt{\mathbf{D_{var}}}$$

$$D_{eddy} = 22.7 \pm 1.2 \text{ km}$$

Vertical Characteristics



Vertical section of the absolute tangential component of the horizontal current (m s⁻¹) for Transect 3



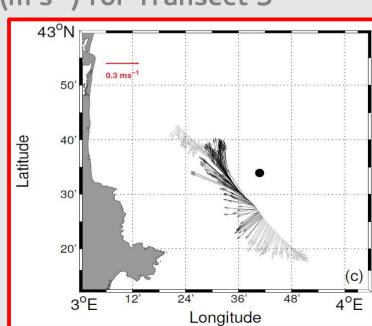
ADCP current at

15 m depth

27 m depth

Depth

30-35 m



Conclusions - Part II

Near real- time determination of eddy center

• Local **characterization** of coastal **mesoscale** features

Horizontal and vertical characteristics

Solid body rotation

Tilted axis of rotation

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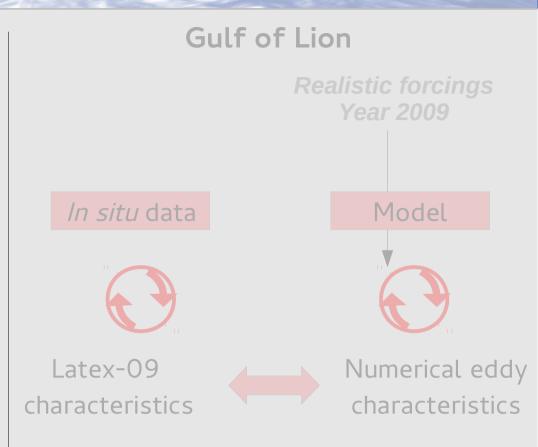
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Strategies

Hawaiian archipelago Climato. forcings Wind Topo Currents *In situ* data Model Opal Numerical eddies characteristics characteristics

Validate the model and find the best set up

What is the **relative importance**of the forcings on the **generation of mesoscale eddies?**



Find a similar numerical eddy at the same period

What is the **generation process** of the Latex-09 eddy?

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Relative importance of the forcings

	Wind Forcing	Advection	Drag coeff.
Run-A	none	Ugeo	Cd
Run-B	none	2Ugeo	Cd
Run-C	QuikSCAT	none	Cd
Run-D	QuikSCAT	Ugeo	none

Inadequacies:

Run-A & Run-B (No wind forcing)

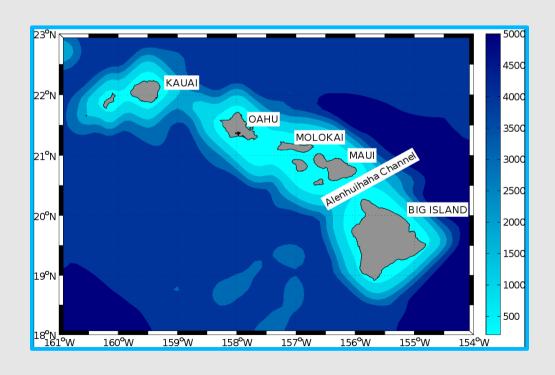
Ocean circulation not realistic Intensity of the eddies is too low

Run-C (No advection)

Stationary eddies- No expected westward drift

Run-D (No Drag. Coeff)

Eddies too large



Relative importance of the forcings

	Wind Forcing	Advection	Drag coeff.
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Run-B	none	2Ugeo	Cd
Run-C	QuikSCAT	none	Cd
Run-D	QuikSCAT	Ugeo	none

Inadequacies:

Run-A & Run-B (No wind forcing)

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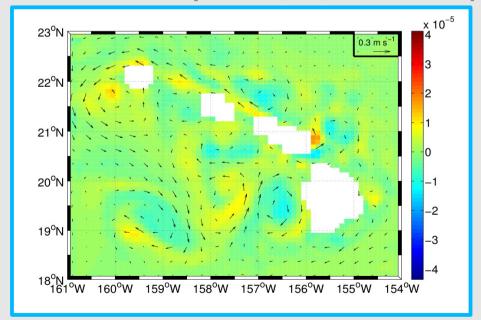
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Eddies too large

Relative vorticity field (s⁻¹) at 20m depth



Relative importance of the forcings

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Run-C	QuikSCAT	none	Cd
Run-D	QuikSCAT	Ugeo	none

Inadequacies:

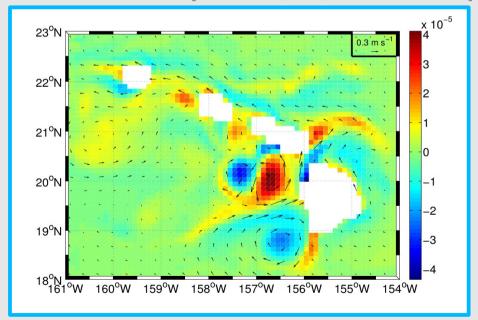
Run-A & Run-B (No wind forcing)
Ocean circulation not realistic
Intensity of the eddies is too low

Run-C (No advection)

Stationary eddies- No expected westward drift

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Eddies too large

Relative vorticity field (s⁻¹) at 20m depth



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<u>Inadequacies:</u>

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Ocean circulation not realistic Intensity of the eddies is too low

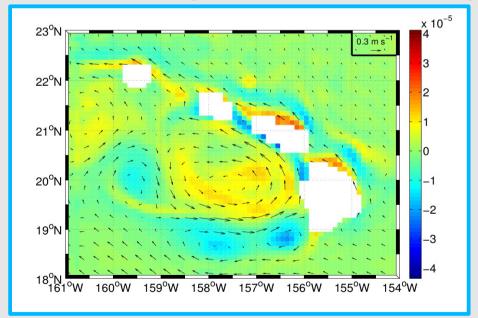
Run-C (No advection)

Stationary eddies- No expected westward drift

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Eddies too large

Relative vorticity field (s⁻¹) at 20m depth



Relative importance of the forcings

			
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Run-C	QuikSCAT	none	Cd
Run-D	QuikSCAT	Ugeo	none

<u>Inadequacies:</u>

Run-A & Run-B (No wind forcing)

Ocean circulation not realistic Intensity of the eddies is too low

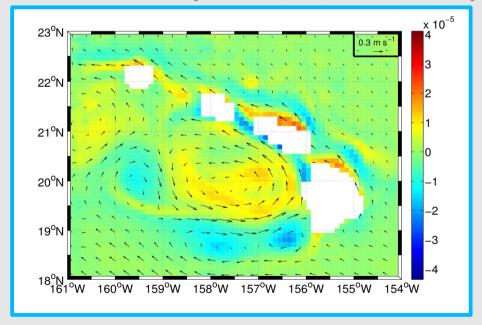
Run-C (No advection)

Stationary eddies- No expected westward drift

Run-D (No Drag. Coeff)

Eddies too large

Relative vorticity field (s⁻¹) at 20m depth





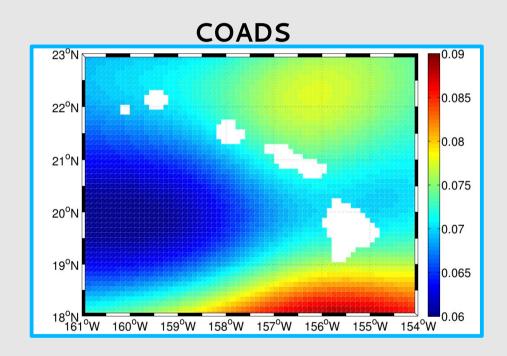
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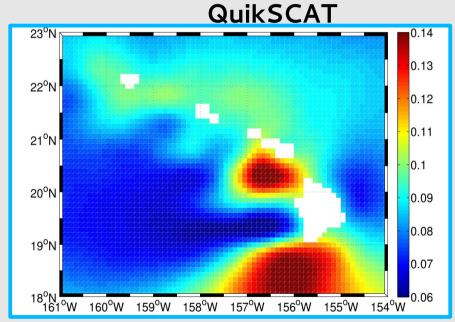
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Resolution of wind forcing

	Wind Forcing	Advection	Drag coeff.
Run-E	COADS (1/2°)	Ugeo	Cd
Run-F	QuikSCAT (1/4°)	Ugeo	Cd

Performed numerical experiments



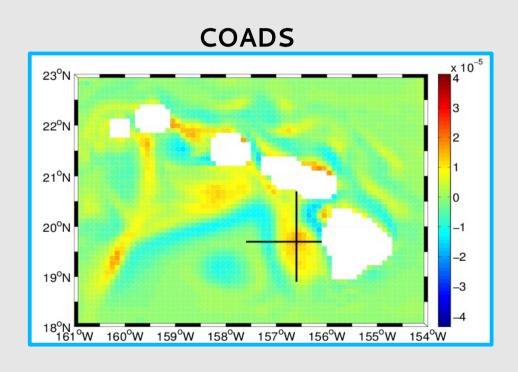


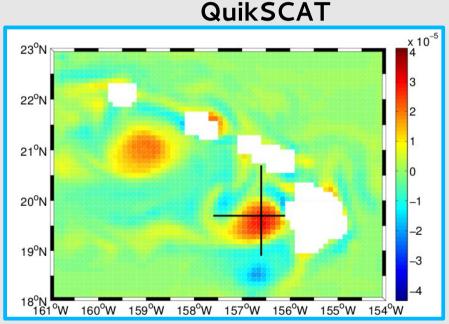
Annually-averaged wind stress values [Nm⁻²]

Resolution of wind forcing

	Wind Forcing	Advection	Drag coeff.
Run-E	COADS (1/2°)	Ugeo	Cd
Run-F	QuikSCAT (1/4°)	Ugeo	Cd

Performed numerical experiments

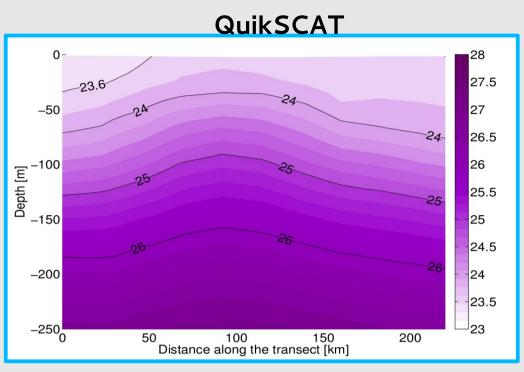


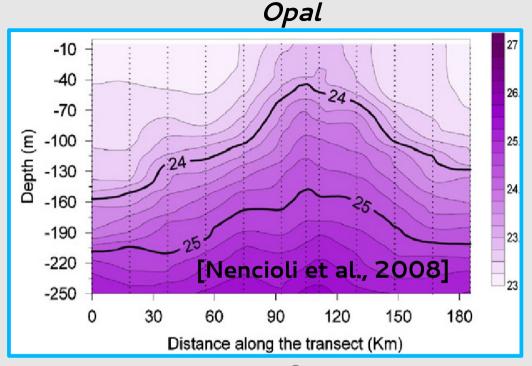


Relative vorticity field (s-1)

Ocean circulation realistic

- at 20m depth
- Focus on numerical cyclones chosen because they are spatially and temporally representative of cyclone Opal studied during the E-FLUX III³⁴





Transect of density (kg m⁻³) + isopycnes

	COADS	QuikSCAT	Opal
Isopycnal outcrop.	σ _{t23.6/23.8}	σ _{t23.6/23.8}	σ _{t23.6}
Depth impact (m)	130±70	>250	>250
Diameter (km)	180±20	180±30	180-200
Velocity (m s ⁻¹)	0.3±0.05	0.53±0.09	0.6

Good representation of mesoscale features in the simulation forced by QuikSCAT



Importance of wind-forcing spatial resolution

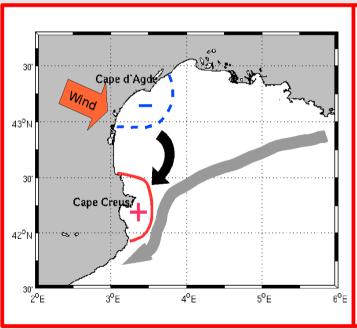
Hypothesis of generation

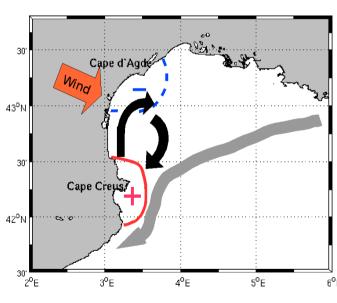
Mesoscale eddies in the western part of the GoL

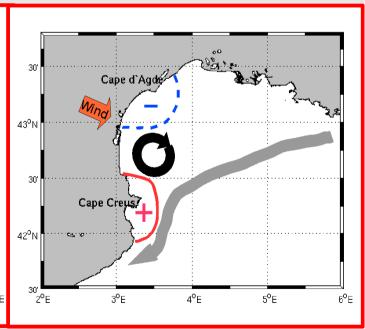
2001-2008 [Hu et al., 2011]

2 conditions : → Persistent & strong northwest wind

→ Strong stratification





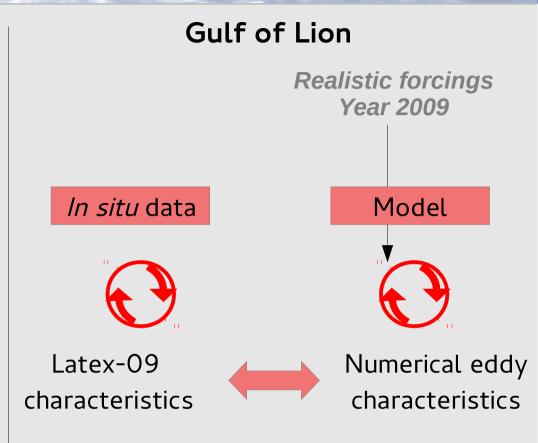


Strategies

Hawaiian archipelago Climato. forcings Wind Topo Currents In situ data Model Opal Characteristics Numerical eddies characteristics

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What is the **relative importance**of the forcings on the **generation of mesoscale eddies?**

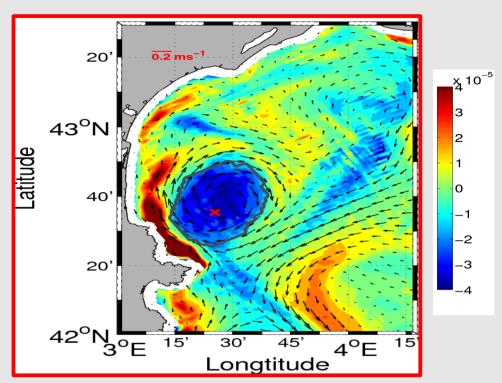


Find a similar numerical eddy at the same period

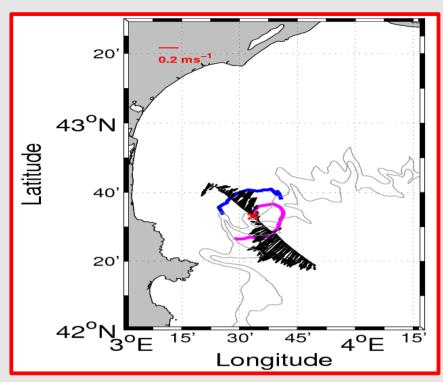
What is the **generation process** of the Latex-09 eddy?

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Part III - Forcings and generation



Eddy detected by wavelet analysis [Doglioli et al., 2007] – Relative vorticity [s⁻¹] 15m depth August 27

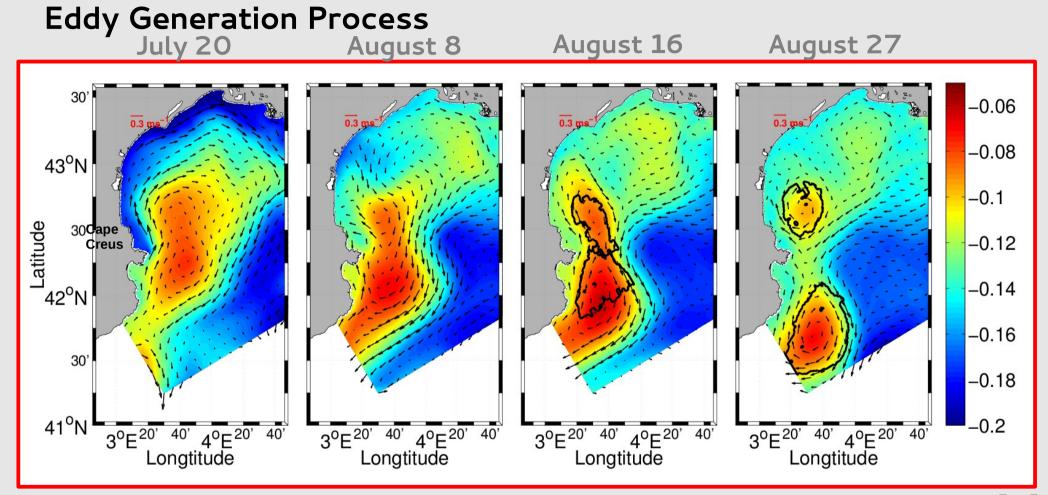


Latex09 ADCP data August 27 +Buoys from August 26-29

	Model	Data
Center	3°26'E – 42°36'N	3°34'E – 42°33'N
Depth impact (m)	37	35
Diameter (km)	28.6±1.4	22.7±1.2



Part III - Forcings and generation



Eddies detected by wavelet analysis - Sea Surface Height [m]

- Pushing and squeezing of an anticyclonic circulation between a meander of the NC and the coast
- Separation in two structures

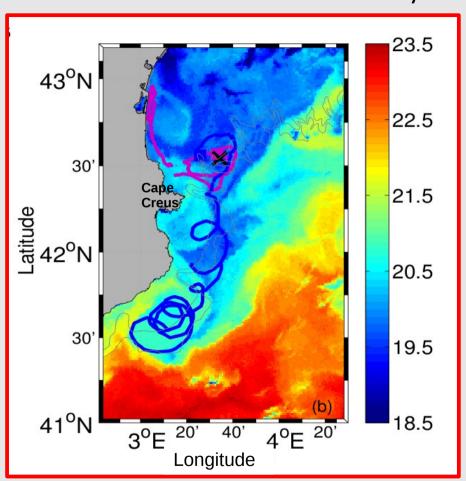
New Generation Process of the Latex-09 eddy

[Kersalé et al., 2013]

Part III - Forcings and generation

Post Generation dynamics

Latex09 feeds the Catalan eddy



<u>x</u>10 43^oN 3 30' Cape Creus Latitude N_o24 0 -2 30' -3 4°F 20' 3°F 20' Longitude

SST (°C) September 12 +Buoys from August 26- September 12

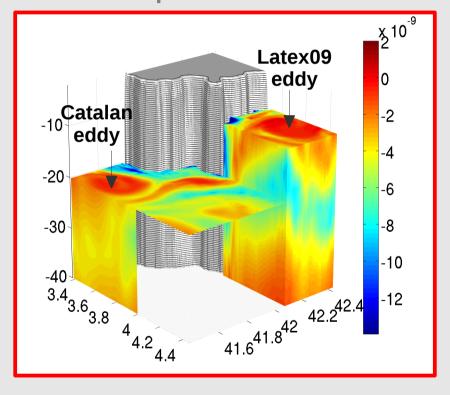
Relative vorticity [s⁻¹] 20m depth September 3

Drifter trajectories explained by the generation of a submesoscale transient structure

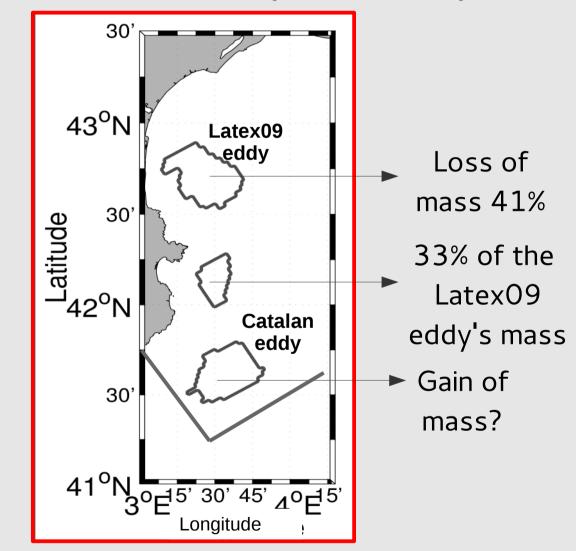
Part III – Forcings and generation

Latex09 - Loss of mass

Potential vorticity [kg.m⁻⁴.s⁻¹] in 3D on September 3



Eddies detected by wavelet analysis



Interactions between the two eddies lead to a transfer of mass and vorticity from the GoL to the Catalan shelf

Conclusions - Part III

- Cumulative effects of external forcings
- Importance of wind forcing resolution

Synergy of model results and in situ data

- New generation process of eddies in the western part of the Gulf of lion
- Observation of transient submesoscale structure
 - Exchanges between coastal areas

Plan

Part I – Tools for study mesoscale features

Numerical codes
In situ experiments

Strategies

Part II - Characteristics of mesoscale eddies

Eddy center detection

Horizontal characteristics

Vertical characteristics

Part III – Forcings and generation

Forcing sensitivity – Wind sensitivity

Process of generation

Generation of submesoscale structures

Impact on coast-offshore exchanges

Part IV – Study of turbulent mixing

In situ experiment

Calculation of diffusion coefficients

Lagrangian experiment

Conclusion & perspectives

Mesoscale activity

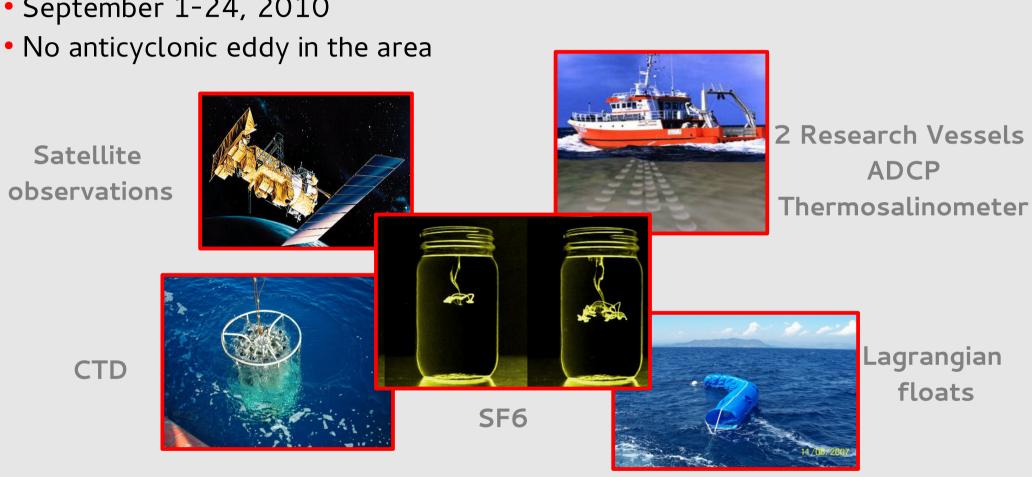
Submesoscale

Turbulence

43

In situ experiment – Field campaign Latex10

- Western part of the GoL
- September 1-24, 2010

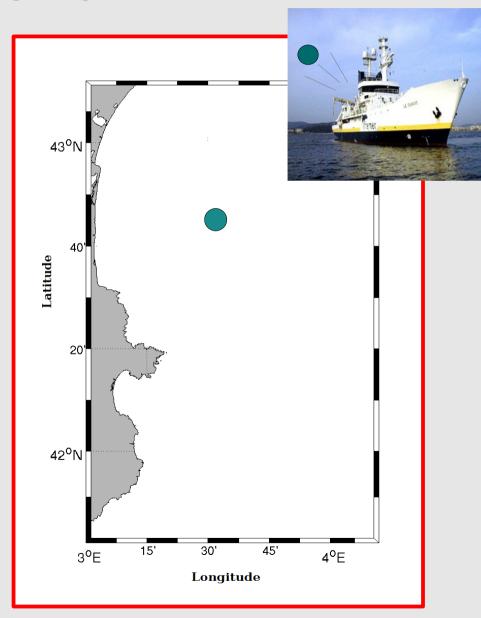


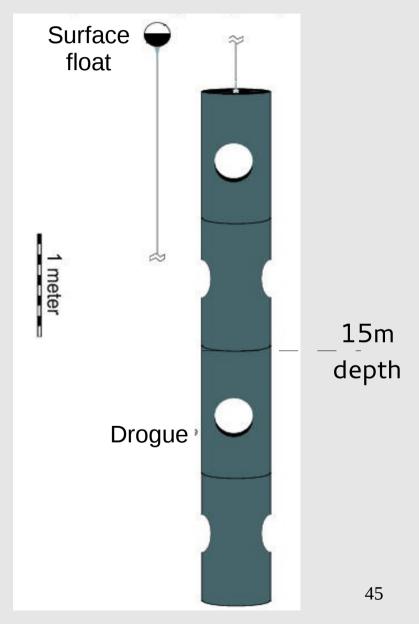
Further characterization of the turbulent mixing in a coastal environment

44

In situ experiment – Field campaign Latex10

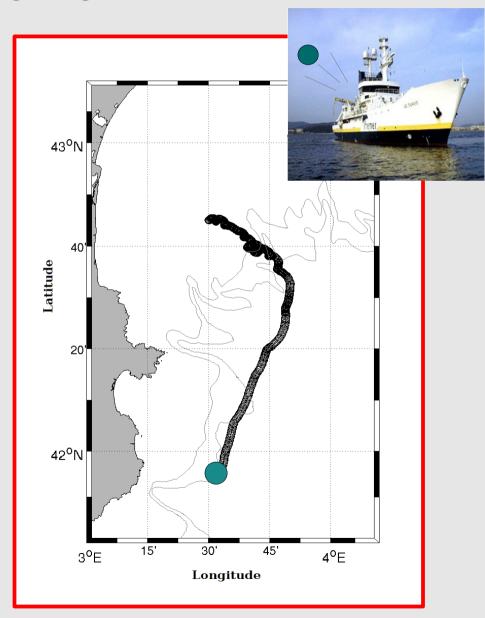
Lagrangian float

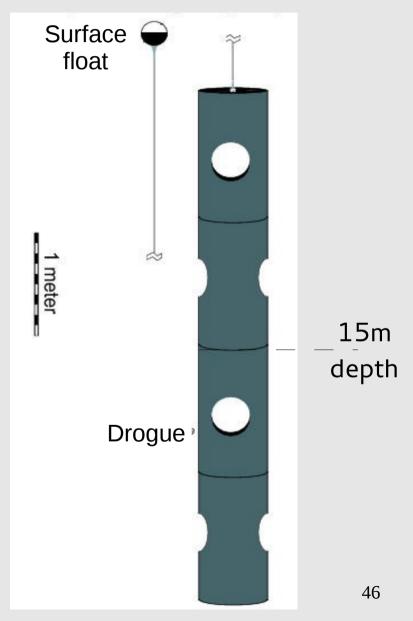




In situ experiment – Field campaign Latex10

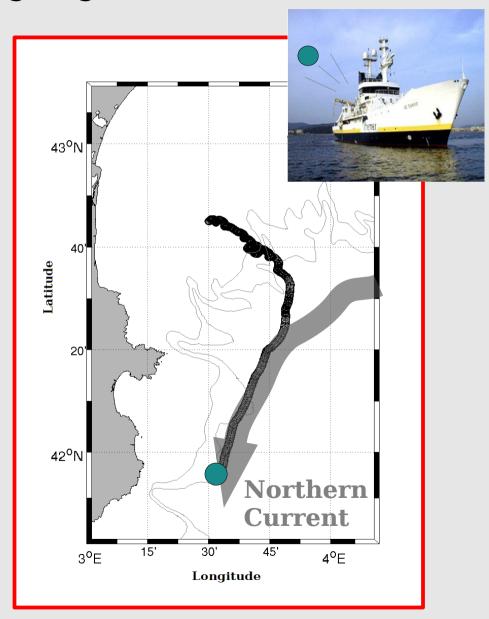
Lagrangian float

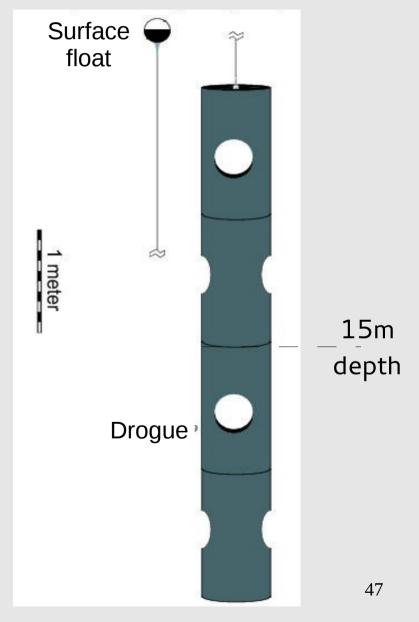




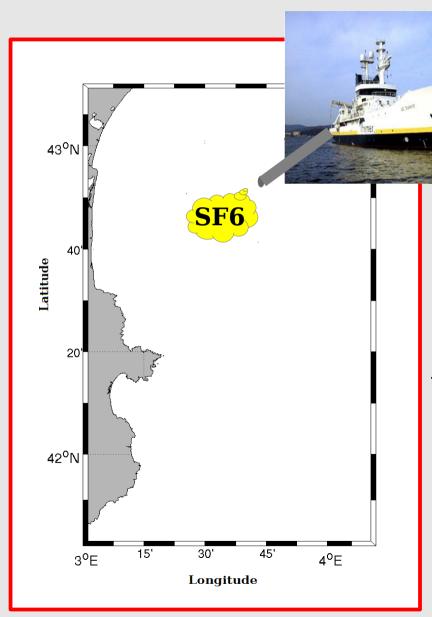
In situ experiment – Field campaign Latex10

Lagrangian float





In situ experiment – Field campaign Latex10 SF6 – Sulfur Hexafluoride





Artificial patch:

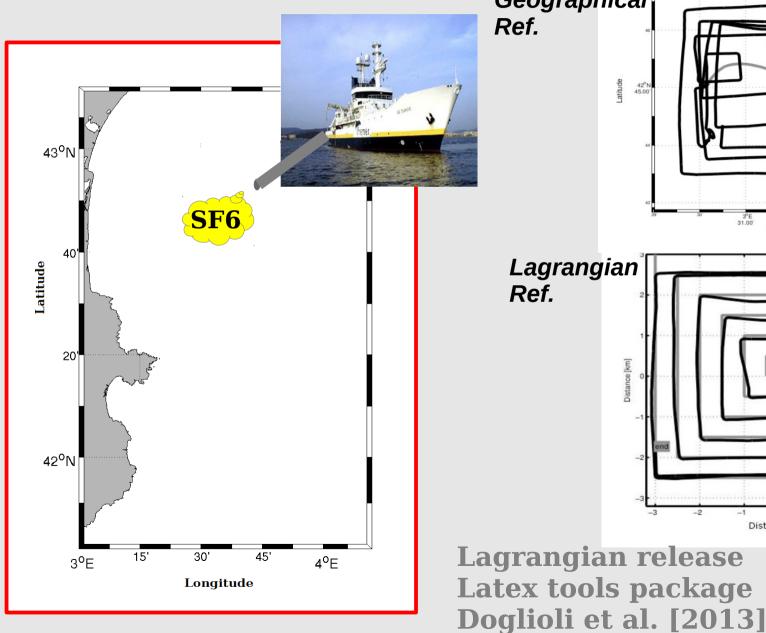
Release of seawater satured with SF6 Inert tracer

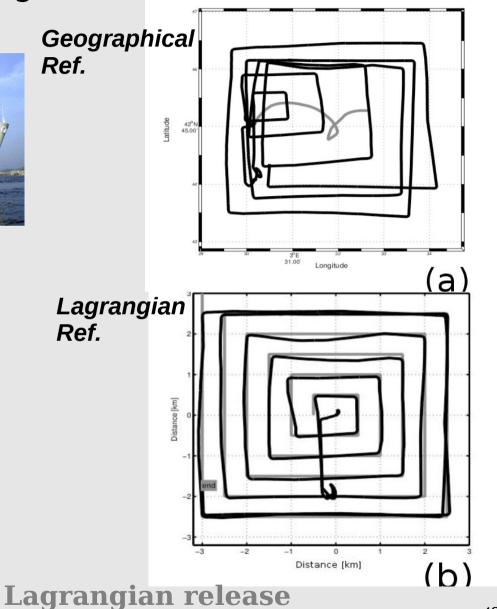
7500 L steel tank of seawater

- + Injection of pure SF6 gaz
- = Satured SF6 solution

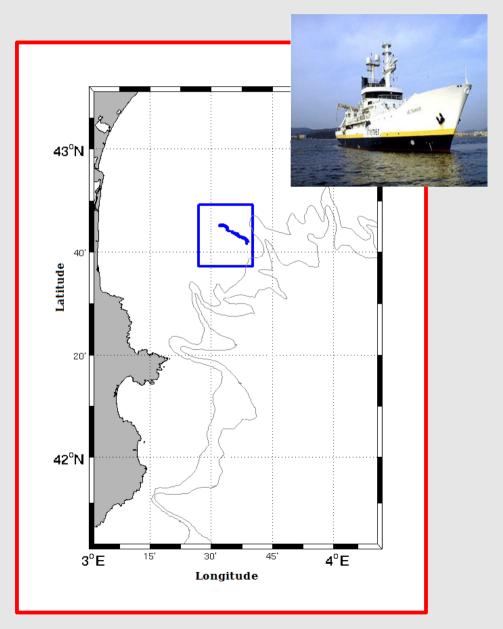
In situ experiment - Field campaign Latex10

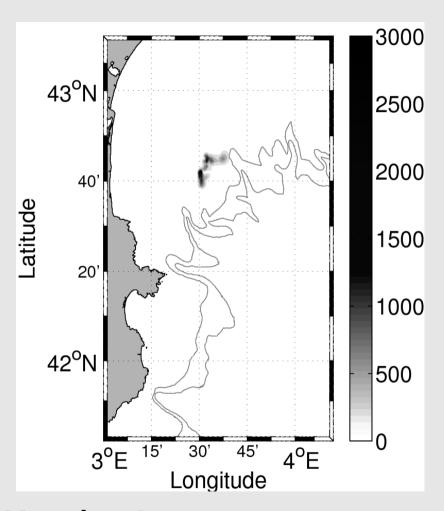
SF6 - Sulfur Hexafluoride



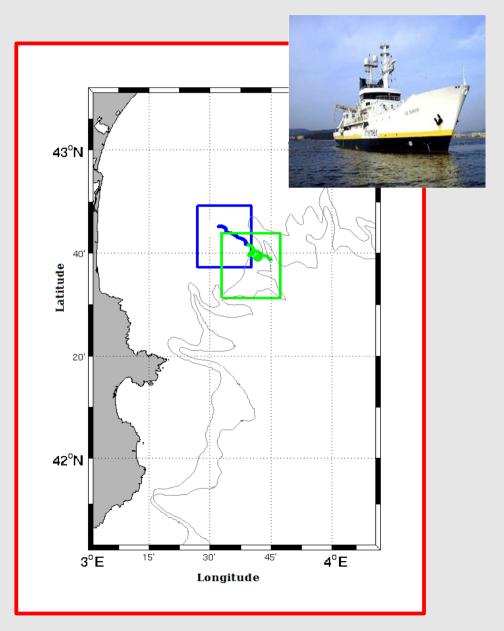


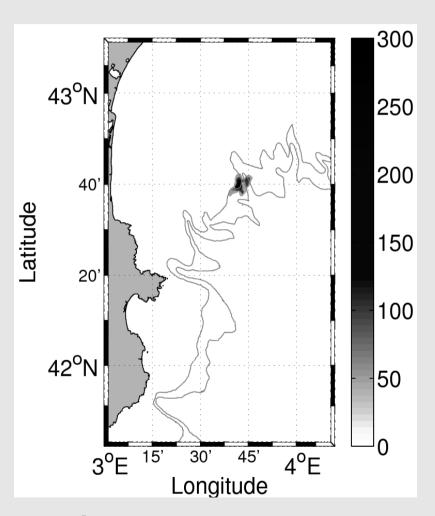
49



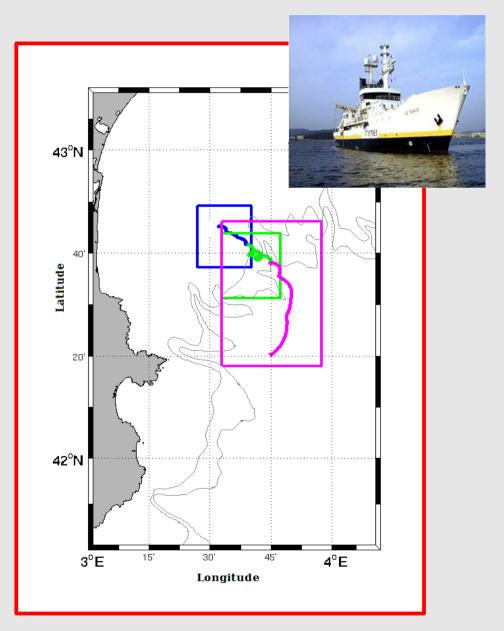


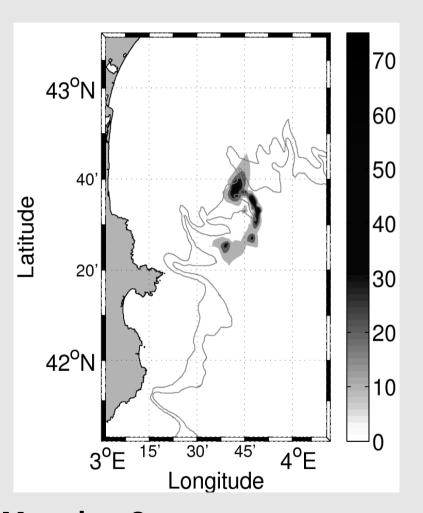
Mapping 1
Concentration of SF6 [fmol I⁻¹] ⁵⁰
1 day after the release



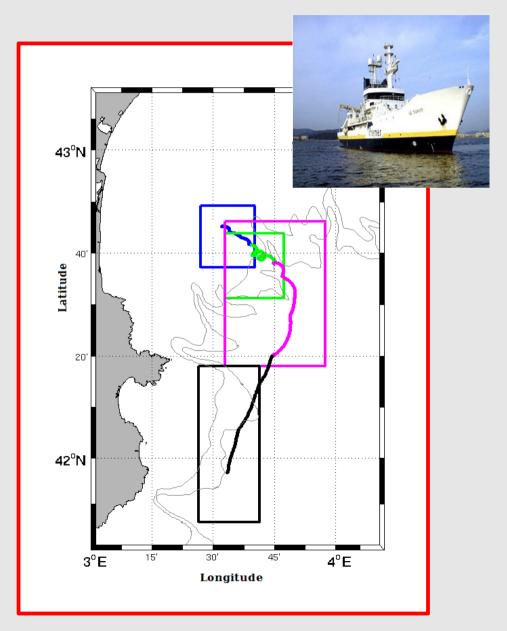


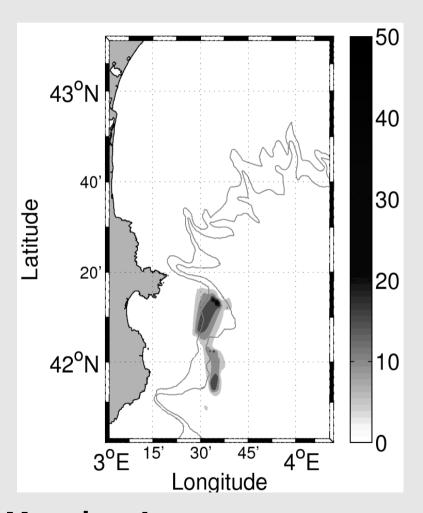
Mapping 2
Concentration of SF6 [fmol I⁻¹] ⁵¹
2 days after the release





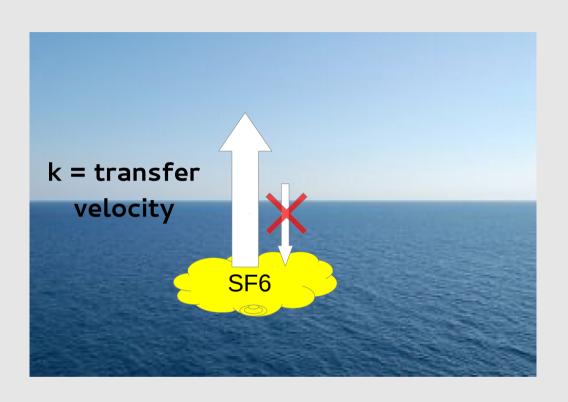
Mapping 3
Concentration of SF6 [fmol I⁻¹] 52
5 days after the release





Mapping 4
Concentration of SF6 [fmol I⁻¹] ⁵³
7 days after the release

Atmospheric loss



- No external source
 Patch is supersaturated with
 respect to the atmosphere
- Loss of SF6 at the surface of the water column due to atmospheric loss

 Variation of SF6 with time due to atmospheric loss modeled as a negative exponential function:

$$C(t) = C_0 e^{\frac{-k}{z_{mix}}t}$$

 C_0 : released concentration SF6

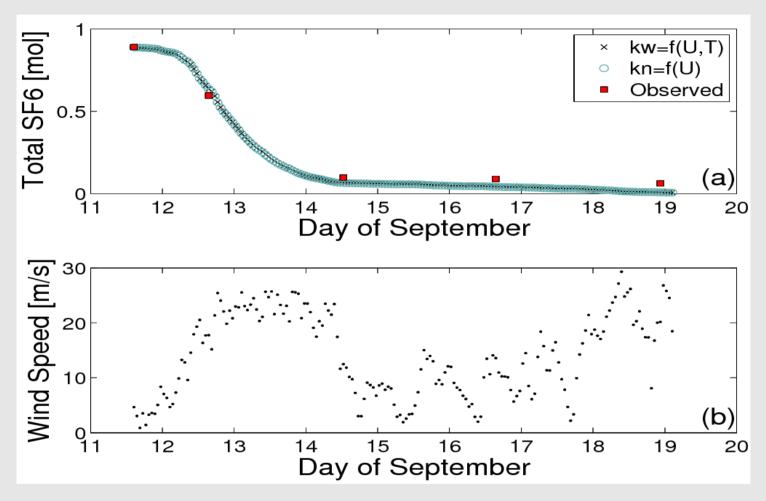
 \mathbf{Z}_{mix} : Depth mixed layer

k: 2 parameterizations for the

transfer velocity

54

Atmospheric loss

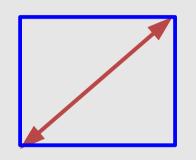


- 2 modeled curves superimposed => Importance of the turbulent diffusivity related to the wind speed
- Good agreement of the observed quantities of SF6



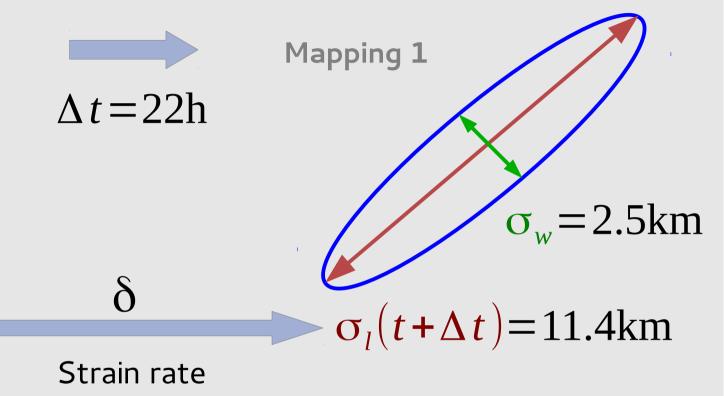
Calculation of the horizontal diffusion coefficient





$$\sigma_{l}(t) = 7.5 \text{km}$$

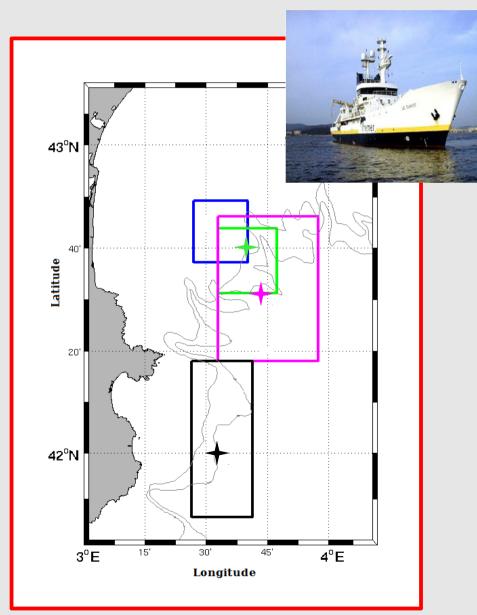
$$\delta = \frac{\ln \frac{\sigma_l(t)}{\sigma_l(t + \Delta t)}}{\Delta t}$$



$$K_h = \sigma_w^2 \delta = 7.6 \, m^2 s^{-1}$$

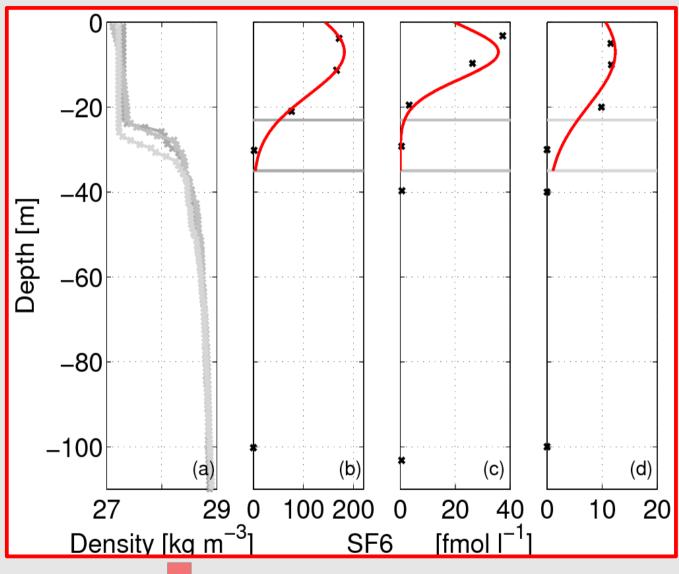
Determination for horizontal scale < 2.5 km for temporal scale < 2 days 56

Calculation of the vertical diffusion coefficient CTD – Niskin bottles used to sampled SF6



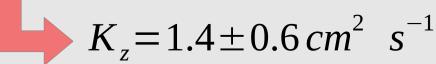


Calculation of the vertical diffusion coefficient

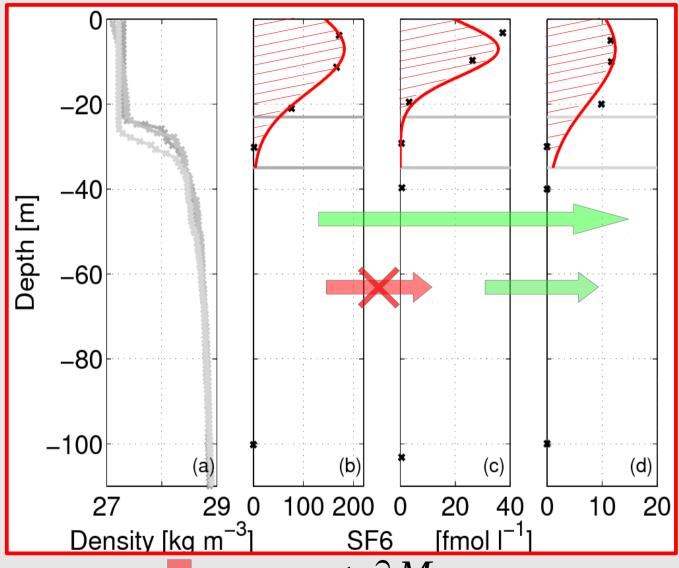


- Gaussian shape for the tracer distribution with depth
- Evolution of the variance of the Gaussian

$$K_z = \frac{\sigma^2}{2\Lambda t}$$





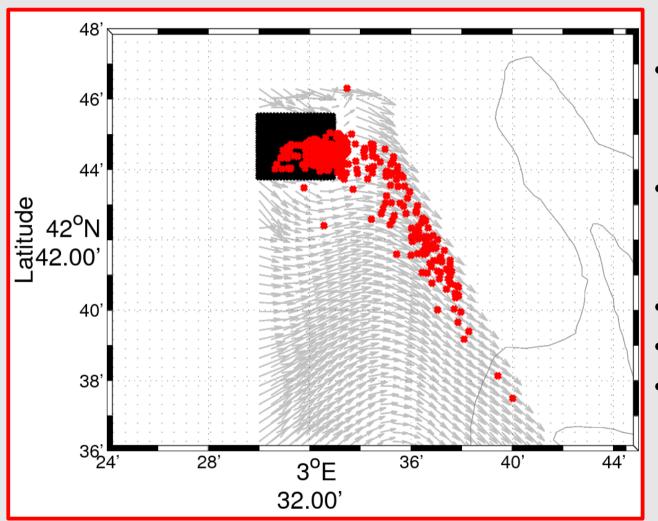


- Gaussian shape for the tracer distribution with depth
- Increase of the second moment M2 of the SF6 profiles with time

$$M_2 = \frac{\int C(z - z_0)^2 \partial z}{\int C \partial z}$$

$$K_z = \frac{1}{2} \frac{\partial M_2}{\partial t} = 1.2 \pm 0.7 \, cm^2 \, s^{-1}$$

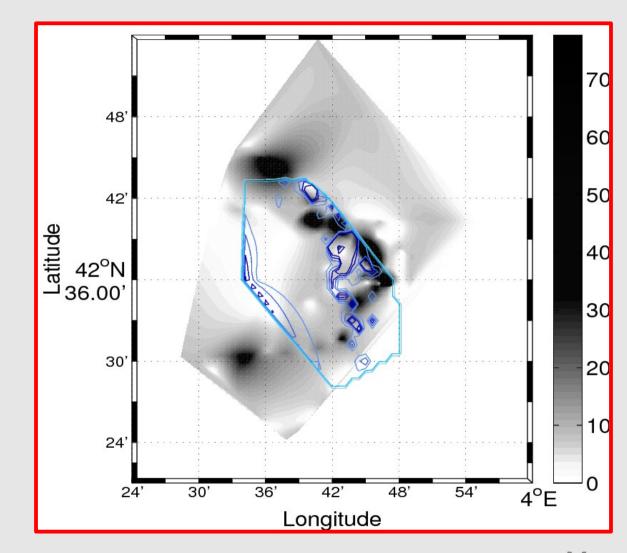
Lagrangian experiment



- Tracer release experiments simulated using a numerical model
- Test estimates of horizontal diffusivities
- Motion of 10 000 particles
- Classic random walk model
- Input velocity field: interpolated ADCP velocities

- Initial position of the particles (end of the release)
- Final position of the particles (end of mapping 1)
- Input velocity field

Lagrangian experiment



- Good agreement between the positions of in situ and numerical patches
- Comparison of the total patch areas

Numeric: 273 km²

In situ: 381 km²

Gray scales [SF6]_{in situ} – Contour [SF6]_{num.} Mapping 3

Conclusions - Part IV

- Investigation of the **turbulent mixing** in a **coastal** environment in the western part of the Gulf of Lion
 - Quantification & validation of the gas exchange
 from the ocean to the atmosphere
 - Estimation of turbulent mixing coefficients
 - Test of the horizontal diffusivity coefficient
 in Lagrangian experiments

MAIN CONCLUSIONS

- Comprehension of mesoscale dynamics on two areas
- Cumulative effects of external forcings Importance of wind resolution
- Local characterization & analysis of generation processes of mesoscale features
- Exchanges between coastal areas : Observation of transient submesoscale structure
- Synergy of model results and in situ data
- Estimation of turbulent mixing coefficients in a coastal environment

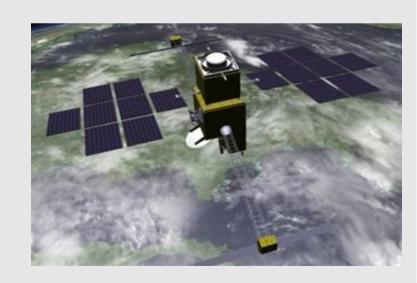
PERSPECTIVES

Short term perspectives

- Tracers distribution & connectivity between regions
- Quantification of coast offshore exchanges with Lagrangian tools
- New regional altimetry products => Precision on coastal areas

Long term perspectives

- Use of numerical field output Symphonie test and validate the SWOT data (GoLSWOT)
- Latex campaign strategies applied in already scheduled campaigns: GoLSWOT and P2B2M



Local & regional impact on ecosystems

THANKS

