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In-situ estimate of submesoscale horizontal eddy diffusion coefficients across a front

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1. Intro

Submesoscale structures





(from http://oceanservice.noaa.gov/)



⁽from Dickey et al., J. Mar. Syst, 2003)

Fronts, jets and eddies: Horizontal scales ~(100m - 10km) Time scales ~(days - week)

Key role for : Energy transfer Horizontal and vertical transport Biogeochemical cycles

Submesoscale structures



Numerical model studies
 Favored by:
 Advances in computational power
 Development of regional models



Submesoscale structures



Numerical model studies Favored by: Advances in computational power Development of regional models

In-situ observations Challenging due to small spatial and temporal scales



Limited estimates of key physical parameters

Submesoscale structures



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In-situ observations Challenging due to small spatial and temporal scales



Limited estimates of key physical parameters

Focus of this study

New approach to estimate horizontal eddy diffusion coefficients (Kh) from in-situ observations

2. Latex10

Lagrangian Transport EXperiment



Latex10 campaign (September 1-24, 2010)



- Western part of Gulf of Lion
 (NW Mediterranean)
- <u>Main goals</u> : (Sub)mesoscale dynamics Cross-shelf exchanges
- Adaptive sampling strategy based on :
 Satellite data
 - Ship-based current measurements
 - Iterative Lagrangian drifter releases





Example:

→ Drifter array deployment "Lyap01" (Sep. 15, 2010)



⁽from Nencioli et al., GRL, 2011)

- Identified in-situ Lagrangian Coherent Structures (LCSs)
- Evidenced inaccuracy of altimetry in coastal region
- LCS associated with a frontal structure

2. Latex10

Front development



→ AVHRR SST + 3-day drifter trajectories (Sep. 8 to 15, 2010)



2. Latex10

Front development



→ AVHRR SST + 3-day drifter trajectories (Sep. 8 to 15, 2010)



Convergence of warmer (eastern outer shelf) and colder (western inner shelf) water masses



Colder and warmer water masses converging along attractive LCS





Colder and warmer water masses converging along attractive LCS





 Shape of T and S profile across the front results from balance between convergence and horizontal mixing



Colder and warmer water masses converging along attractive LCS





 Shape of T and S profile across the front results from balance between convergence and horizontal mixing

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$

 Analytical solution to 1D advection diffusion equation for a tracer T

⇒ Analogous to Flament et al. 1985, Ledwell et al. 1998 (from satellite)

Advection-diffusion equation



1D equation for a tracer *T*

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$

Advection-diffusion equation



1D equation for a tracer *T*

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$
$$-\gamma(x-\mu)\frac{dT}{dx} = K_H \frac{d^2 T}{dx^2}$$

Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction

Advection-diffusion equation



1D equation for a tracer *T*

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$
$$-\gamma(x-\mu)\frac{dT}{dx} = K_H \frac{d^2 T}{dx^2}$$

with

Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction
 - y : Strain rate (Lyapunov exponent)
 - μ : Position of LCS axis



Advection-diffusion equation

with



1D equation for a tracer *T*

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$
$$-\gamma(x-\mu)\frac{dT}{dx} = K_H \frac{d^2 T}{dx^2}$$

Boundary Conditions

$$T(x = -\infty) = T_1;$$

$$T(x = +\infty) = T_2;$$

Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction
 - y : Strain rate (Lyapunov exponent)
 - μ : Position of LCS axis



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$



$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} \mathrm{d}t$$





































Cross-front transects







SST and SSS from ship thermosalinograph

Cross-front transects







² <u>SST and SSS from ship</u> thermosalinograph



Cross-front transects







SST and SSS from ship thermosalinograph



- SST -

Cross-front transects





4. Kh



² <u>SST and SSS from ship</u> <u>thermosalinograph</u>



September 15, 2013

Cross-front transects





- SST -



² <u>SST and SSS from ship</u> <u>thermosalinograph</u>



Cross-front transects





4. Kh



- ^{38.2} SST and SSS from ship
 ^{38.1} thermosalinograph
 - Identified 30 cross-front transects
- Transects projected on the direction normal to the LCS axis





Example: Transect 11



4. Kh Curve fitting



Example: Transect 11

Front equation

$$T(x) = C_1 + C_2 \operatorname{erf} (C_3 (x - C_4))$$









Strain rate





Dispersion patterns of drifter arrays

Strain rate







Dispersion patterns of drifter arrays

 For each deployment, computed fastest separation rate between buoy couples (analogous to Lyapunov exponent)

Strain rate





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Dispersion patterns of drifter arrays

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Eddy diffusivity coefficients



<u>T Front</u>





S Front

Strain rate



5. Results

Eddy diffusivity coefficients

5. Results





Eddy diffusivity coefficients

5. Results







New approach relatively simple and cheap (i.e. compared to passive tracer release experiments)

→ In-situ estimates of Kh at the submesoscale in line with values used in high-resolution numerical models

Tail of high values affects Kh statistics (mean and standard deviation) => check starting assumptions:

- Steady state
- Uniform strain rate
- Vertical motions
- •



- Further dedicated in-situ experiments
- Test approach from high-resolution models
- Extend analysis of Kh over wider regions/the global ocean using remote sensed datasets

5. Results **Perspectives**



- Further dedicated in-situ experiments
- Test approach from high-resolution models
- Extend analysis of Kh over wider regions/the global ocean using remote sensed datasets



Airborne SWOT sensor tested over key ocean regions (2013-2020) Surface Water and Ocean Topography NASA – CNES mission

- New generation, high-resolution (1Km) altimeter
- Launch: Fall 2020







This work has been developed within the project:



Lyapunov Analysis in the CoaSTal Environment (LACOSTE)

Marie Curie Intra-European Fellowship Call: FP7 - PEOPLE - 2011 - IEF Leading PI: F. Nencioli



F. Nencioli, F. d'Ovidio, A. Doglioli, A. Petrenko Surface coastal circulation patterns by in-situ detection of Lagrangian Coherent Structures. Geophysical Research Letters, 38, L17604, doi:10.1029/2011GL048815

LATEX website: www.com.univ-mrs.fr/LOPB/LATEX



Numerical solution



First order upwind scheme





Numerical solution



First order upwind scheme





- Fast adjustment to equilibrium (within 1 day)
- However Kh proportional to square of width
- Even small errors in width could affect estimate of Kh