

Modellizzazione della dispersione in mare.

*Introduzione ai modelli numerici ed esempi di applicazioni
nelle acque costiere liguri*

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Maître de Conférences

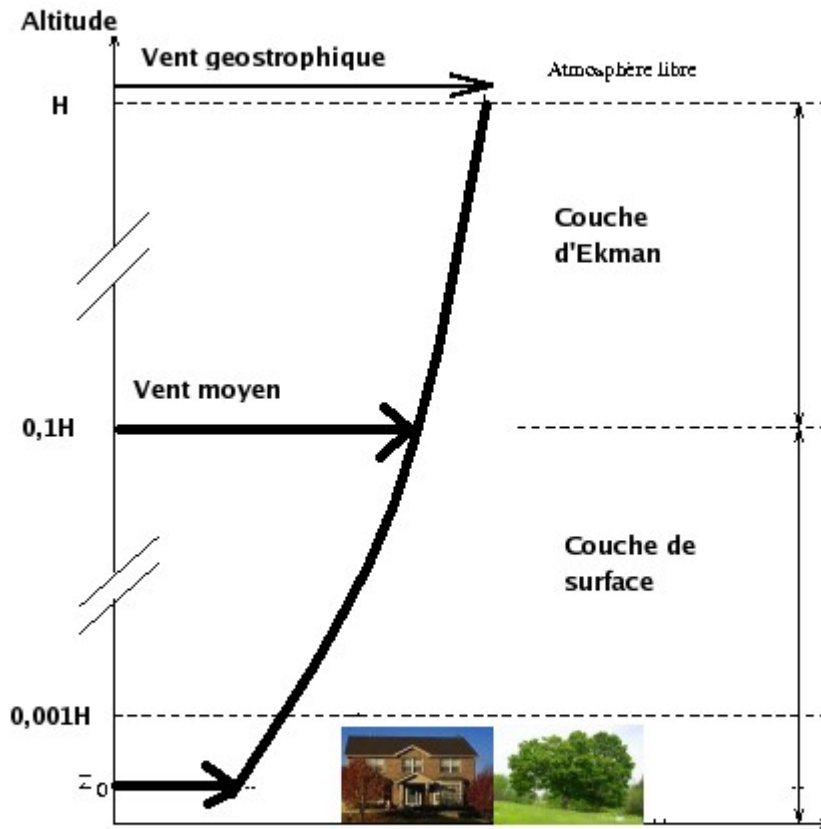
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In atmosfera... (ripasso)

COUCHE LIMITE ATMOSPHERIQUE



H = HAUTEUR DE LA COUCHE

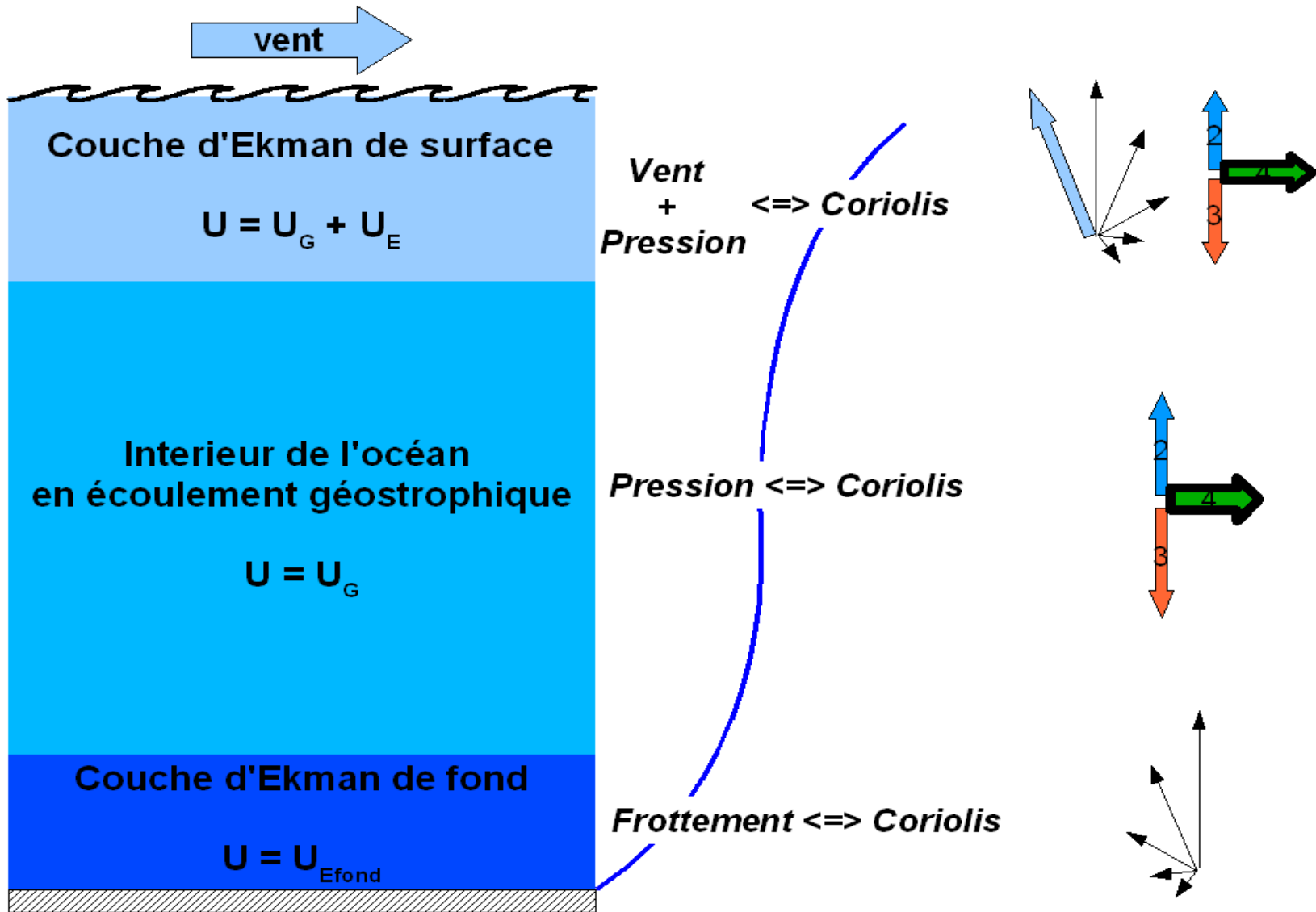
http://fr.wikipedia.org/wiki/Couche_limite

En météorologie, on appelle couche limite planétaire la zone de l'atmosphère entre la surface (terre ou mer), où la friction ralentit le déplacement de l'air, et l'atmosphère libre où cette dernière devient négligeable. Elle varie entre 0,5 et 3 km d'épaisseur selon la stabilité de l'air et la rugosité de la surface. Elle est en moyenne de 1 500 mètres. L'étude théorique de cette tranche d'atmosphère divise en fait la couche limite planétaire comme la superposition de deux couches dont les épaisseurs sont très inégales:

La **couche d'Ekman** dans laquelle le vent est causé par un équilibre entre le gradient de pression, la force de Coriolis, due à la rotation quotidienne de la Terre, et une portion de la friction diminuant graduellement jusqu'à l'atmosphère libre. La vitesse et la direction au sommet de cette couche est approximativement celle du vent géostrophique alors qu'elle diminue graduellement et tourne vers la plus basse pression à mesure qu'on descend vers le sol

La **couche de surface** ou **couche limite de turbulence atmosphérique** immédiatement au contact du sol et dont l'épaisseur ne dépasse pas le dixième de celle de l'ensemble de la couche limite. La vitesse de l'air y est causée par la convection due aux différences de températures et par les effets dynamiques du reliefs. Le flux y est turbulent. On parle également d'une sous-couche rugueuse tout près de la surface, qui varie de quelques centimètres à quelques dizaines de mètres selon les aspérités du relief. La vitesse y tend vers zéro.

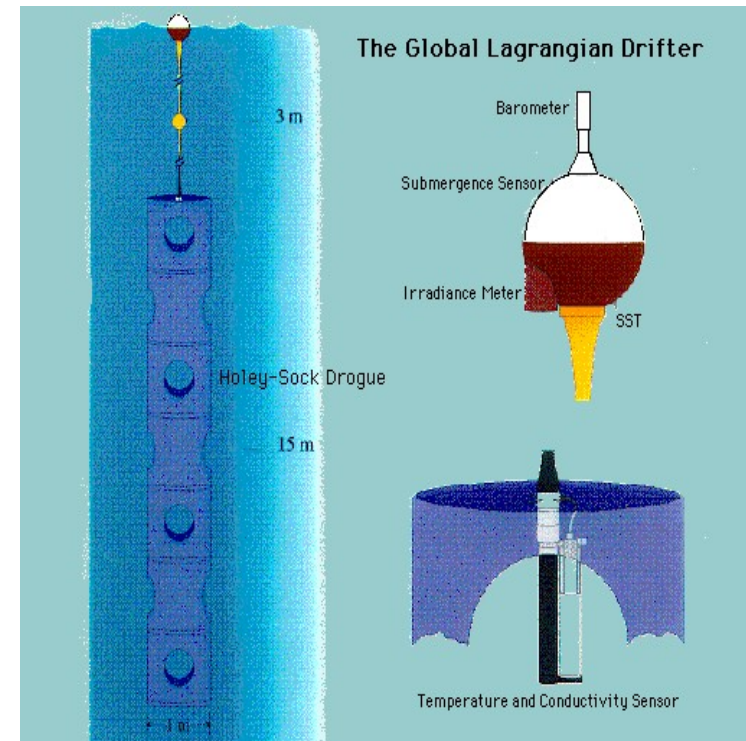
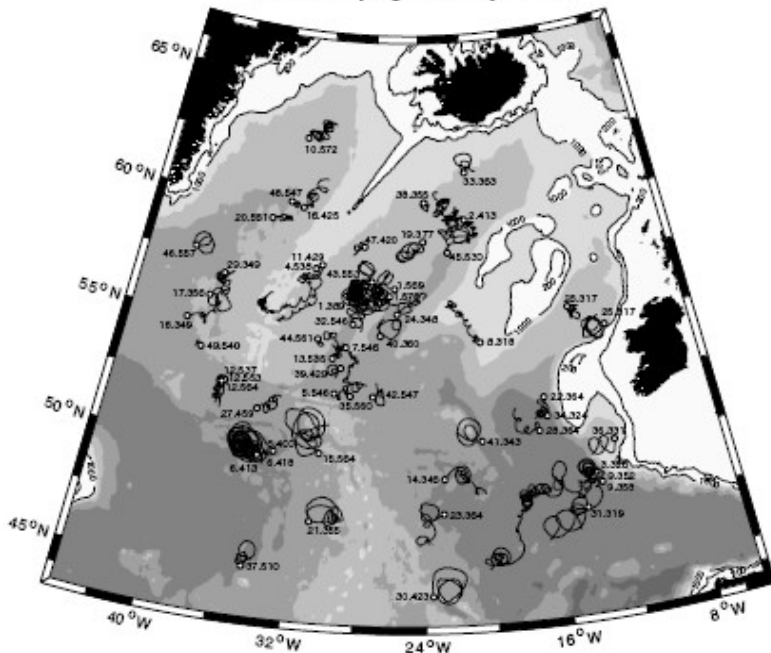
In mare

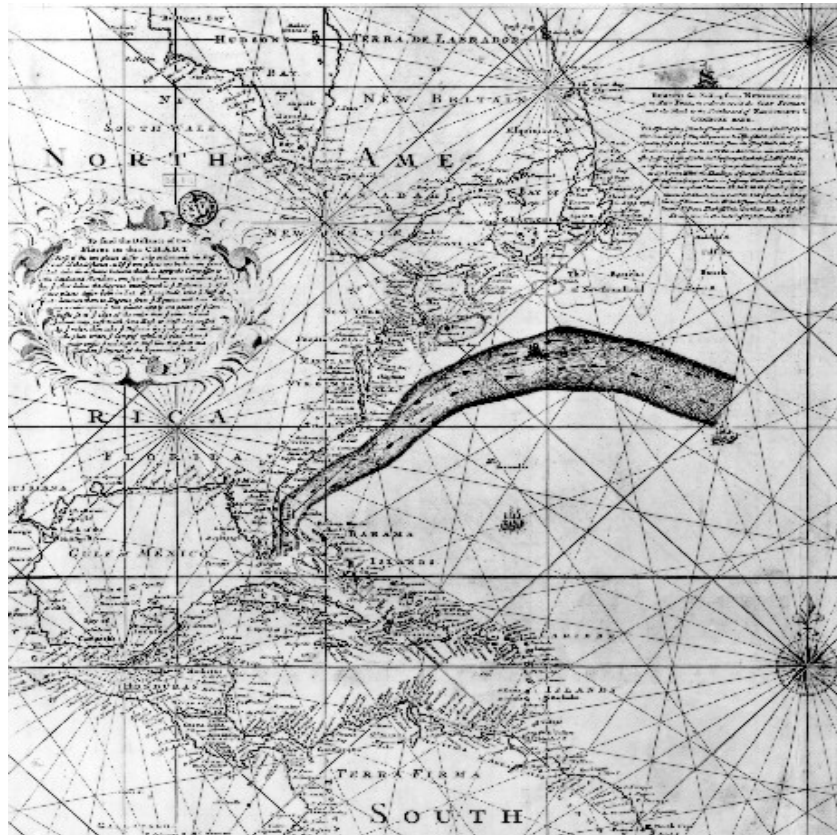


Even though systematic observations began in the 1880s with pioneering observations by Nansen and others, the seagoing and theoretical efforts were mainly oriented toward describing large-scale circulation, which was often regarded as steady for lack of more detailed information. It was not until the 1960s, when long-distance tracking of drifting buoys at mid-depth showed currents to be highly variable on quite small spatial scales (5), that oceanographers became aware of the immensity of their task.



Floats Looping in Anticyclones





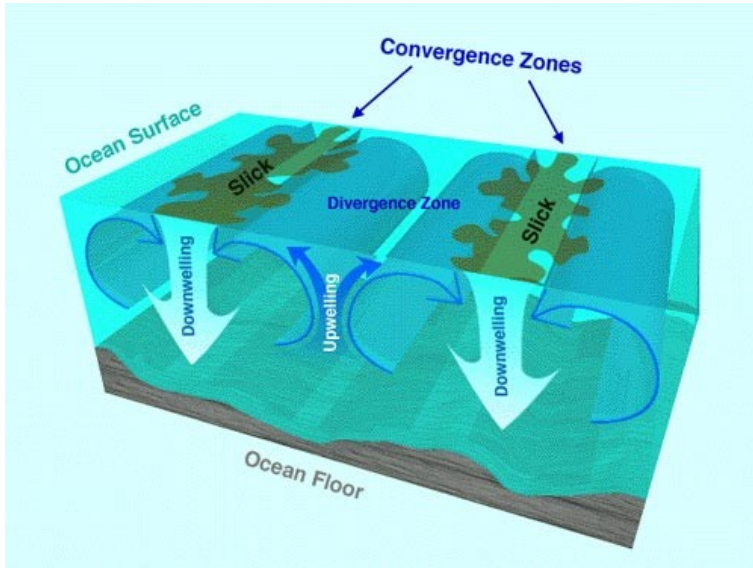
The Franklin-Folger map of the Gulf Stream, printed in 1769-1770.

This early map of the Gulf Stream location was produced by B. Franklin for the mail service from England, based on information from whaling captain Timothy Folger. This map was rediscovered by P. Richardson (1980), and is remarkably accurate. This image is from R. Peterson et al. (1996), article in Progress in Oceanography.

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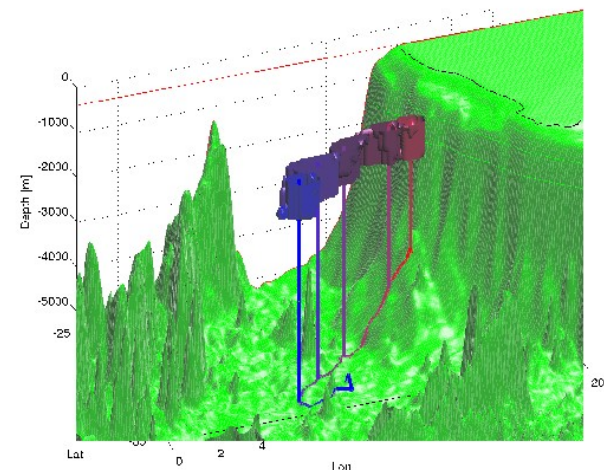
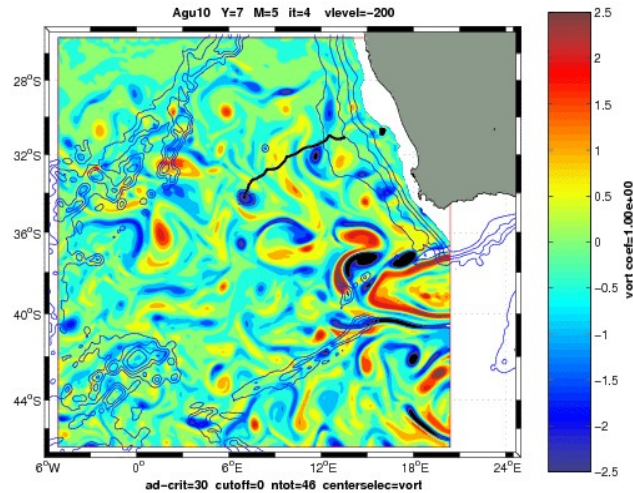
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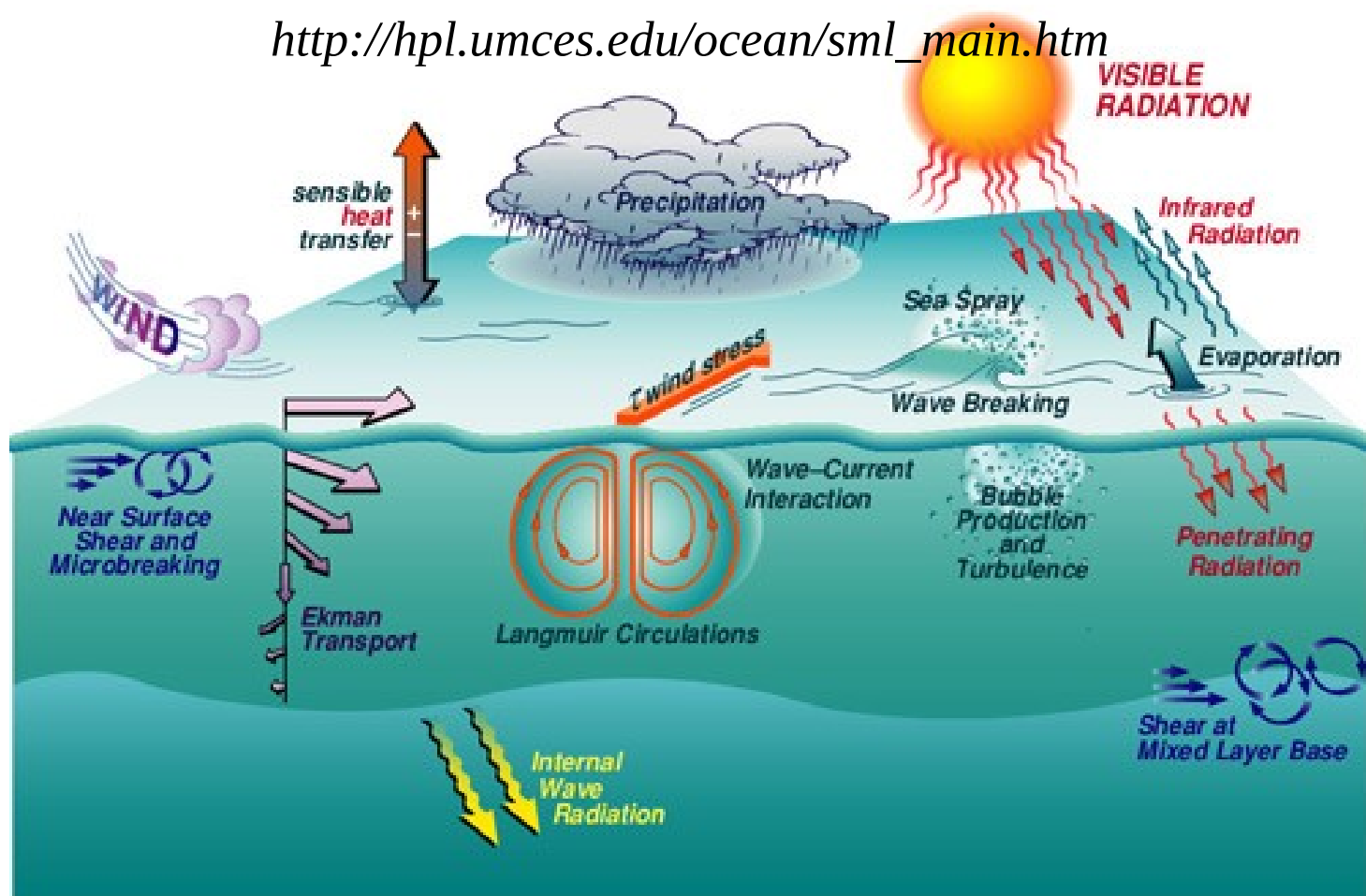
Strato limite oceanico di superficie: *Cellule di Langmuir*



http://sealevel2.jpl.nasa.gov/jr_oceanographer/oceanographer-dgiacomo.html

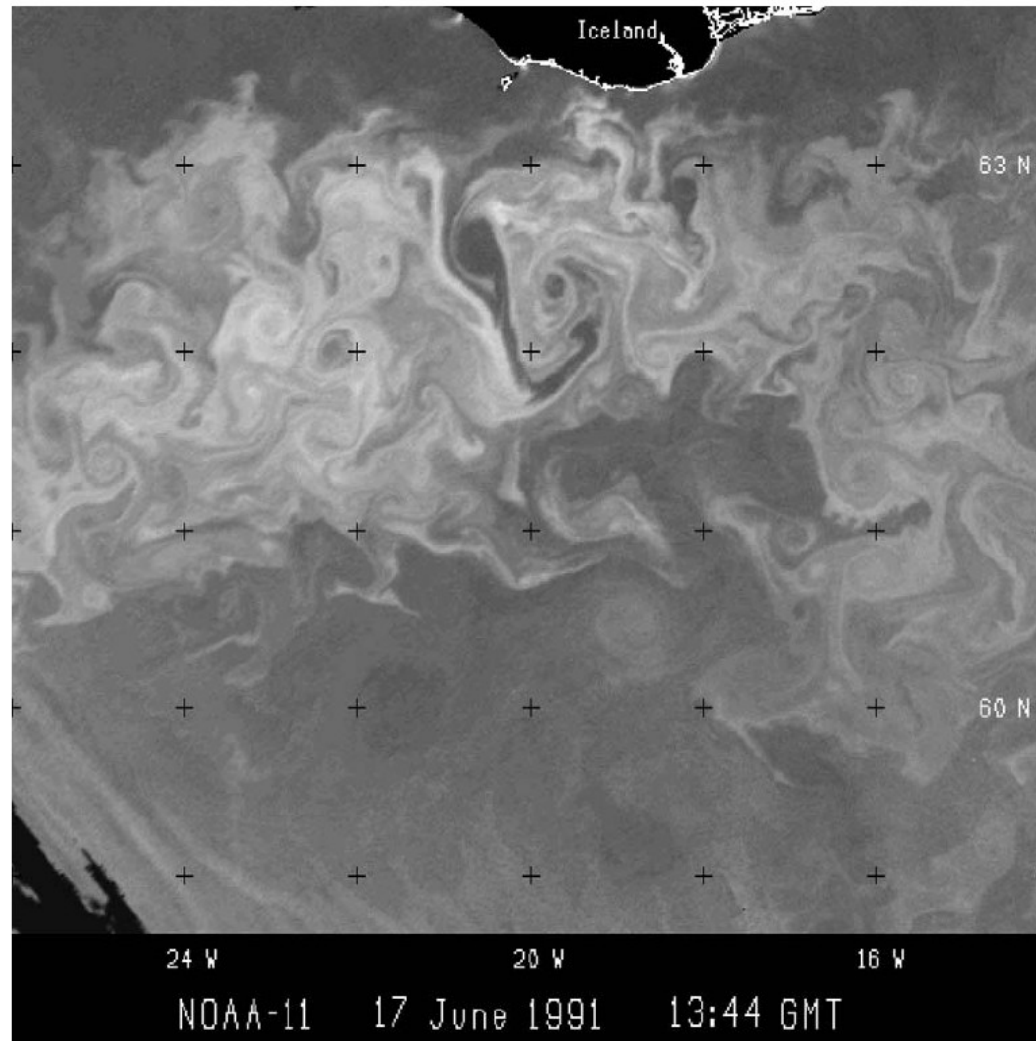
Strutture di mesoscala: *Agulhas Rings*





The upper region of the ocean typically exhibits of a surface mixed layer with a thickness of a few to several hundreds meters. This mixed layer is a key component in studies of climate, biological productivity and marine pollution. It is the link between the atmosphere and deep ocean and directly affects the air-sea exchange of heat, momentum and gases. Moreover, turbulent flows in the mixed layer affect biological productivity by controlling both the supply of nutrients to the upper sunlit layer and the light exposure of phytoplankton.

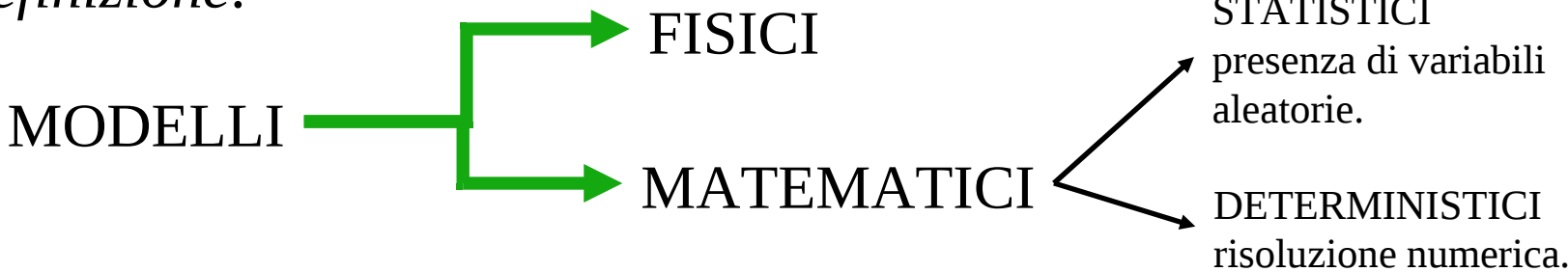
Several processes contribute to turbulent mixing in the mixed layer. Thermal convection can be generated by the ocean losing heat through longwave back radiation or evaporative cooling. The shear generated in wind-driven currents can produce Kelvin-Helmholtz billows. The interaction between surface waves and wind-driven shear current also produces Langmuir circulation, consisting of counter-rotating vortices with their axes aligned roughly in the wind direction.



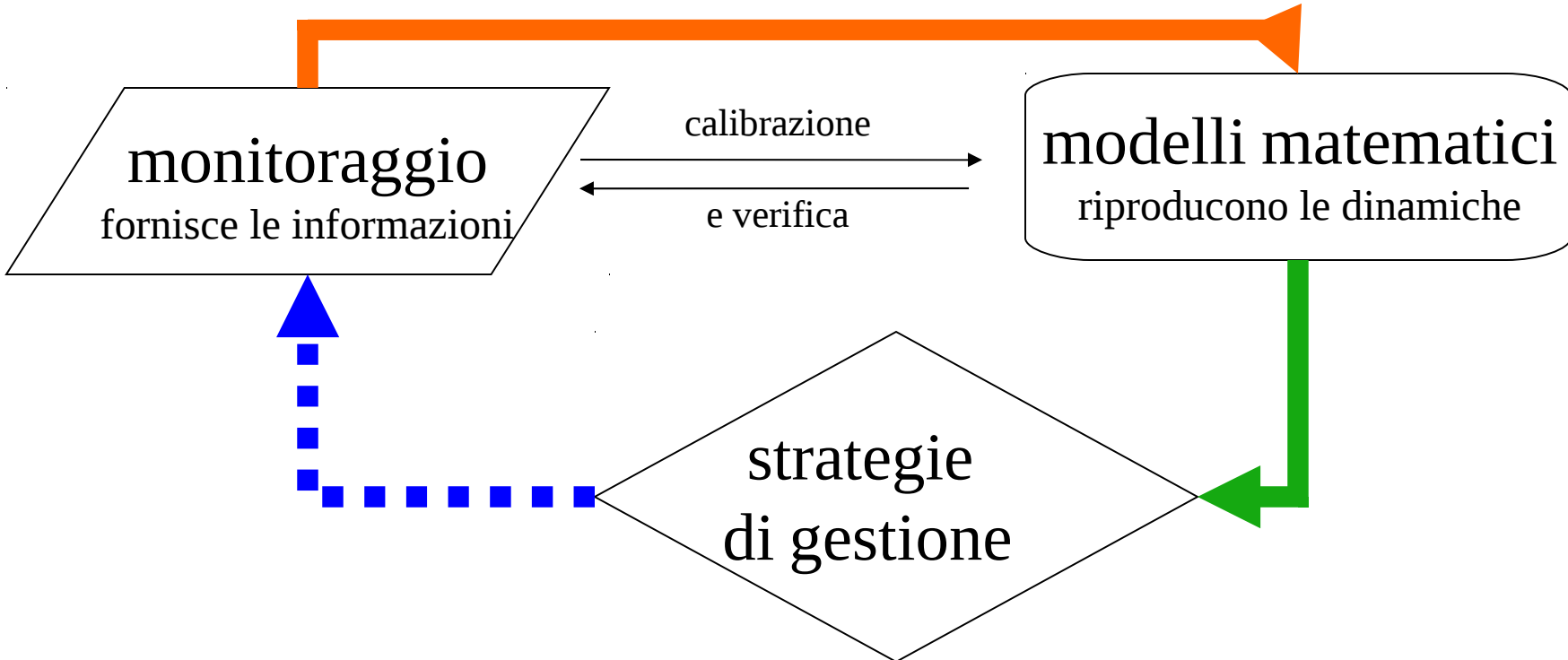
Advanced Very High Resolution Radiometer (AVHRR) image of a phytoplankton bloom south of Iceland in June 1991. For each of the crosses is approximately 110 km apart. Crudely speaking the paler a region the higher the concentration of plankton cells within it. The light is being reflected by microscopic plates grown as a covering by the phytoplankton (*Emiliana*). Aside from the large scale population explosion or “bloom” between 61N and 63N there is a great deal of patchiness at smaller scales. In particular, note the strong signature of the local currents, especially eddy features and thin filaments. Movie 1, an animation of the AVHRR observations of this area from 15–23 June, can be found at [http://dx.doi.org/10.1016/S0079-611\(03\)00085-1](http://dx.doi.org/10.1016/S0079-611(03)00085-1) reveals the dynamic nature of the patchiness. The raw data was received by Dundee Satellite Receiving Station. Steve Groom, NOAA, Plymouth Marine Laboratory processed the images.



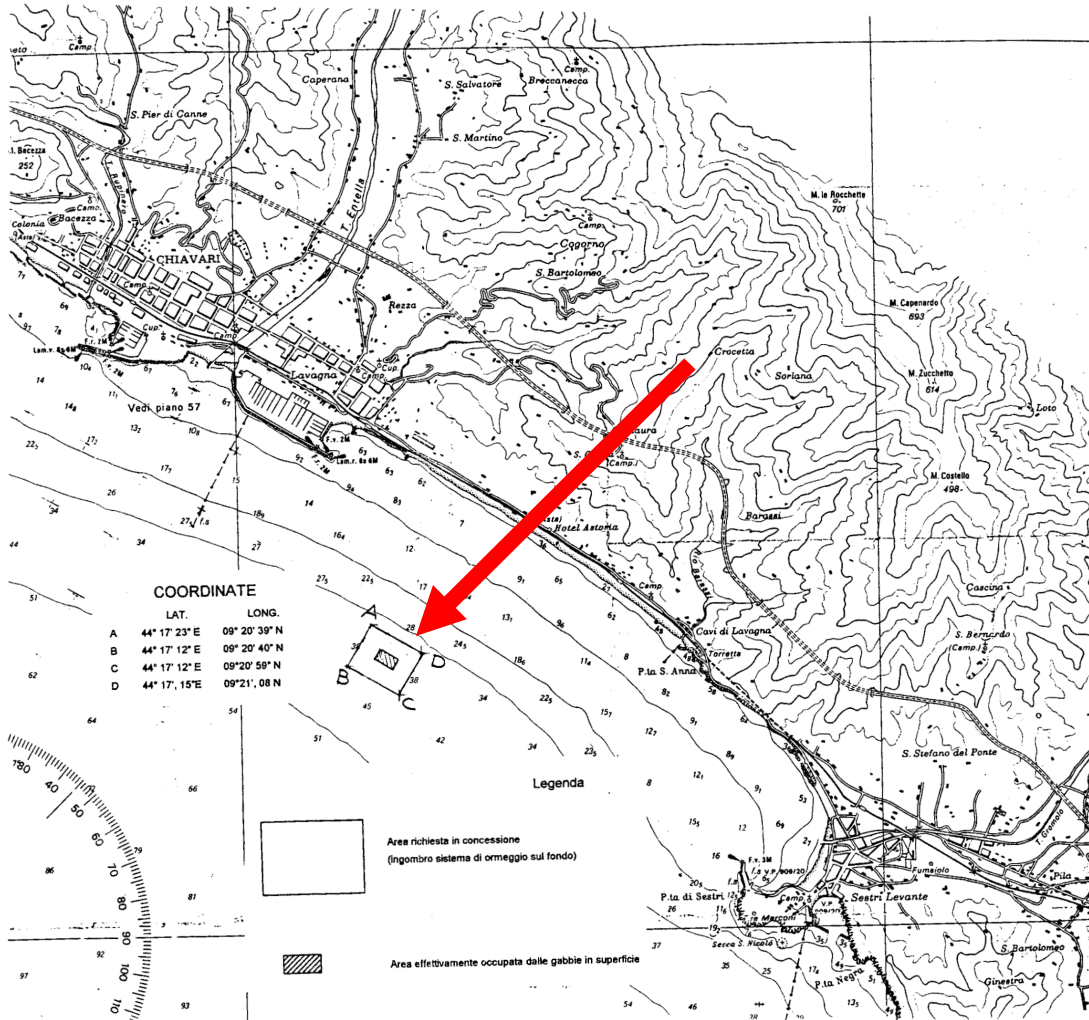
Definizione:



Utilizzo:



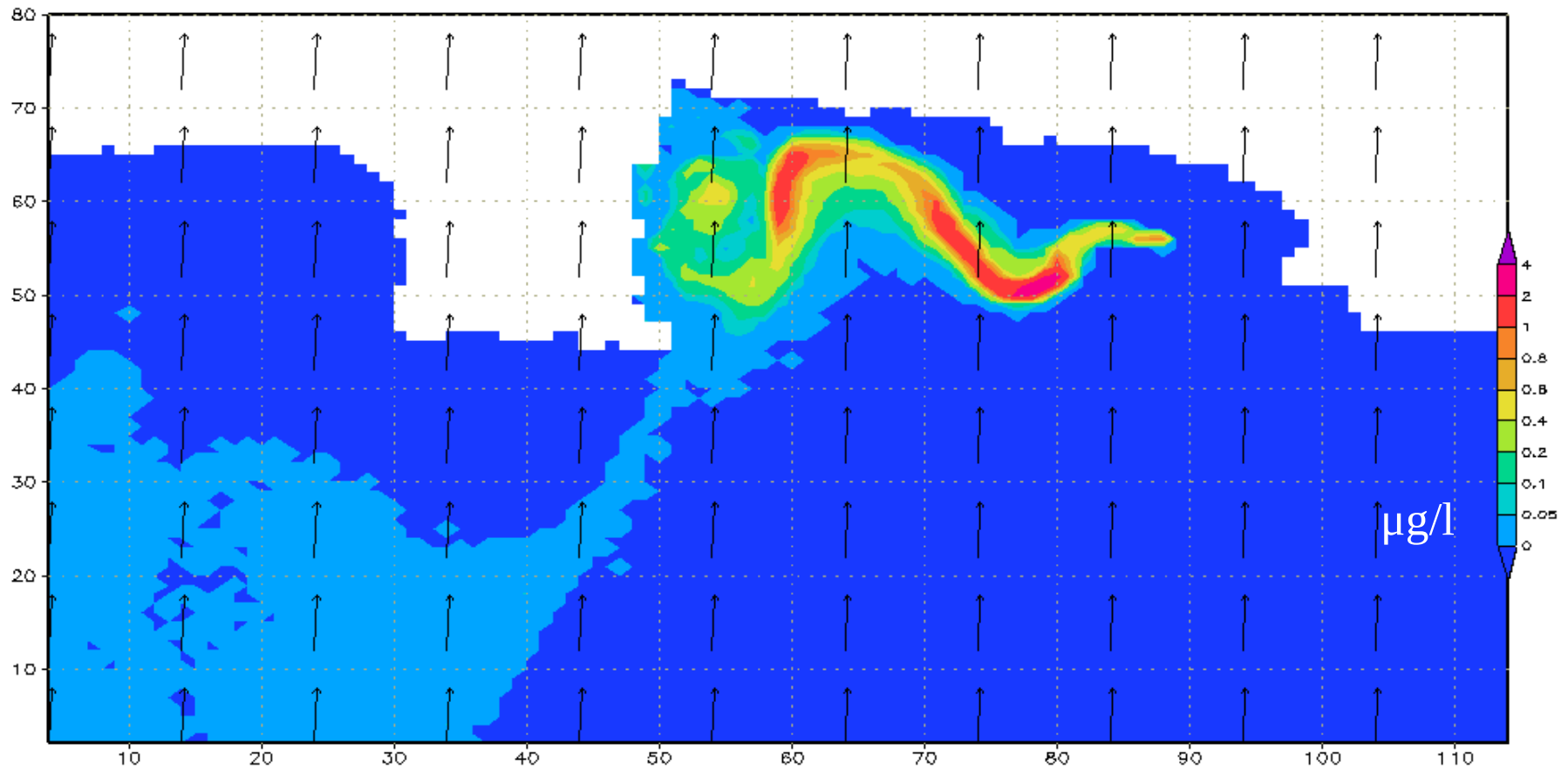
L'impianto di maricoltura



Caratteristiche:

- Impianto off-shore (35-40 m)
- Target produttivo pari a 200 tonnellate di pesce annue
- Orate e spigole all'inizio
- Successiva diversificazione delle specie ittiche
- Prodotto finito: pesci di taglia pari a 300-400 g o superiori

Simulazioni avvettivo dispersive per i nutrienti



Densità
di massa

Velocità

Pressione

Sforzi
molecolari

Coriolis

$$\left\{ \begin{array}{l} \frac{d\rho}{dt} + \rho \nabla \cdot \mathbf{v} = 0 \\ \frac{d\mathbf{v}}{dt} = -\frac{\nabla p}{\rho} + \frac{1}{\rho} \nabla \cdot \mathcal{T} + \frac{1}{\rho} \nabla \cdot \mathcal{R} - \nabla \Phi - \nabla \Phi_m + 2\boldsymbol{\Omega} \wedge \mathbf{v} \\ \frac{dT}{dt} = \kappa_T \nabla^2 T + \frac{Q}{\rho c_\alpha} \\ \frac{ds}{dt} = \kappa_s \nabla^2 s + \frac{\Sigma}{\rho} \\ \alpha = \alpha(T, \rho, s) \end{array} \right. \begin{array}{l} \text{(eq. continuità)} \\ \text{(eq. momento)} \\ \text{(eq. calore)} \\ \text{(eq. salinità)} \\ \text{(eq. di stato)} \end{array}$$

Sforzi turbolenti

Temperatura

Salinità

Volume
specifico

Geopotenziale

Potenziale di marea

Processi di trasporto

AVVEZIONE

Movimento dell'inquinante dovuto a processi di flusso risolti

DISPERSIONE

Distribuzione dell'inquinante nel fluido dovuta a processi di flusso non risolti

DIFFUSIONE

Moti molecolari e turbolenza

Fick (1855)

Coeff. di diffusione

Coeff. di dispersione

$$M \propto \nabla c$$

Taylor (1921)

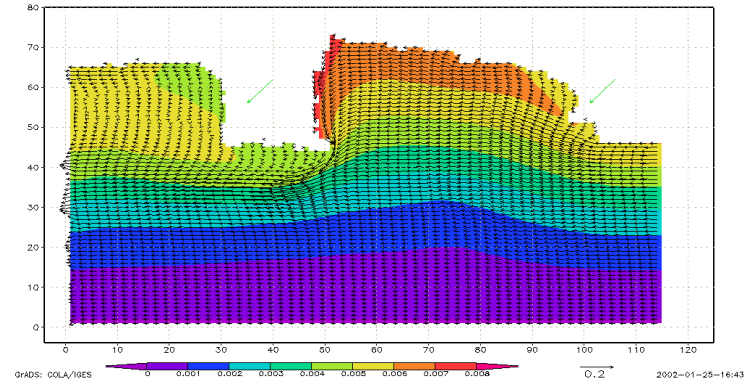
SHEAR

Gradienti spaziali di velocità

Applicazione dei modelli

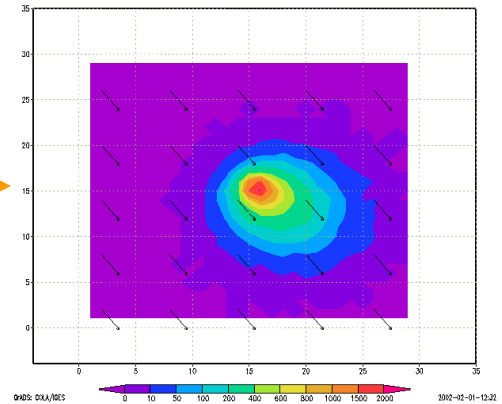
POM2D

Calcola le
velocità mediate
in verticale



LAMP3D

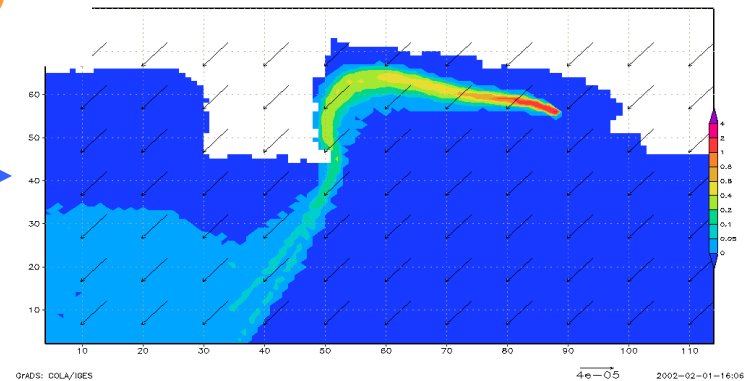
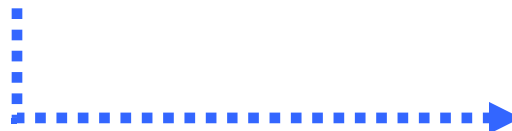
Calcola gli
spostamenti
delle particelle



POM2D

+

LAMP3D



Processi di trasporto

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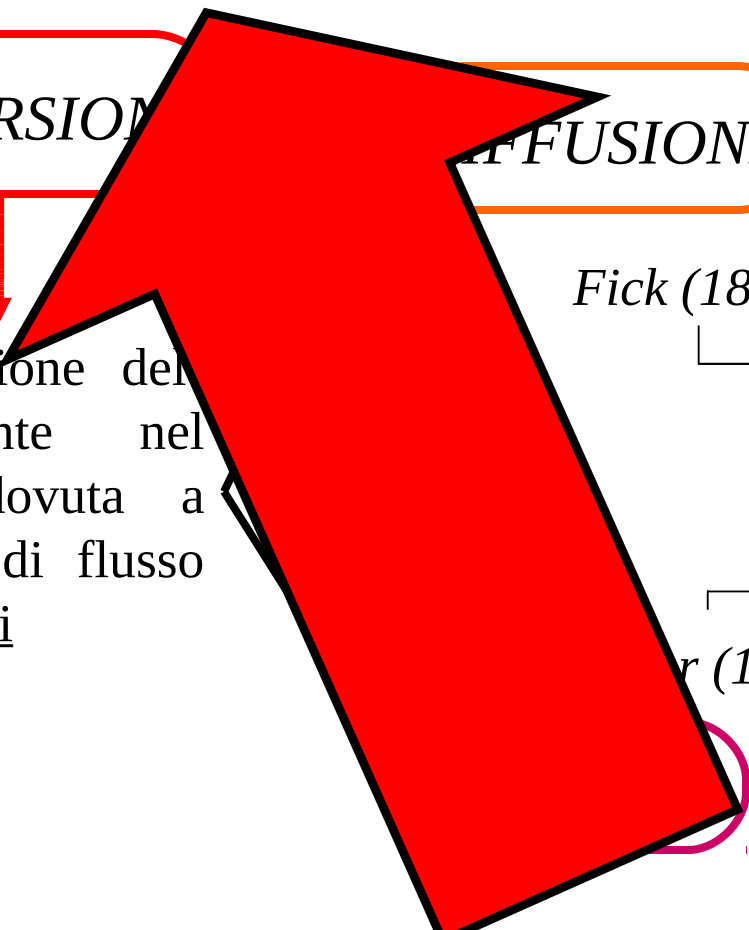
Coeff. di diffusione

Coeff. di dispersione

Taylor (1921)

Gradienti spaziali di velocità

$$M \propto \nabla c$$





POM (Princeton Ocean Model)

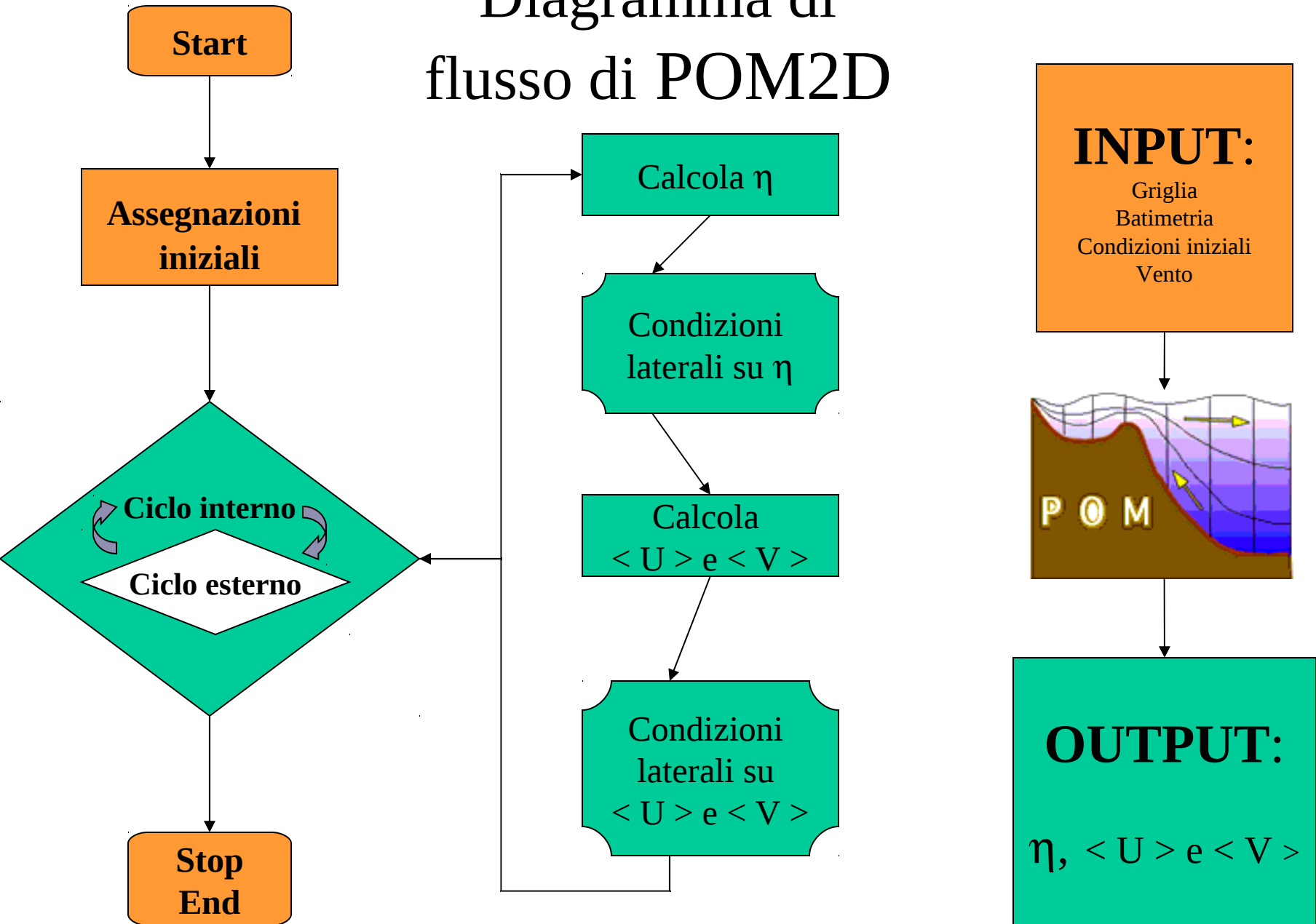
www.aos.princeton.edu/WWWPUBLIC/htdocs.pom.

- Finite difference
- Free surface.
- Primitive equations (hydrostatic, Boussinesq).
- Horizontal grid: Curvilinear orthogonal coordinate.
- Mode split technique:
 - external mode (2D depth averaged, barotropic) $\Rightarrow EL(i,j), UA(i,j), VA(i,j)$
 - internal mode (3D, baroclinic) $\Rightarrow U(i,j,k), V(i,j,k), T(i,j,k), S(i,j,k)$

BAROTROPIC APPROXIMATION:

- Simpler to calibrate;
- More numerically stable;
- Faster as computing time.

Diagramma di flusso di POM2D

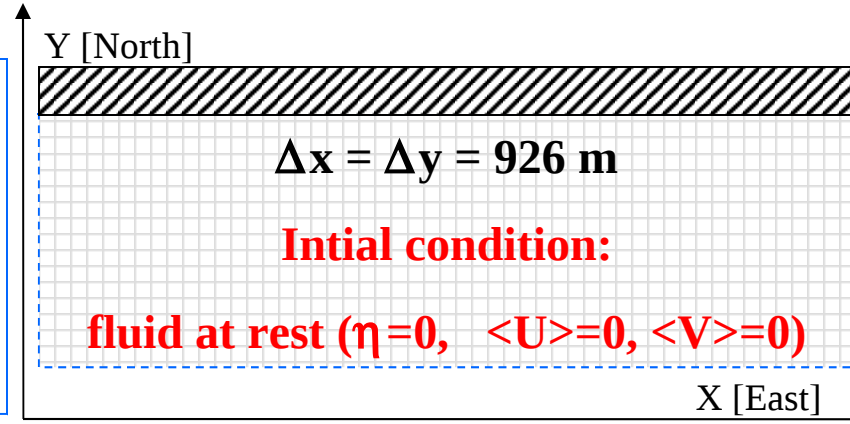


WIND TESTS with POM-2D mode

BCOND settings

Western Boundary:

- η : radiation (gravity wave velocity)
- $\langle U \rangle$: zero gradient
- $\langle V \rangle$: zero gradient



Eastern Boundary:

- η : radiation (gravity wave velocity)
- $\langle U \rangle$: zero gradient
- $\langle V \rangle$: zero gradient

Southern Boundary

- η : clamped $\eta=0$
- $\langle U \rangle$: zero gradient $\langle V \rangle$: zero gradient

Vertically-integrated and linearized momentum equation for a homogeneous fluid with surface and bottom stresses

[Chapman 85]

$$\frac{\partial P}{\partial t} - fQ = -gH \frac{\partial \eta}{\partial x} + \frac{\tau_x^S}{\rho} - \frac{\tau_x^B}{\rho}$$

$$\frac{\partial Q}{\partial t} + fP = -gH \frac{\partial \eta}{\partial y} + \frac{\tau_y^S}{\rho} - \frac{\tau_y^B}{\rho}$$

where

$$P = \int_{-H}^0 u dz$$

$$Q = \int_{-H}^0 v dz$$

TEST 2

Linear Bathymetry: $H=c*y$, Cross-shelf wind: $-\tau_y^s/\rho = -10^{-4} \text{ m}^2 \text{ s}^{-2}$

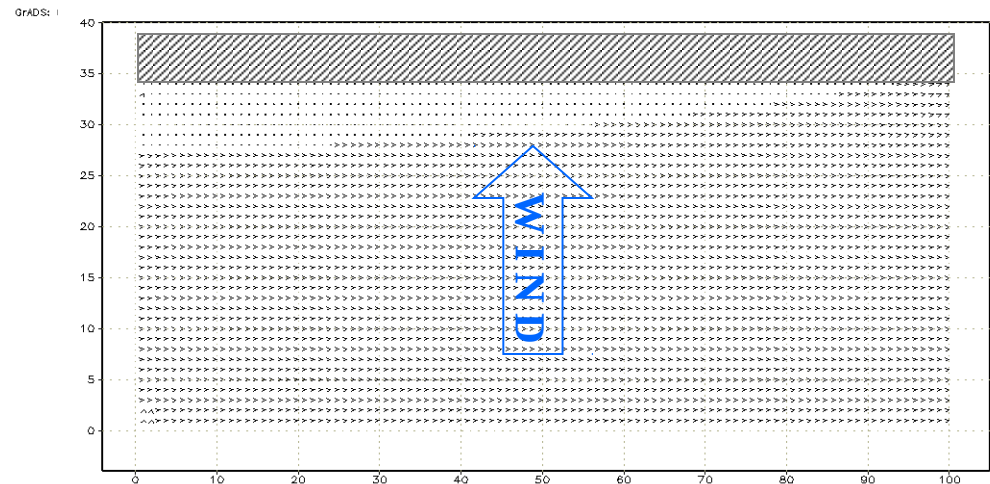
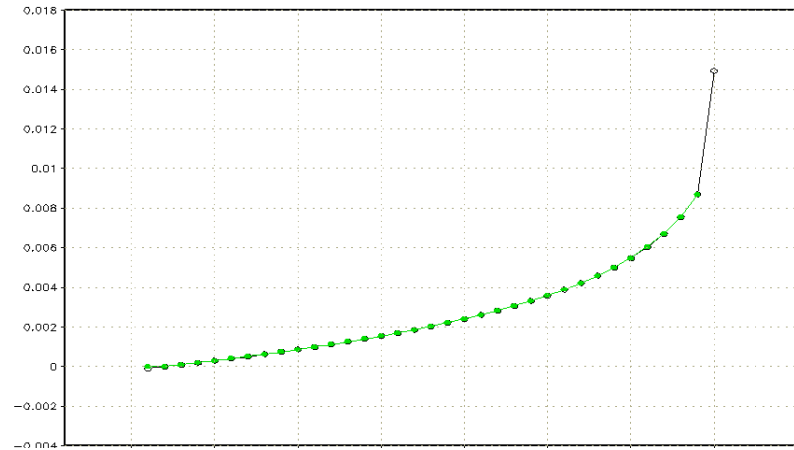
Hypothesis:

- Steady state
- $P = Q = 0$ everywhere
(balance between wind stress and horizontal pressure gradient)
- Bottom stress negligible

Analytical solution

$$\eta(y) = -\frac{\tau_y^s}{\rho g m} \ln \left[\frac{H(y)}{H(-L)} \right]$$

$$\frac{\partial \eta}{\partial x} = 0$$



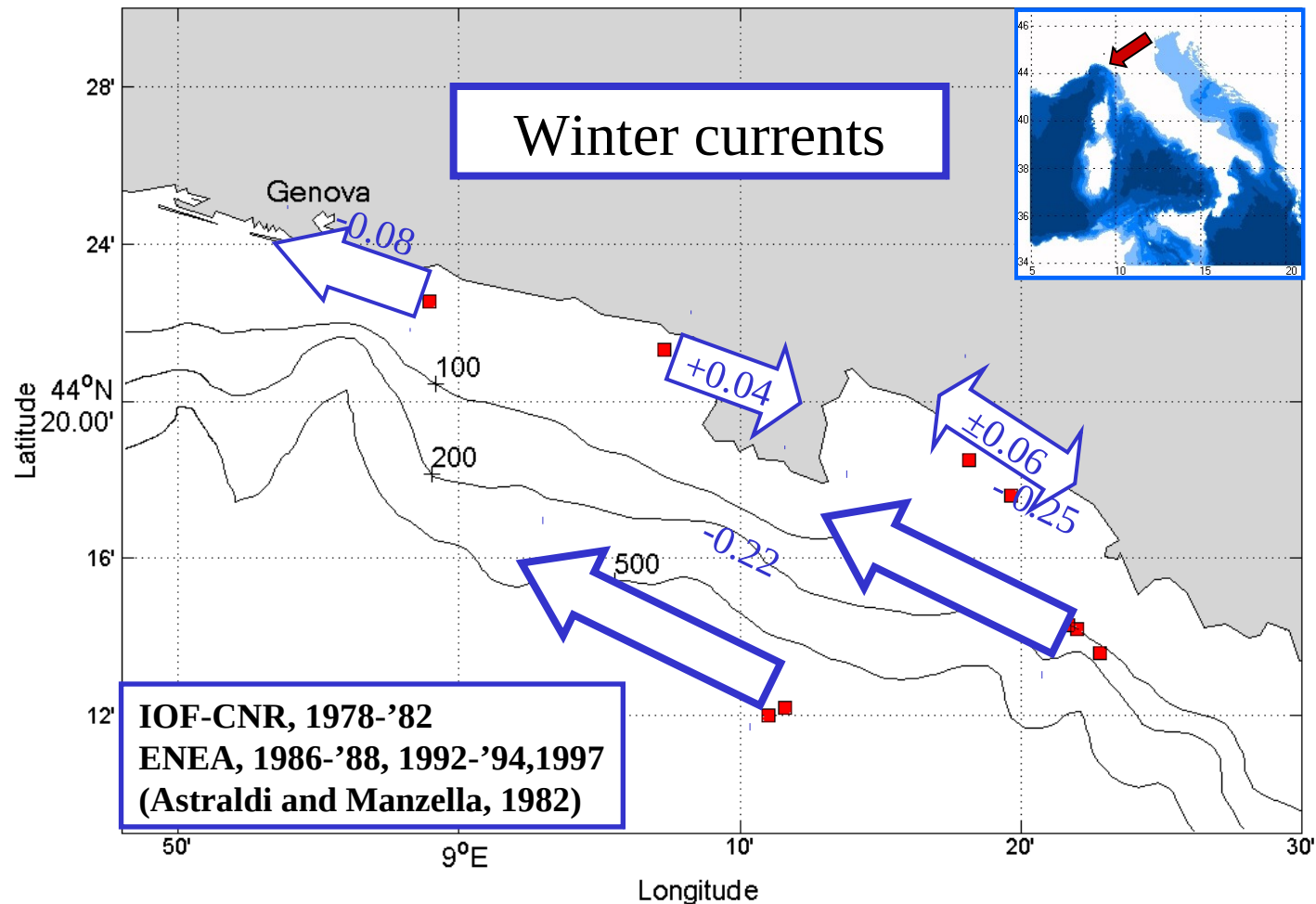
Study area

Bathymetry

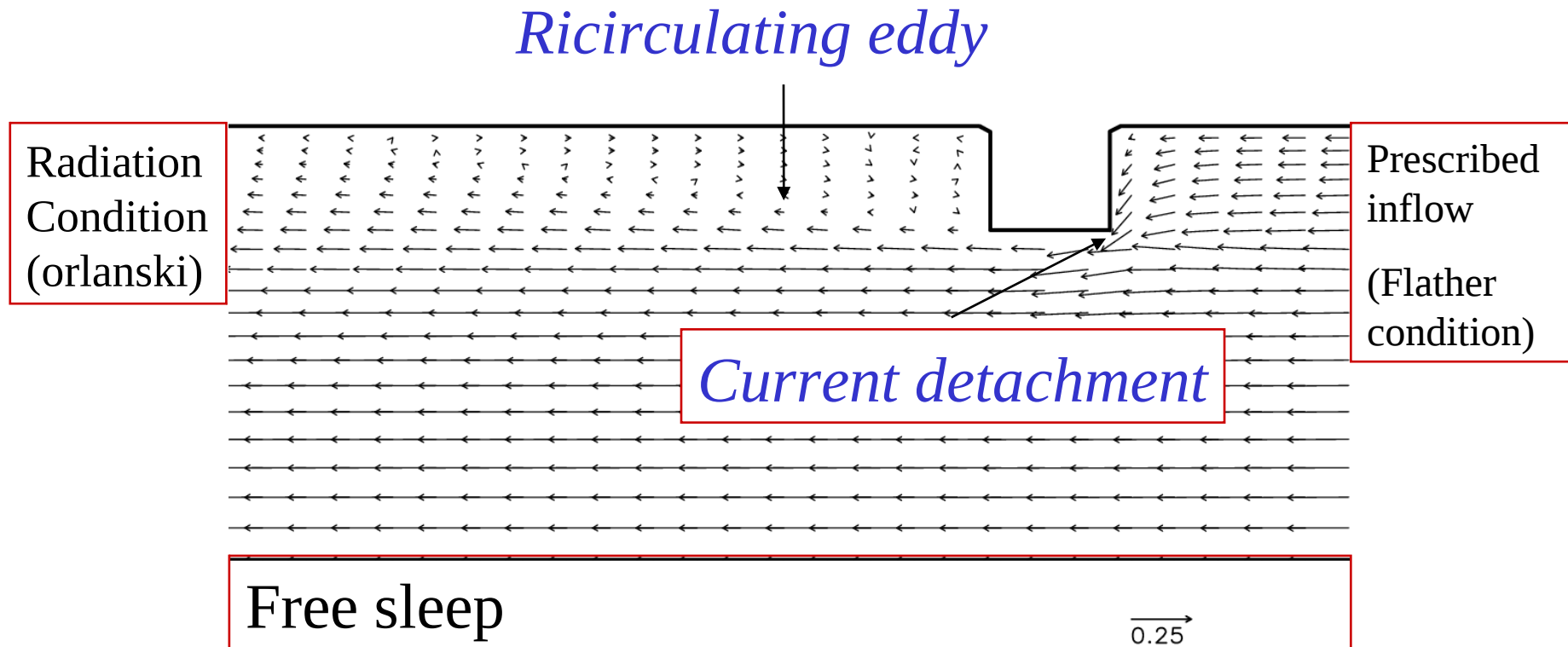
- headland Promontorio di Portofino;
- abrupt coast;
- narrow shelf and steep slope.

Current measurements (winter time)

- persistent geostrophic inflow;
- persistent countercurrent in the Golfo Paradiso;
- oscillating direction in the Golfo del Tigullio;

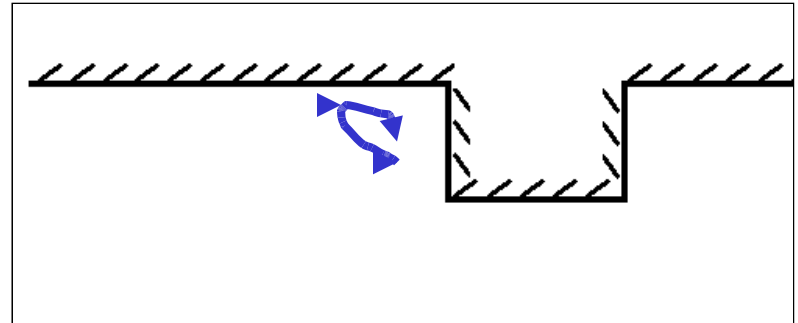
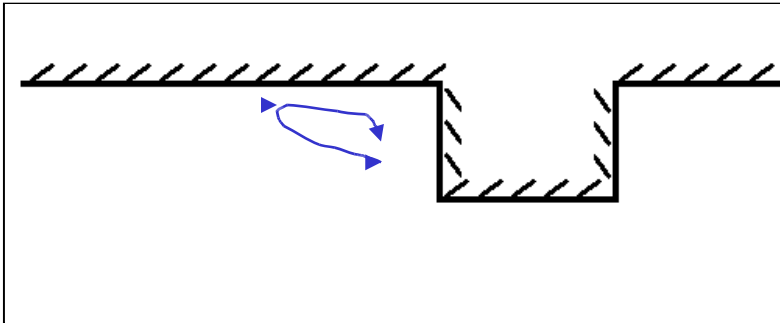
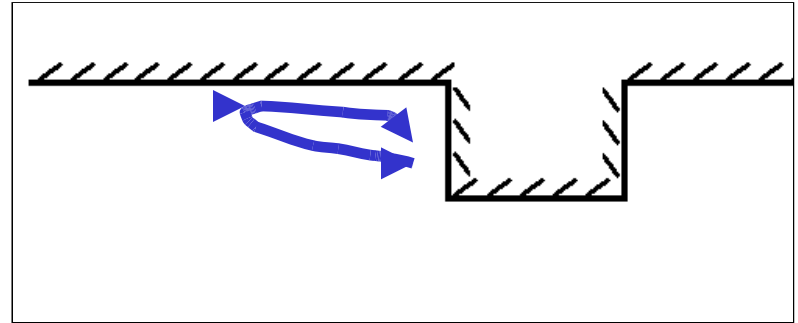
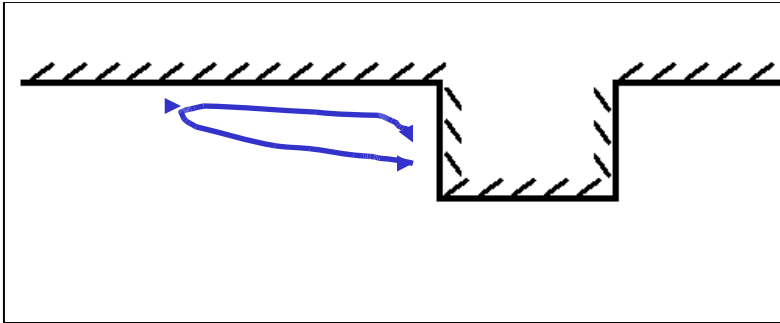
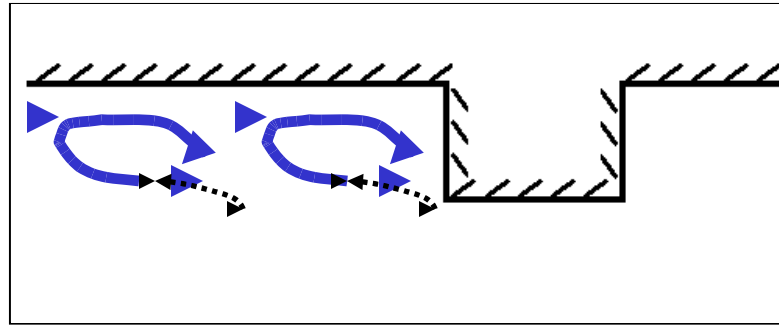


Numerical setting and typical solution



Summary scheme

Re_f

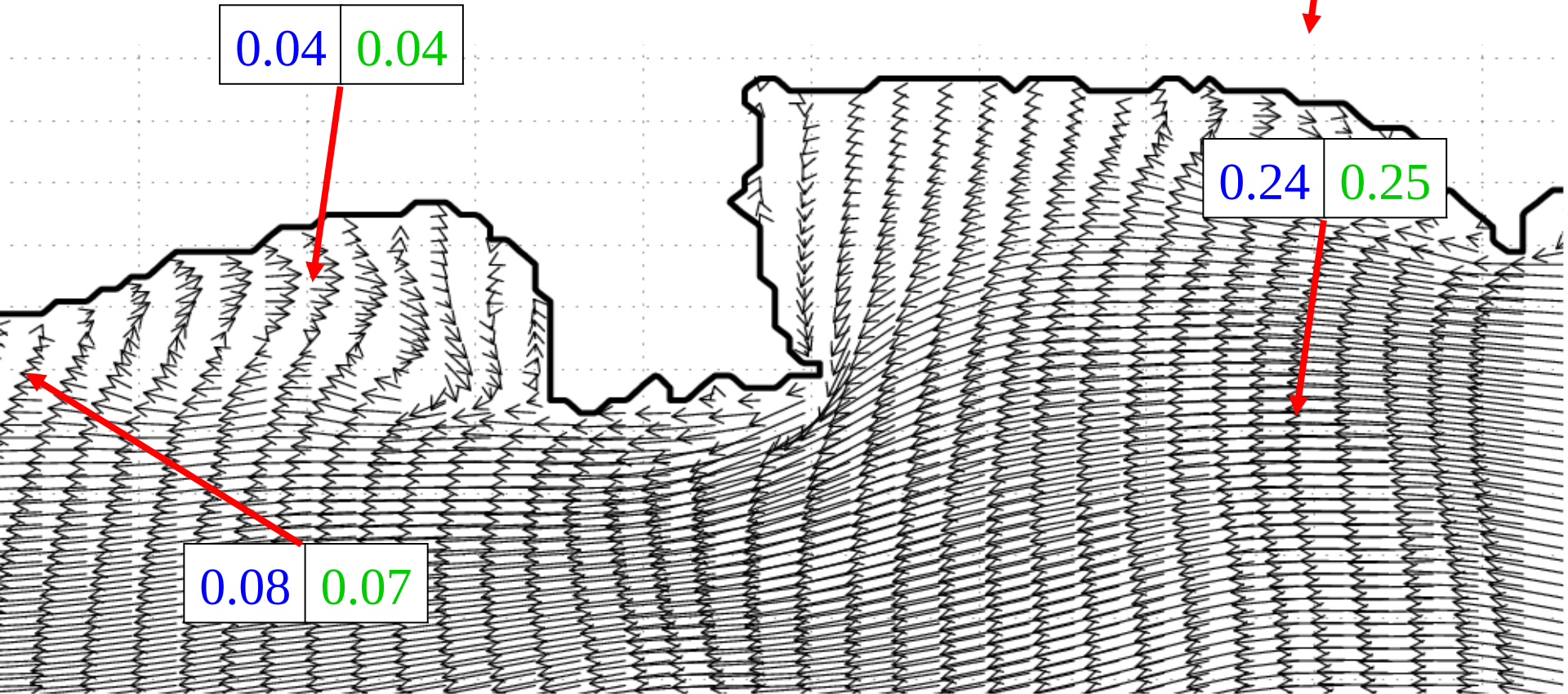


K_M

Real bathymetry REf=40

3D *Good agreement with experimental data*

data	model
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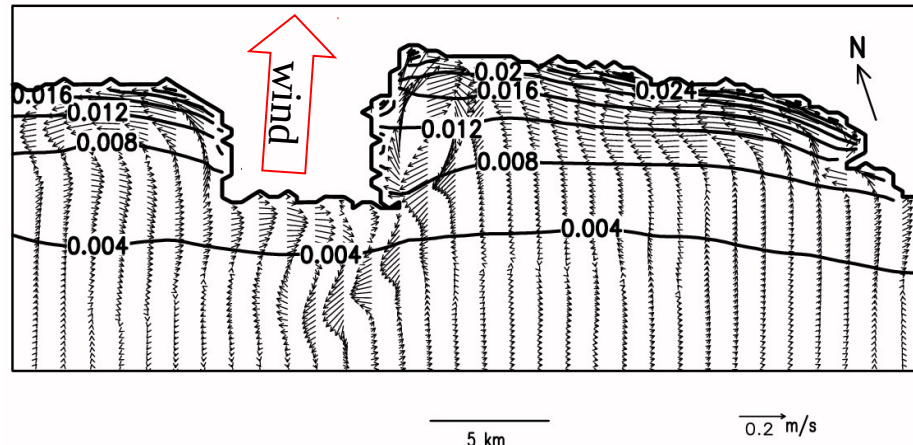
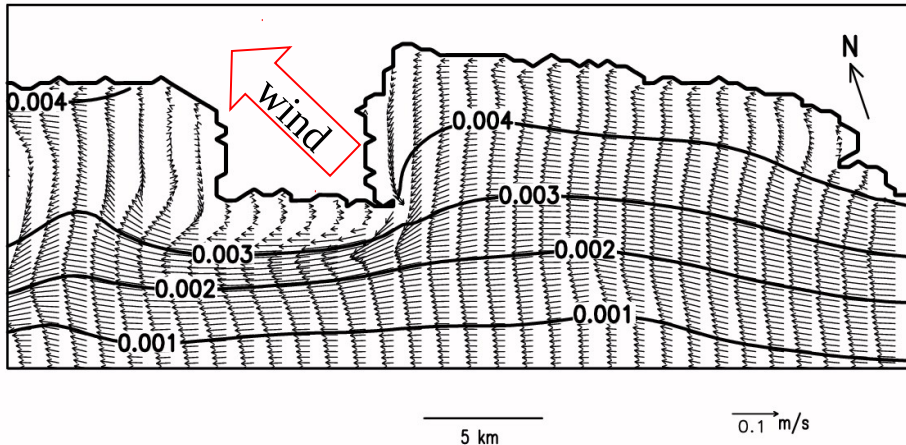
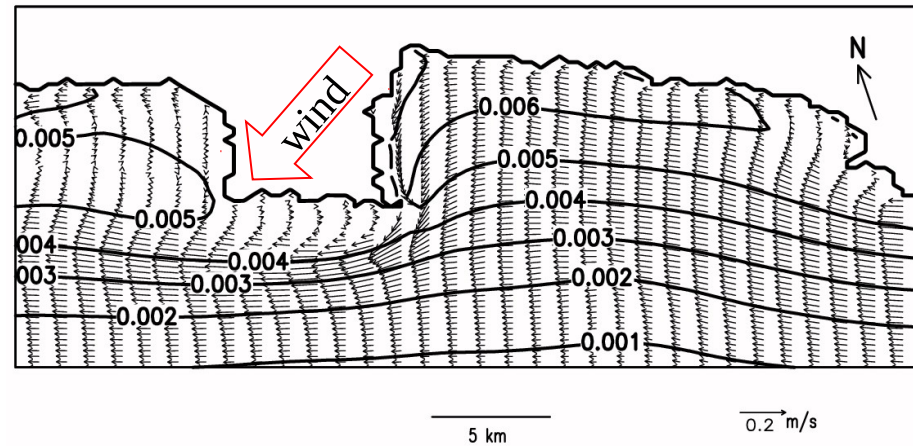
5 km $\xrightarrow{0.2}$ m/s

Water circulation

12 days simulation with
a typical local wind sequence

LEGEND:

Free surface elevation [m] and
current depth averaged velocity [m/s]



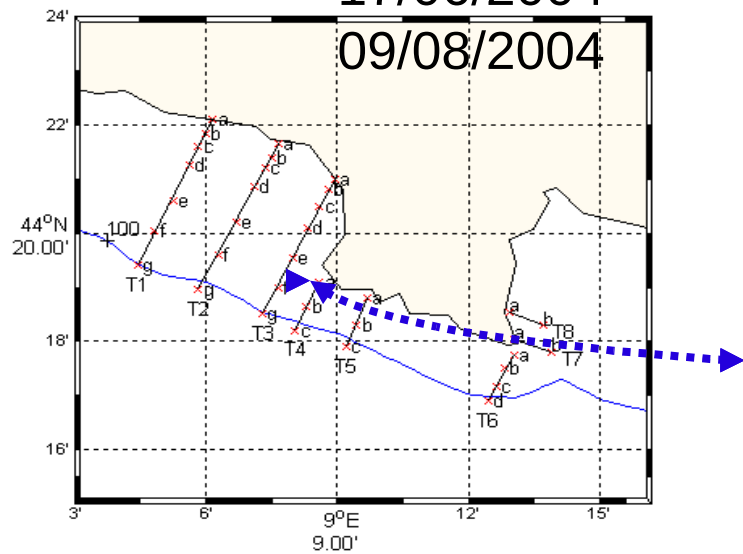
Along shelf currents calculated by the model are consistent with:

- *Measurements*
- *Other numerical experiments*

Sonde CTD

17/06/2004

09/08/2004

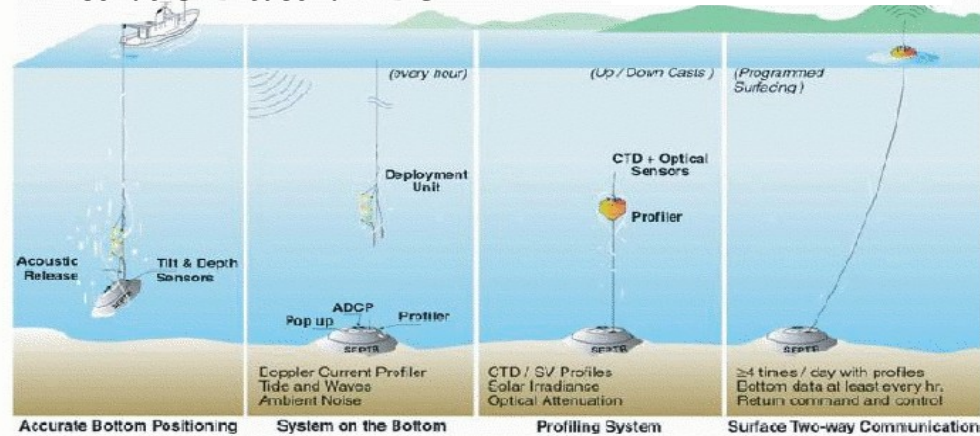


SEPTR du NURC-OTAN

23jul-18aug2003

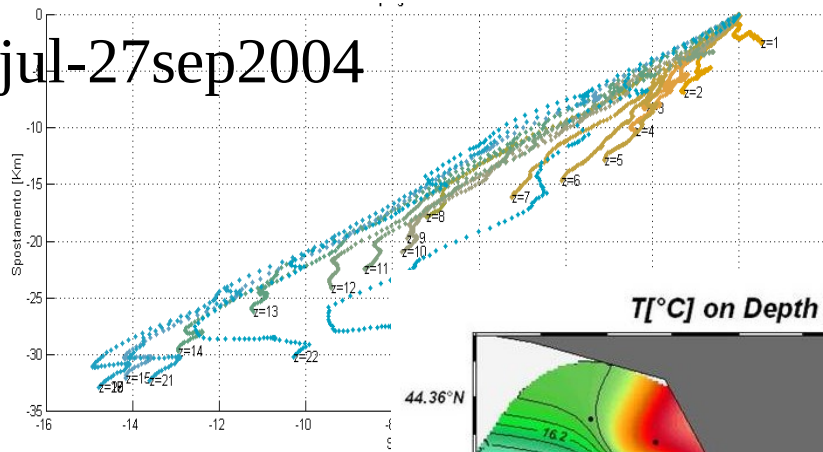
01jul-27sep2004

sonde CTD et sonar ADCP

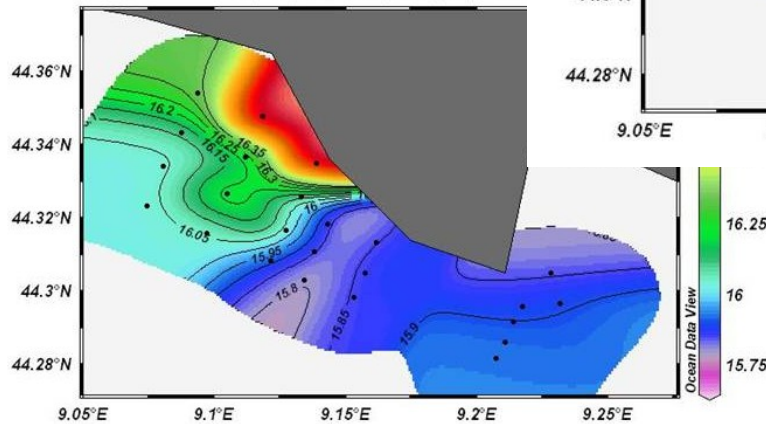


SEPTR operational scenario: deployment and operations

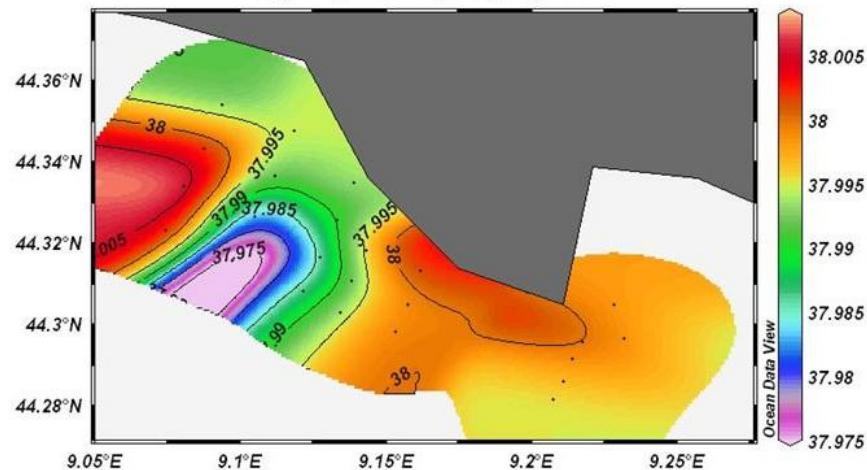
01jul-27sep2004



$T[^\circ\text{C}]$ on Depth $[m]=40$



$S[\text{psu}]$ on Depth $[m]=40$



09/08/2004

Processi di trasporto

~~AVVEZIONE~~

Movimento dell'inquinante dovuto a processi di flusso risolti

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Distribuzione dell'inquinante nel fluido dovuta a processi di flusso non risolti

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$$M \propto \nabla c$$

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SHEAR

Gradienti spaziali di velocità

Approccio Euleriano (alla dispersione)

- descrizione delle caratteristiche del sospeso/soluto tramite campi di concentrazione media $\langle c(\mathbf{r},t) \rangle$;
- evoluzione del campo $\langle c(\mathbf{r},t) \rangle$ basata sull'equazione di avvezione-dispersione.

Equazione di avvezione-dispersione

- $\mathbf{v} = \langle \mathbf{v} \rangle + \mathbf{v}'$ e $c = \langle c \rangle + c'$;
- ipotesi di ergodicità;
- teoria K .

$$\frac{\partial \langle c \rangle}{\partial t} + \mathbf{v} \cdot \nabla \langle c \rangle = K_{SH} \nabla^2 \langle c \rangle + K_{SV} \frac{\partial^2 \langle c \rangle}{\partial z^2} + \frac{\xi}{\rho}$$

Approccio Lagrangiano (alla dispersione)

- descrizione delle caratteristiche del sospeso/soluto tramite particelle;
- dinamica probabilistica di ogni particella basata sulla “probabilità di transizione” $P(\mathbf{r},t | \mathbf{r}_0,t_0)$ per ogni specie di inquinante.

Probabilità di transizione:

Probabilità che una particella che si trova in \mathbf{r}_0 all'istante t_0 si trovi in \mathbf{r} all'istante t .

$$\langle c(\mathbf{r}, t) \rangle = \int_{-\infty}^t \int_V P(\mathbf{r}, t | \mathbf{r}_0, t_0) \xi(\mathbf{r}_0, t_0) d\mathbf{r}_0 dt_0$$

Modelli a particelle

$$\mathbf{v}_e = \bar{\mathbf{v}} + \mathbf{v}'$$

da misure euleriane di corrente
oppure
fornita da un modello idrodinamico

‘velocità di diffusione’

*Random
walk*

approssimazione AR(0)

$$\mathbf{v}'_n = \boldsymbol{\mu}$$

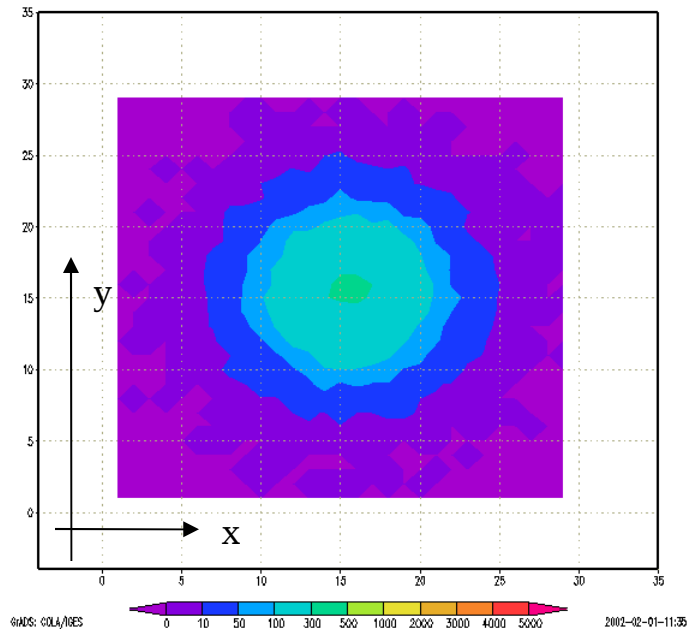
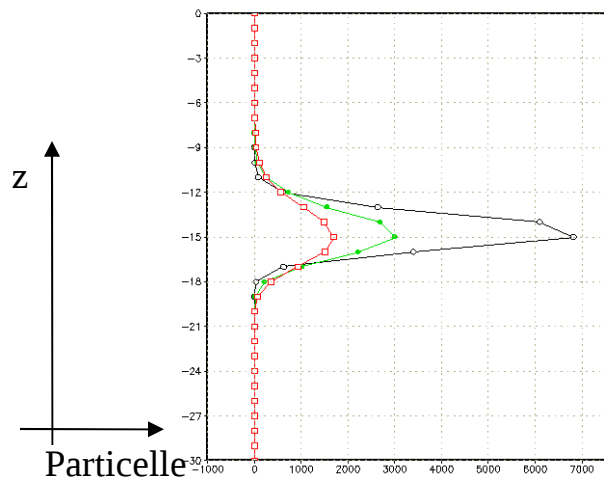
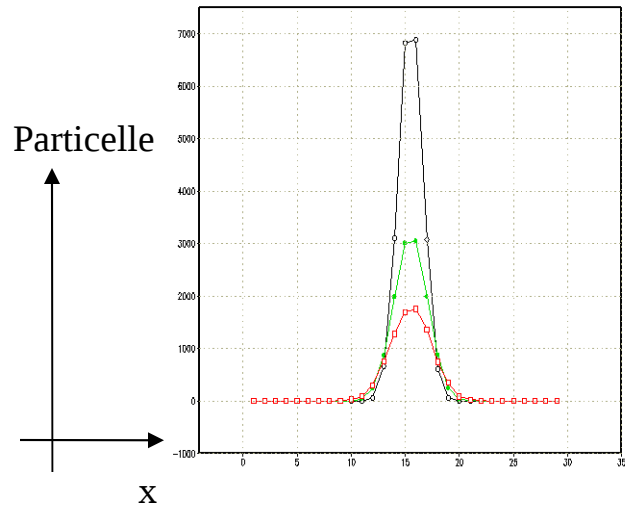
$$\boldsymbol{\varrho} = \boldsymbol{\mu} \Delta t$$

$$\mathbf{r}_{n+1} - \mathbf{r}_n = \bar{\mathbf{v}} \Delta t + \boldsymbol{\varrho}$$

estrazione

$\mathbf{G}(0, \boldsymbol{\sigma})$

TEST di *LAMP3D*



Velocità nulla
Test di diffusione

INPUT

LAMP2D

Legge di decadimento

$$c = c_0 \cdot 10^{-\frac{t}{T_{90}}}$$

Ciclo per le particelle

Ciclo per i passi temporali

C_n

Posizione
Distanze e pesi

v_{det}

Δr_{det}

Δr_{ran}

C_{n+1}

Ciclo per i passi temporali

Ciclo per le particelle

Controllo
posizione

Controllo
posizione

si

no

no

si

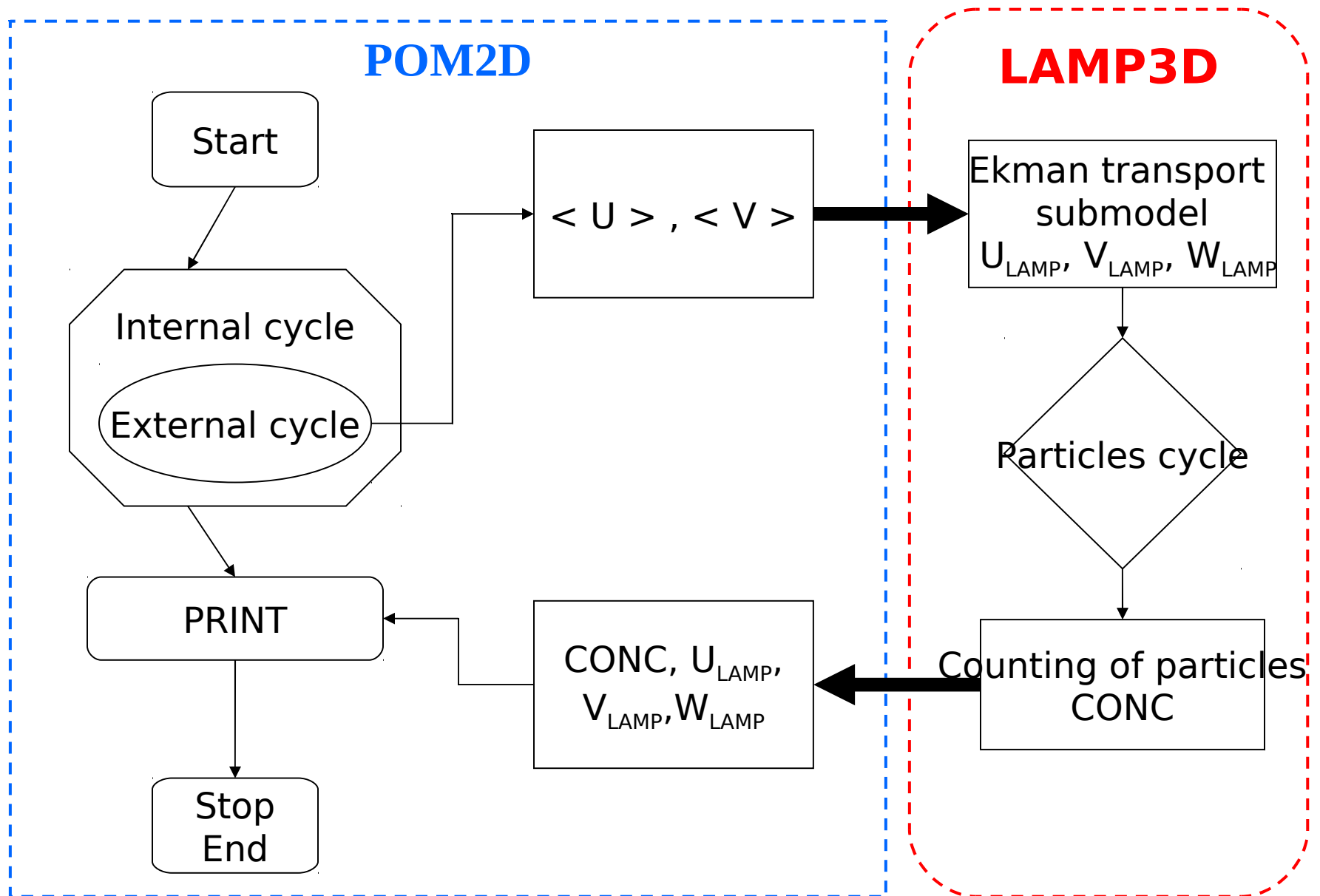
Particelle
decadenti

si

OUTPUT

LAMP2D

POM2D-LAMP3D coupled model



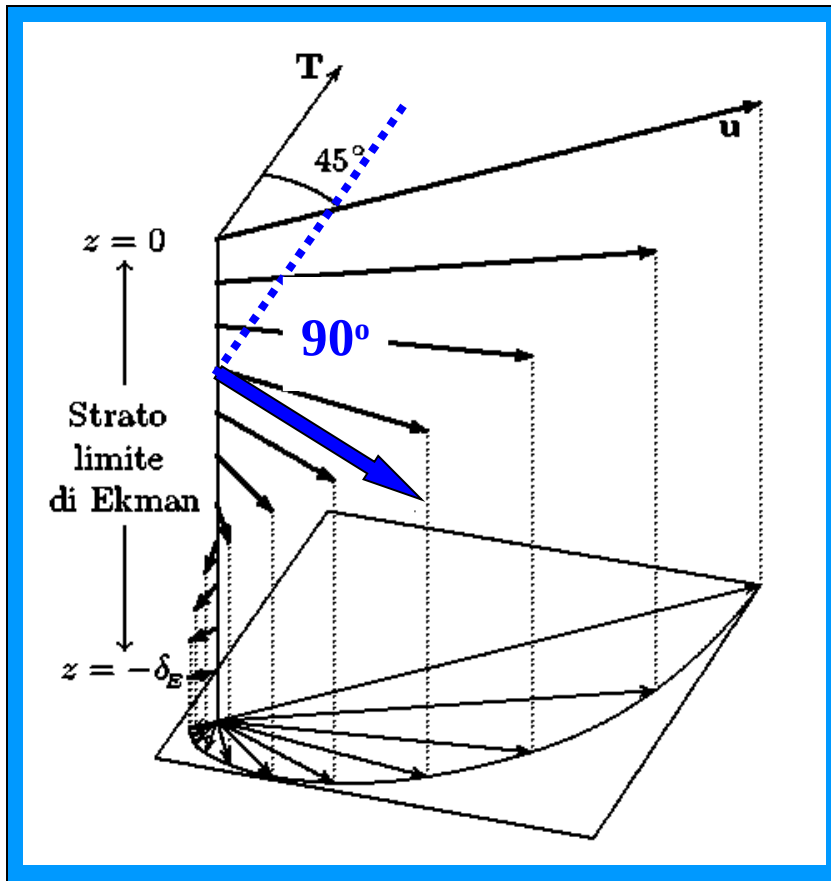
La spirale di Ekman

Profilo velocità

- fluido omogeneo ed *infinito*
- gradiente p orizzontale nullo
- moto stazionario
- vento costante

$$u + iv = V e^{\frac{\pi}{\delta_E} z} e^{i\left(\frac{\pi}{4} + \frac{\pi}{\delta_E} z\right)} \quad (z \leq 0)$$

$$\delta_E = \pi \sqrt{\frac{2A_V}{f}}$$



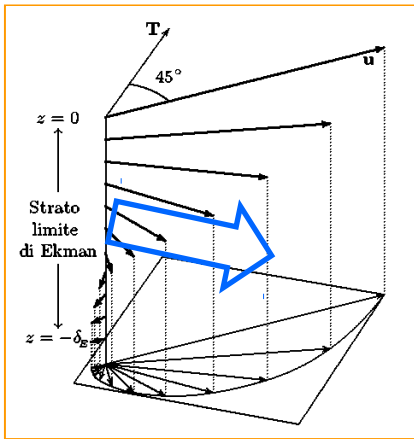
Trasporto

- integrale tra $z = 0$ e $z = \delta_E$
- trascurare sforzo del vento a $z = \delta_E$

$$\mathbf{M} = -\frac{1}{f} \mathbf{k} \wedge \mathbf{T}$$

$$\mathbf{M} = \int_{-\delta_E}^0 \mathbf{v}_H dz$$

$$\mathbf{T} = \left(A_V \frac{\partial \mathbf{v}_H}{\partial z} \right)_{z=0}$$



$$u + iv = V e^{\frac{\pi}{\delta_E} z} e^{i\left(\frac{\pi}{4} + \frac{\pi}{\delta_E} z\right)} \quad (z \leq 0)$$

Ekman transport submodel

Spiral profile
in a limited
layer

$$(u + iv) = \langle u + iv \rangle \kappa H \mathcal{F}(z)$$

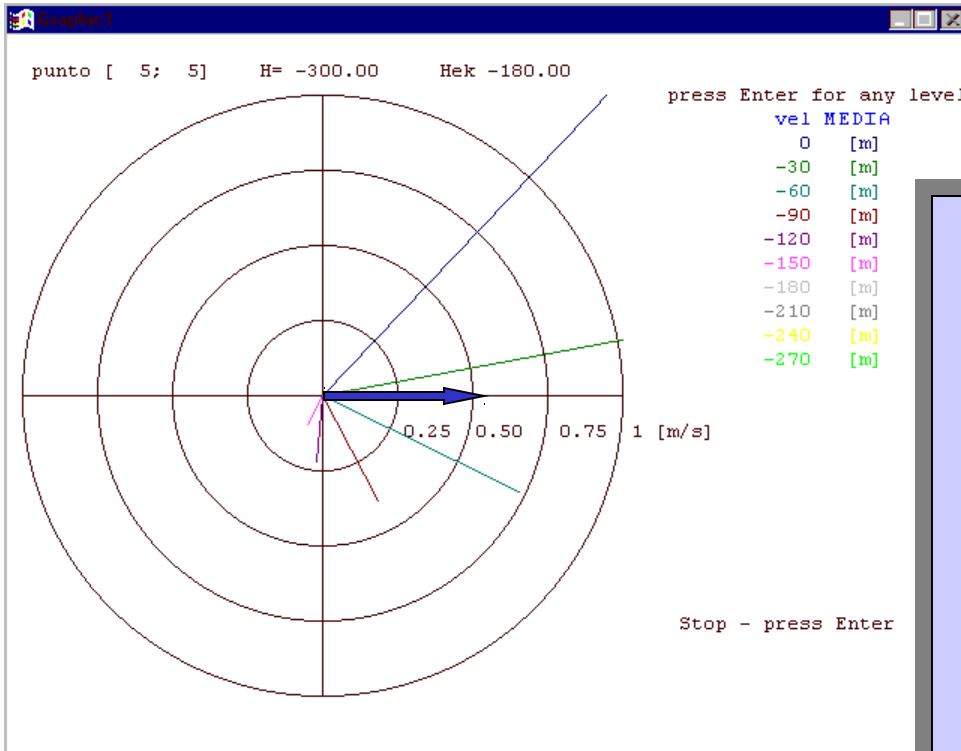
$u(x,y,z)$
 $v(x,y,z)$

 $w(x,y,z)$

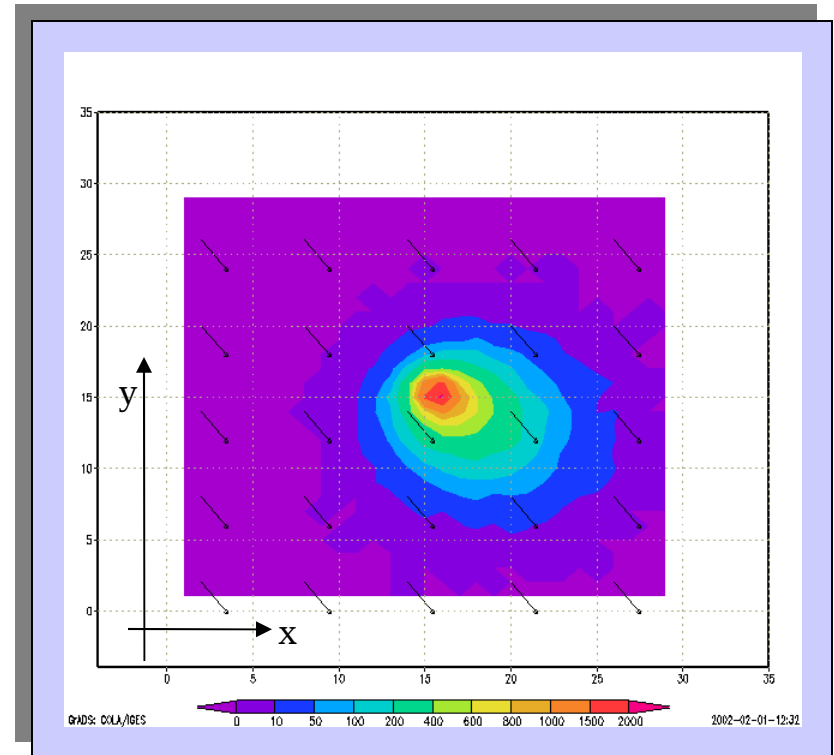
Mass
conservation

$$\frac{\partial w}{\partial z} = -\mathcal{Re} \left\{ \frac{\partial \mathbf{v}}{\partial x} \right\} - \mathcal{Im} \left\{ \frac{\partial \mathbf{v}}{\partial y} \right\}$$

TEST di *LAMP3D*

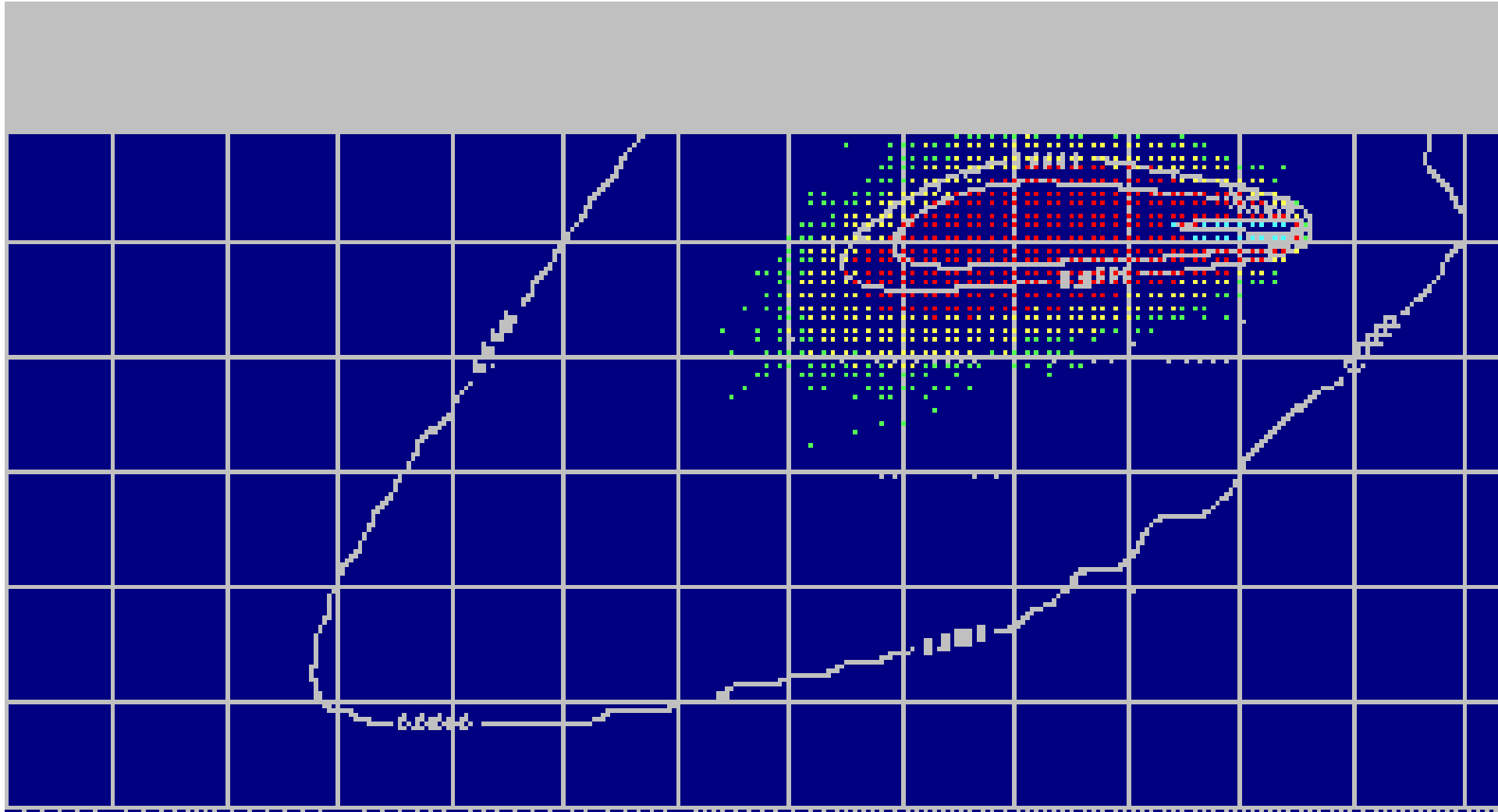


Corrente “orientale”
(Ovest-Est)
di 0.5 m/s



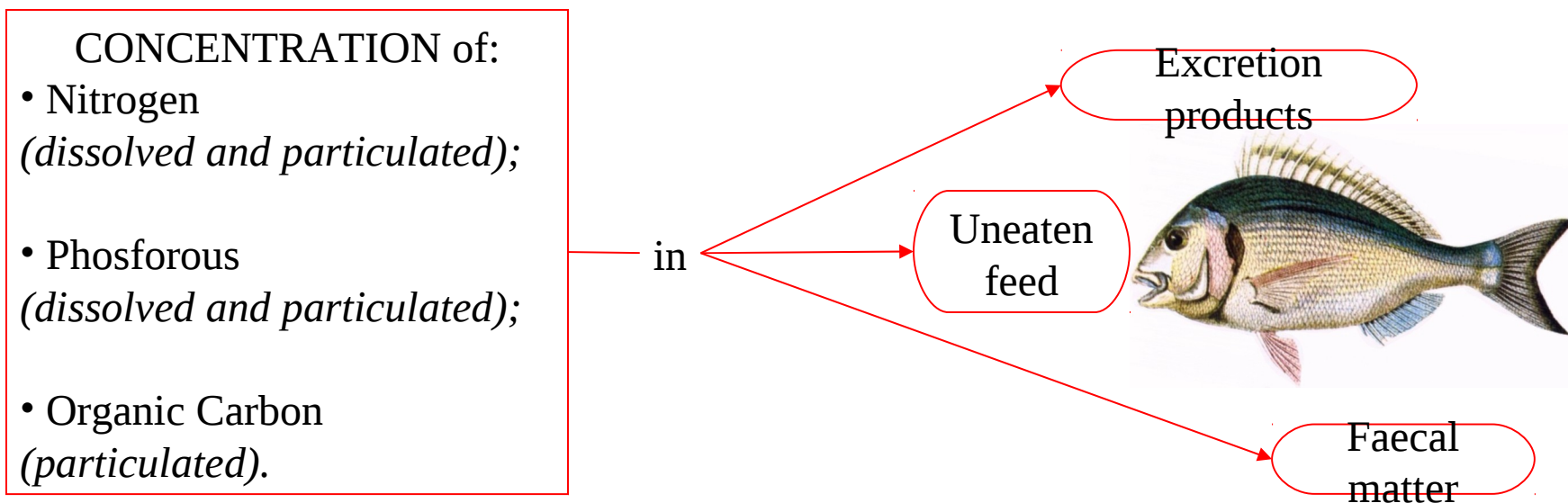
Tipica forma a goccia per un
rilascio continuo

Comparison between LAMP (dots) and MIKE21 (contour lines)



CONC< 5 5 <CONC> 50 50 <CONC> 500

Wastes Indicators



to assign a weight to particles

From literature:

- % of uneaten feed on the feed supplied;
- Velocities of sedimentation for particulated matter;
- % of each nutrient in feed (if not provided by feed producer);
- the soluble and sedimentable fractions of metabolic N and P;
- faecal production of fishes;

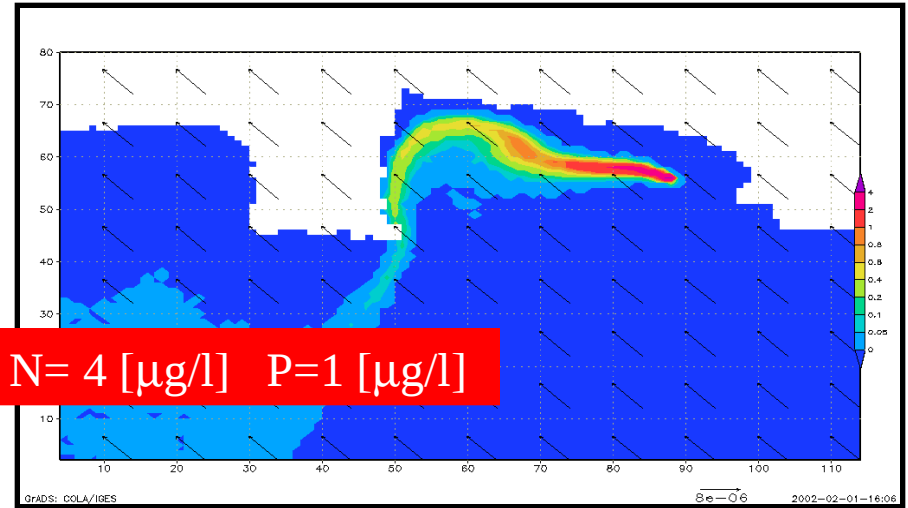
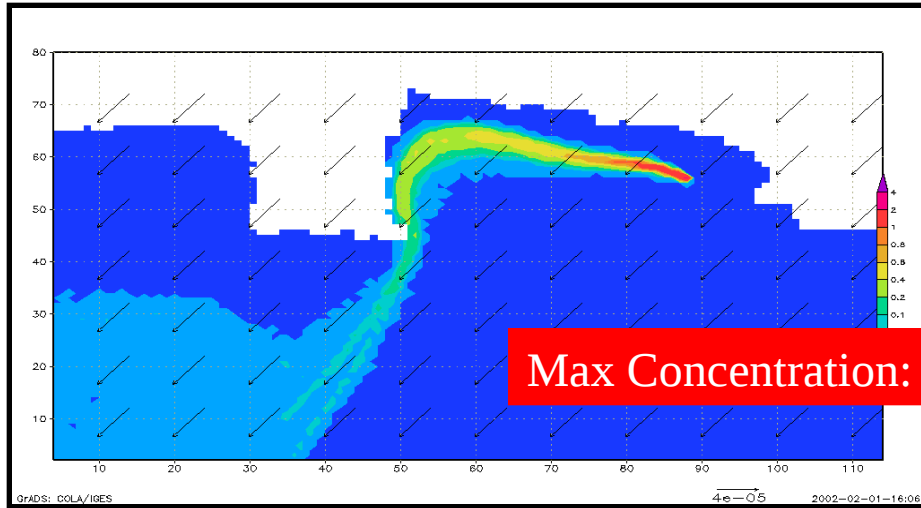
Assumptions:

- fish production
 $P = 200$ ton/year
- feed conversion factor
 $F_c = 1.3$ kg pellet/kg fish

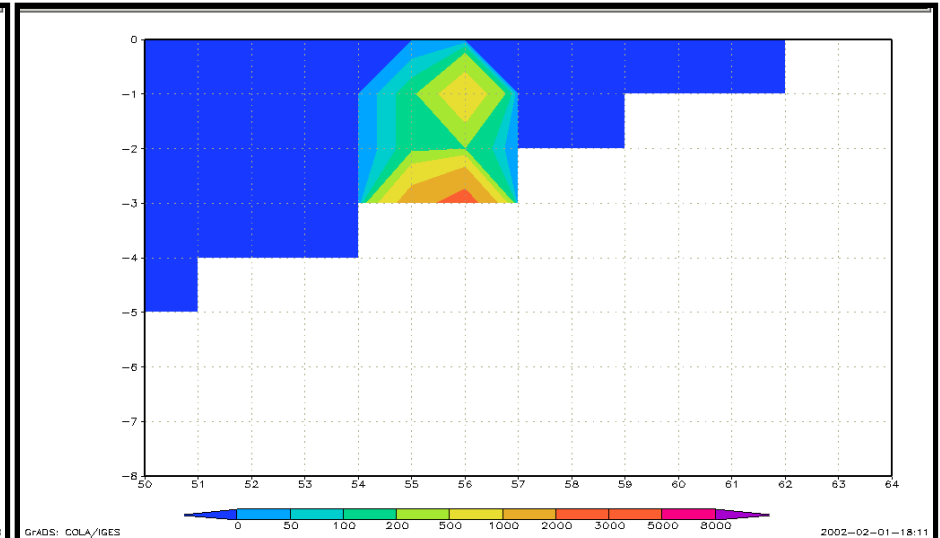
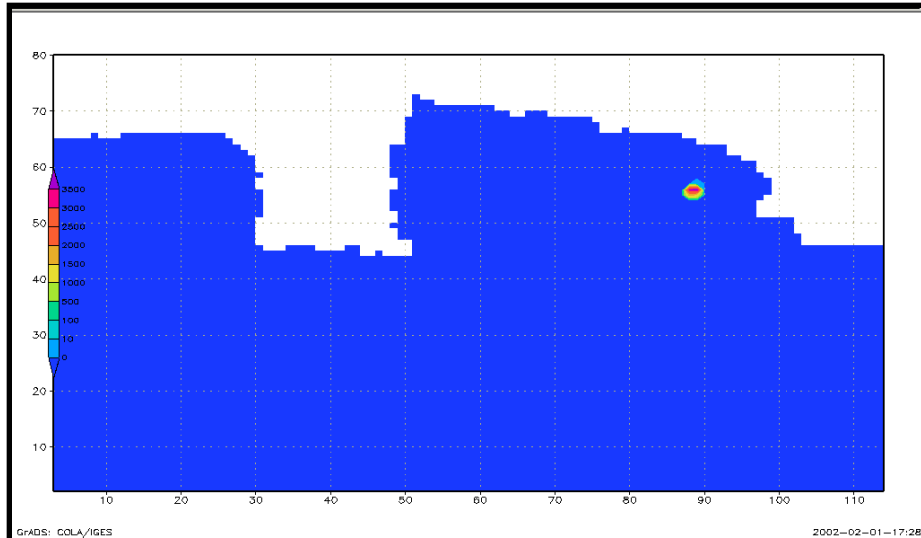
Weights of particles

	Uneated feed % of feed supplied = 5% Periodical release $V_s=0.12$ m/s	Faecal pellet Balance of mass $F_c = 1.3$ kg pellet/kg fish Continuous release $V_s=0.04$ m/s	Dissolved matter Balance of mass $F_c = 1.3$ kg pellet/kg fish Continuous release $V_s=0.$ m/s
N	% in feed = 6.6% <i>1 particle = 2.35 g</i>	% particulated = 22% <i>1 particle = 0.932 g</i>	% dissolved = 78% <i>1 particle = 3.32 g</i>
P	% in feed = 1.35% <i>1 particle = 0.48 g</i>	% particulated = 79% <i>1 particle = 0.814 g</i>	% dissolved = 21% <i>1 particle = 0.22 g</i>
C	% in feed = 45% <i>1 particle = 16 g</i>	Faecal production = 1.9 g/kg fish % organic carbon= 28% <i>1 particle = 0.04 g</i>	----

Dissolved nutrients



Particulated matter

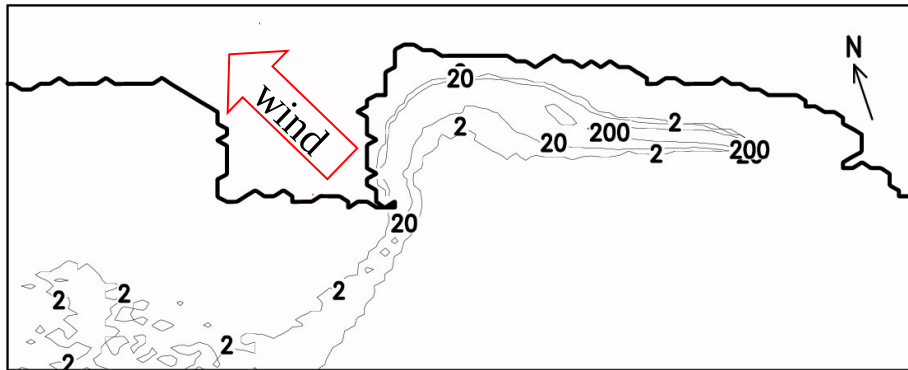
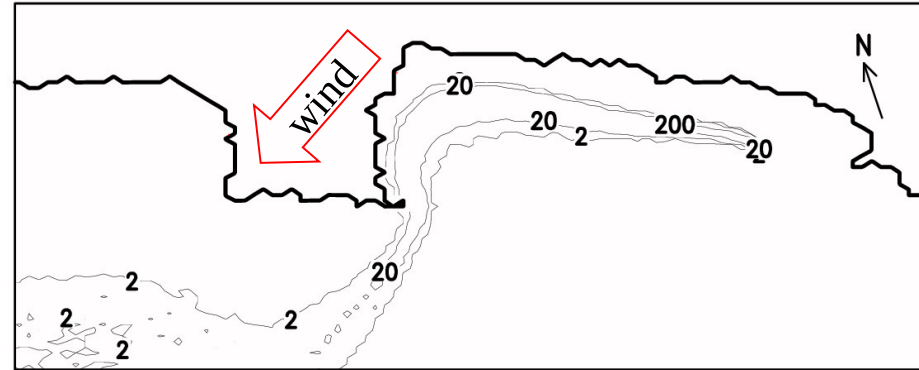


Dissolved nutrients

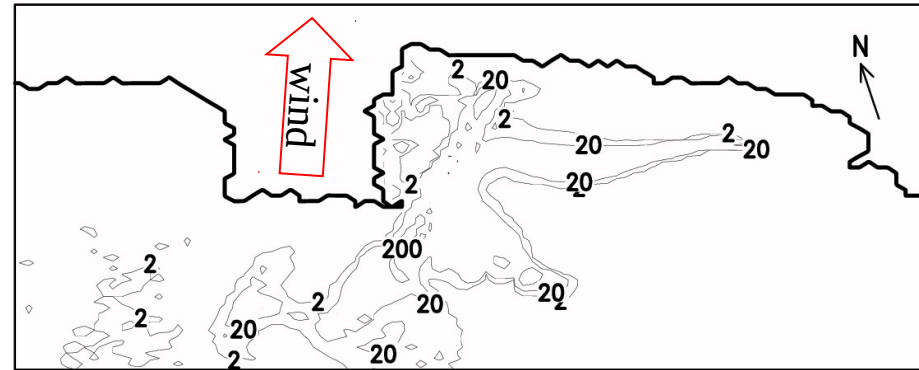
LEGEND:

Number of particles / grid cell

200 = $\begin{cases} 0.83 \mu\text{gN/l} \\ 0.05 \mu\text{gP/l} \end{cases}$



5 km



5 km

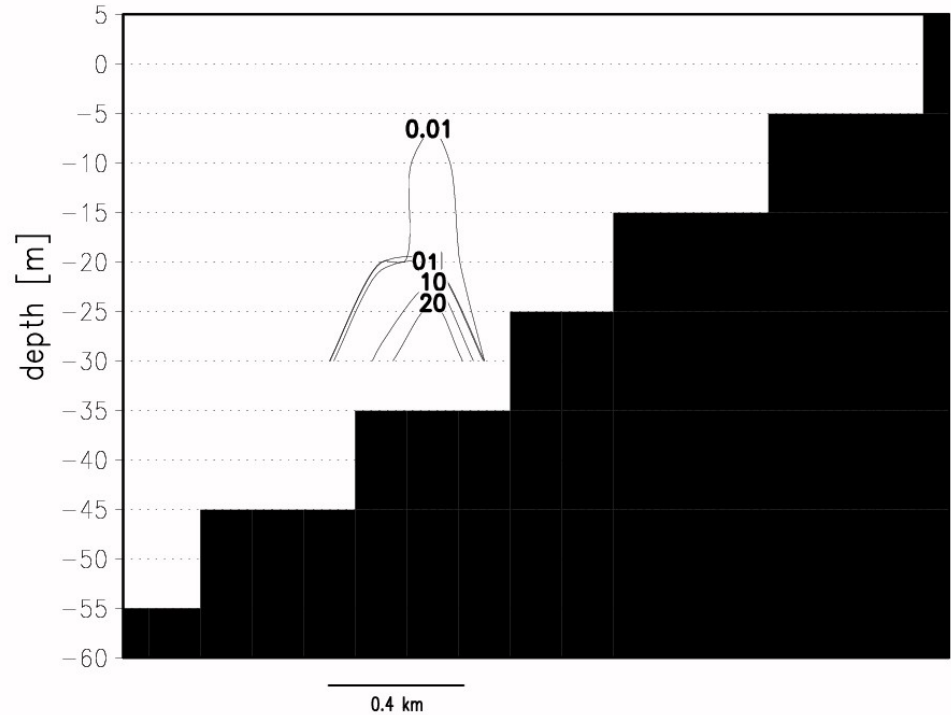
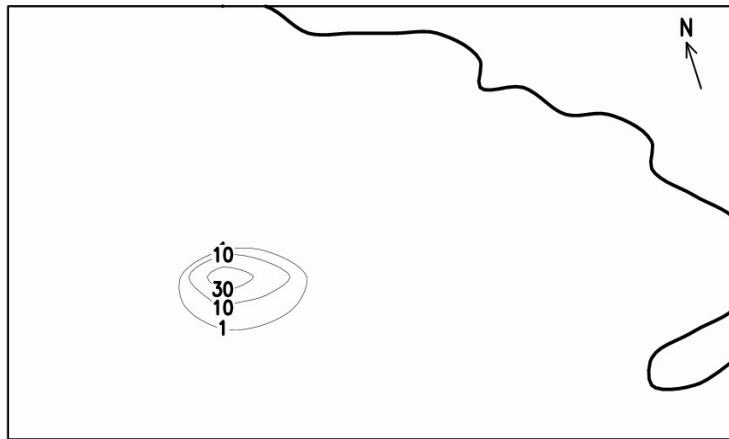
- *Important role of headland;*
- *Best dilution with NE wind;*
- *Nutrients trapped in eddies with SSW wind.*

Particulated matter

LEGEND:

Number of particle /grid cell $\times 10^3$

30 = 0.6 gC/m²



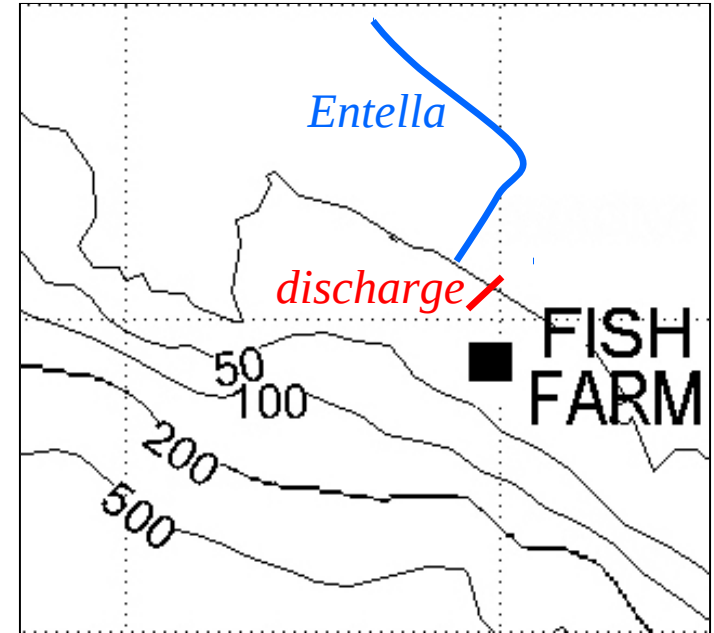
- *Sinking velocity comparable to local current velocity;*
- *The sedimentable matter of any origin remains in the zone of sea cages;*
- *Bell-shaped sedimentation.*
- *Maximum loading rate calculated for organic carbon by the model: **0.085 gC/m²/day***

Comparison with data

DISSOLVED MATTER:

data values >> model output

*Actual setup do not take into account the pre-existing environmental concentration (role of the **river Entella** and the **sewage discharge** from the treatment plant of Lavagna)*



PARTICULATED MATTER:

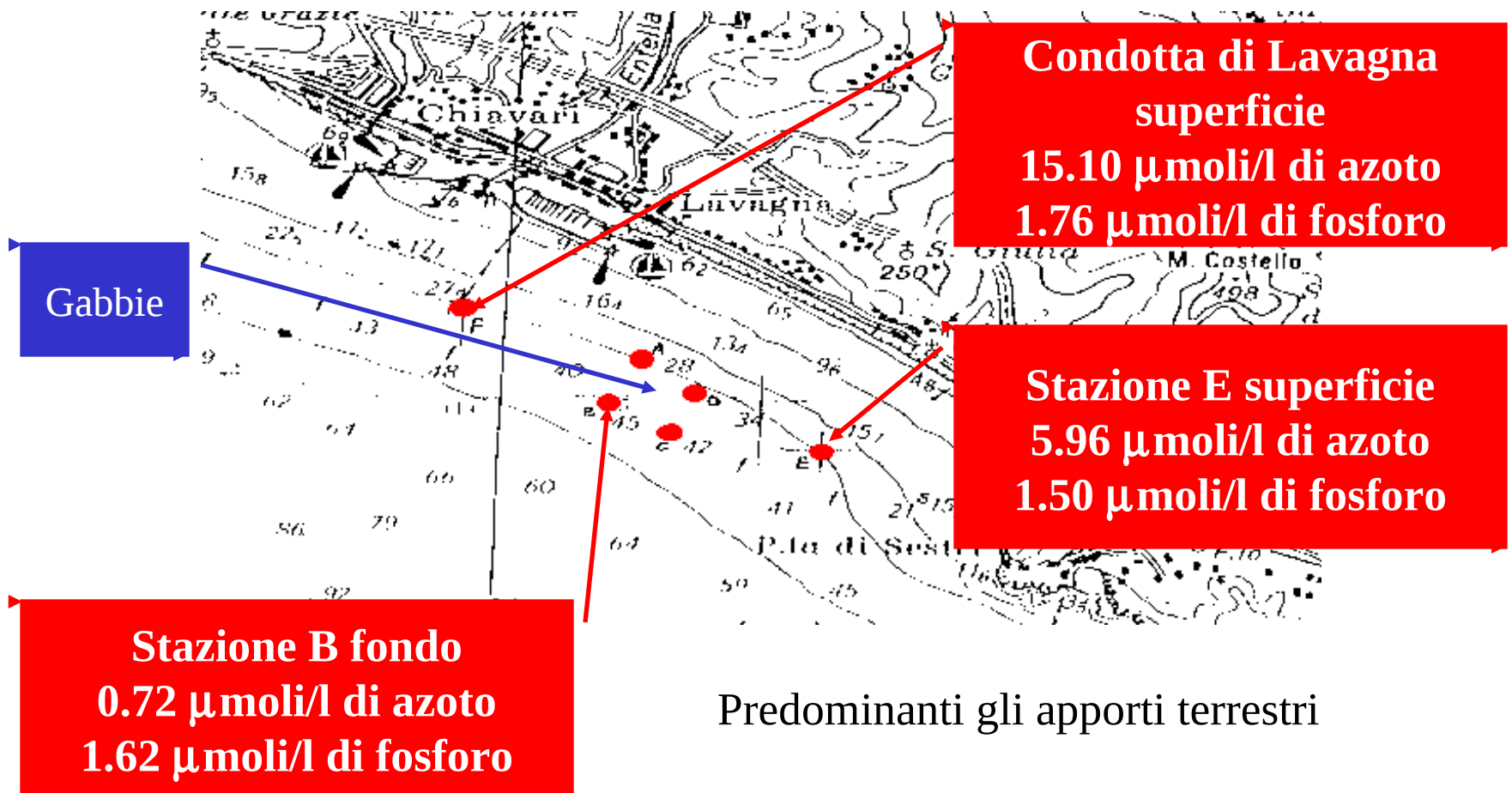
redox potential values indicate an impact restricted to fish farm area

Agreement with the model

Simulazioni avvettivo-dispersive

Le concentrazioni non superano i $4 \mu\text{g/l}$ per l'azoto cioè i $0.065 \mu\text{moli/l}$

Le concentrazioni non superano i $1 \mu\text{g/l}$ per il fosforo cioè i $0.032 \mu\text{moli/l}$



Conclusions

Concentration predicted by the model:

- are not in contrast with the experimental data;
- result low in investigated area in both water column and sediment compartments;
- never exceed the predicted amounts warned in environmental risk assesment.

Outlooks

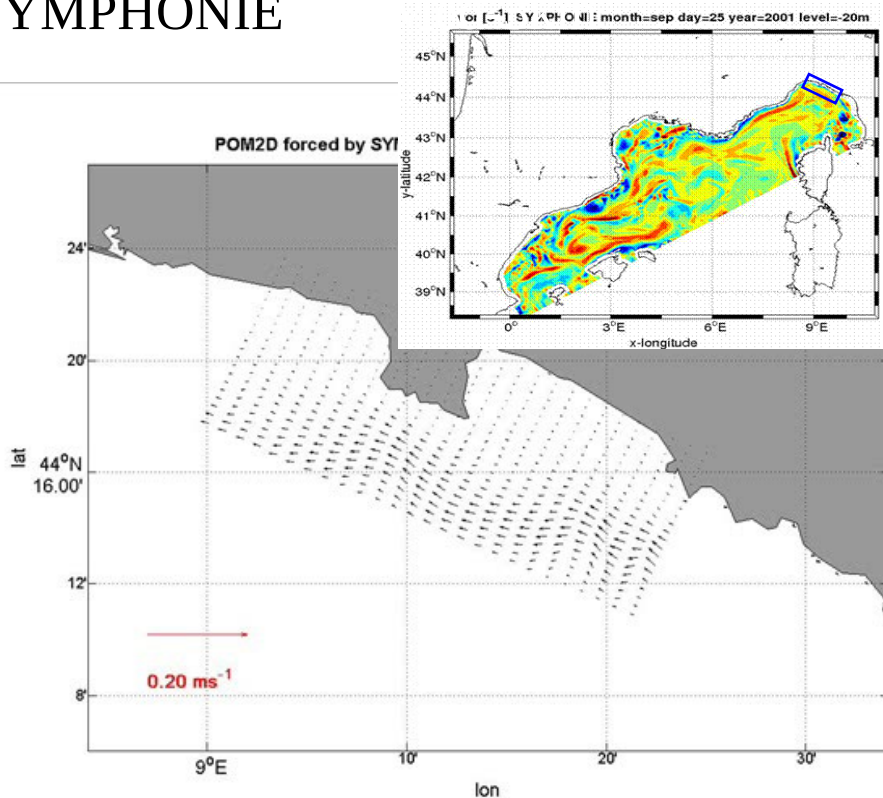
- Model validation with data on a regional scale.
- New experimental monitoring strategy both in water and sediment compartment.
- Implementation of resuspension dynamics.

Travail en cours

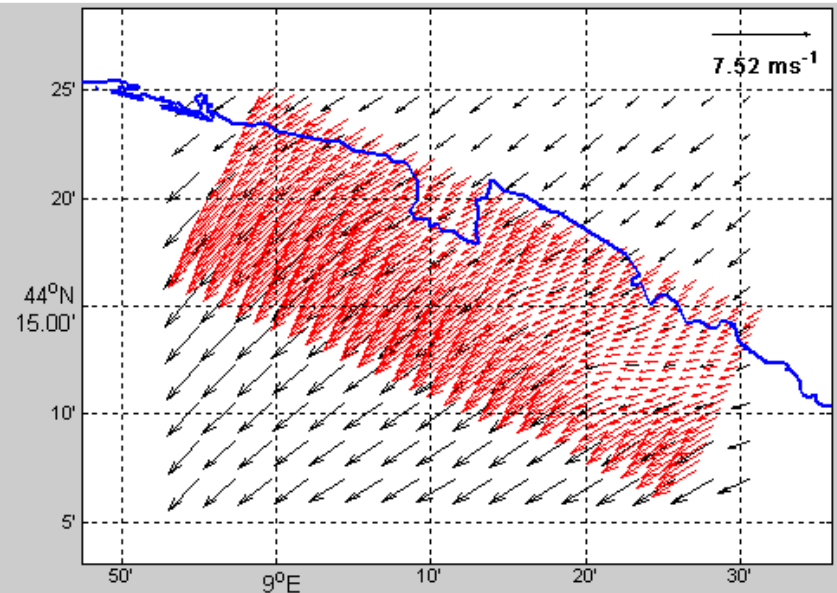
en collaboration avec Departement de Physique-Université de Gênes

Afin d'une simulation réaliste de la circulation, est-il important forcer le model côtier avec un vent à très haute résolution en région à topographie complexe?

coastal model (code POM) forced by SYMPHONIE



WINDS (Wind-field Interpolation by Non-Divergent Schemes) is a diagnostic mass-consistent model for simulation of the three-dimensional wind field in complex terrain at mesoscale



PROJET LATEX (LAgrangian Transport Experiment)

Objectif: rôle de la dynamique couplée physique - biogéochimie à (sub) méso-échelle dans les échanges côte-large

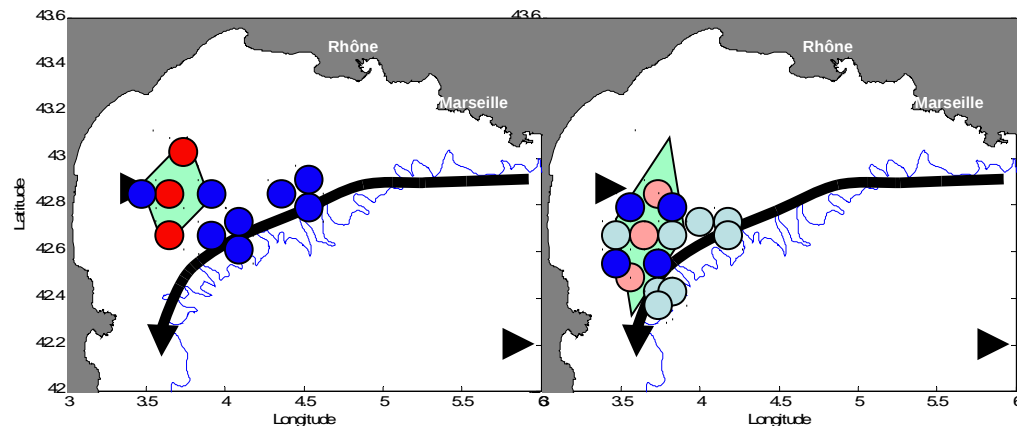
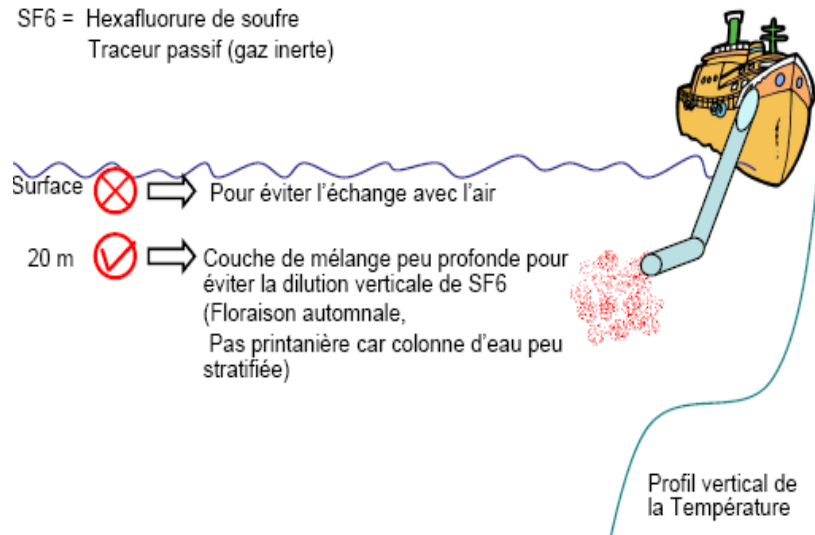
Méthodologie: utiliser une démarche lagrangienne pour le suivi d'une structure tourbillonnaire de (sub)méso-échelle marquée avec un traceur chimique inerte (SF_6)

LATEX = LATEX00 (projet pilote) + LATEX

→
2007

→
2008-2010

SF6 = Hexafluorure de soufre
Traceur passif (gaz inerte)



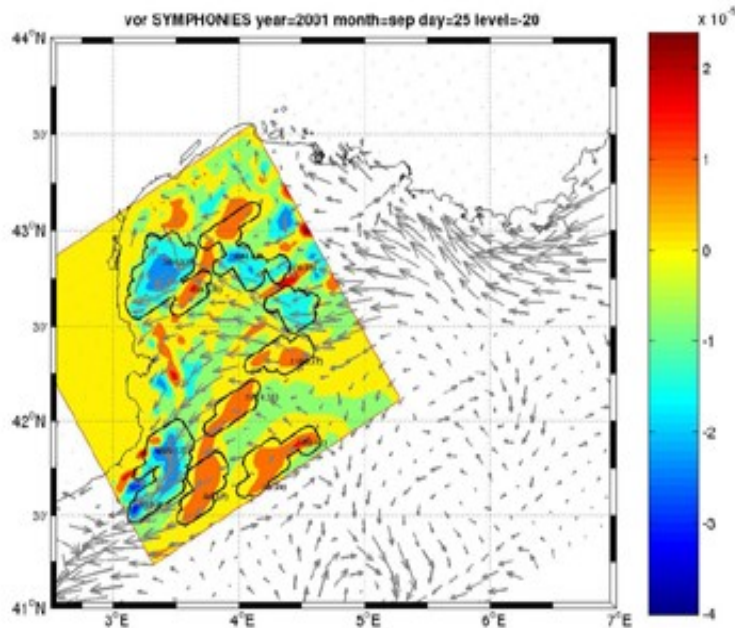
LATEX = projet pilote LATEX00 + projet LATEX

LATEX00 (2007)

Approfondissement des connaissances sur les processus (sub)méso-échelle et mise à point de la stratégie de mesure

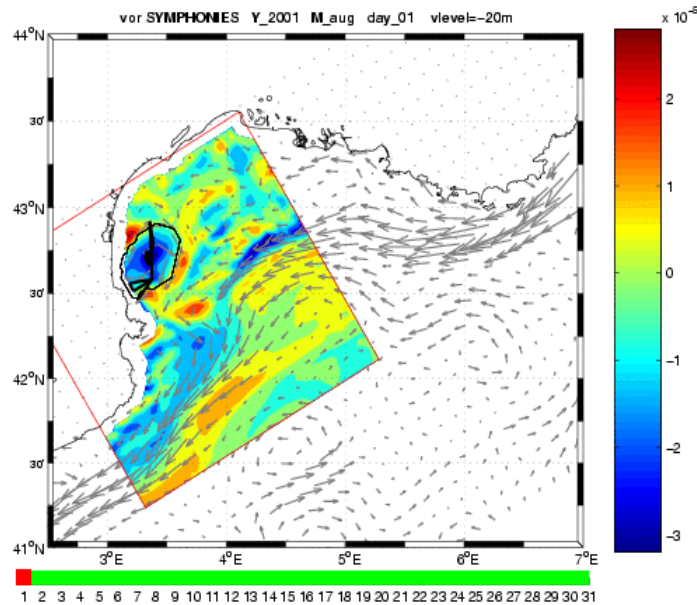
modèle SYMPHONIE +
Analyse en ondelettes +
Flotteurs lagrangiens

Campagne en mer
tests de navigation lagrangienne +
mesures de C_{amb} du SF₆

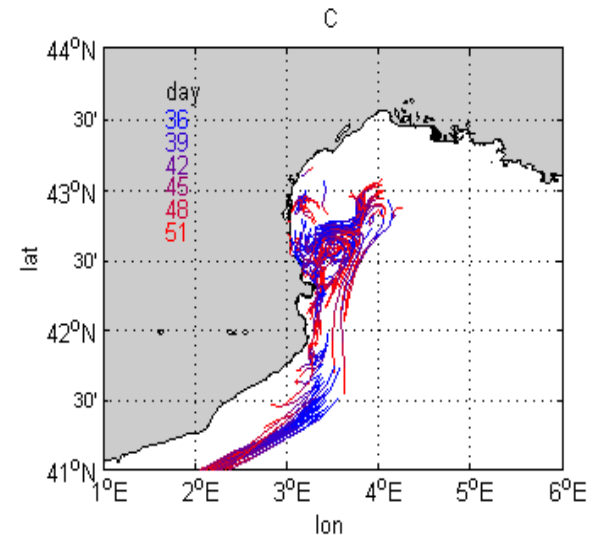
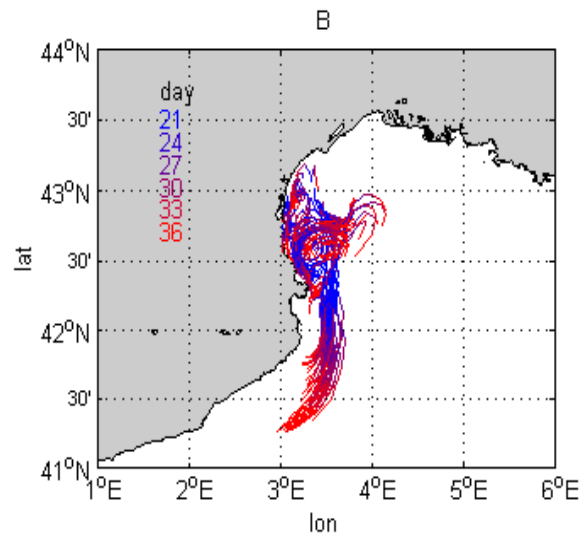
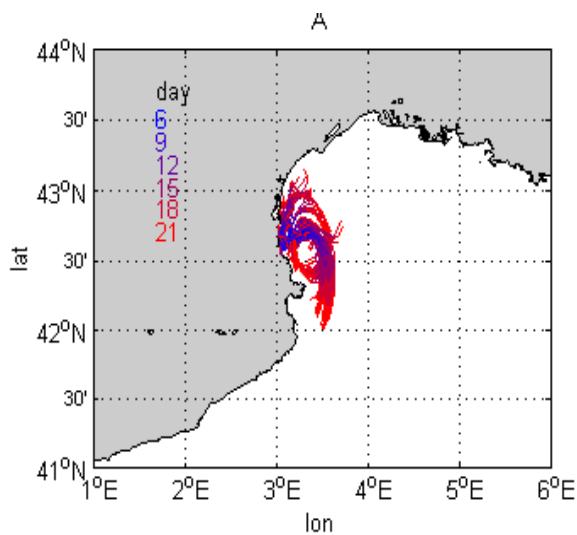


Tourbillon anticyclonique A1(01)

durée: 32 jours animation montrée sur la dernière partie (1er – 18



Site web ZiYuan Hu, master 2007
<http://www.com.univ-mrs.fr/~h307258/>



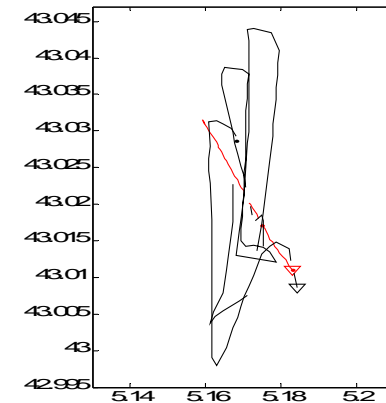
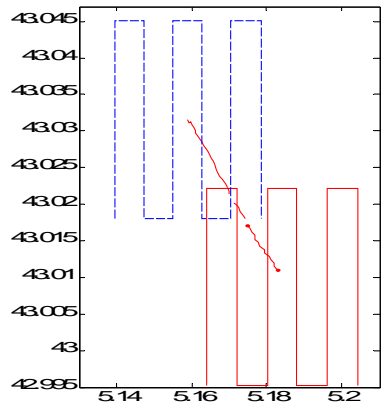
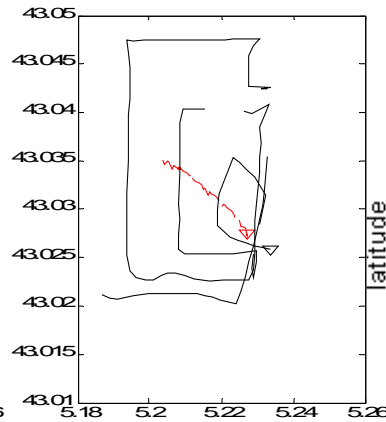
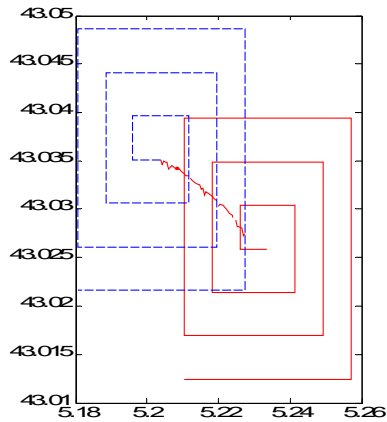
Navigation lagrangienne pour la mise à l'eau du traceur

Résultats des tests, en mer, du logiciel de nav. lagrangienne

LATEX 00 9-11 juin 2007

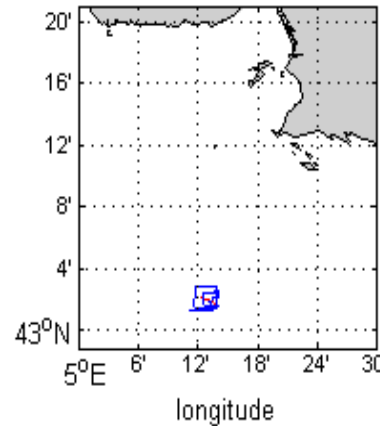
théorique

réelle

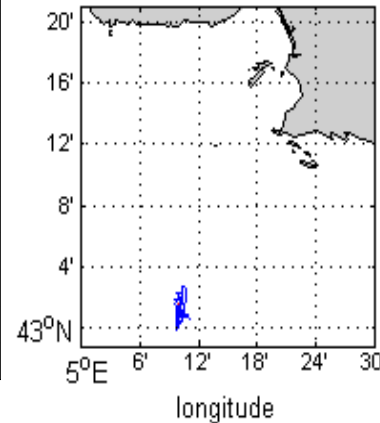


réelle
zoom out

vessel (blue) and buoy (red) tracks

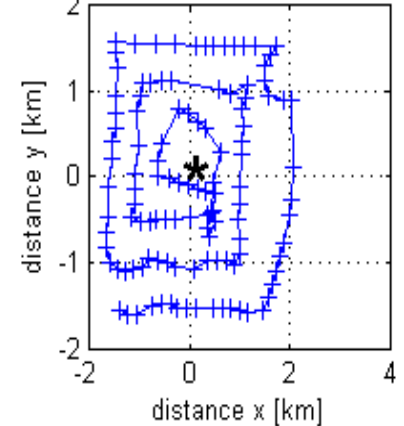


vessel (blue) and buoy (red) tracks

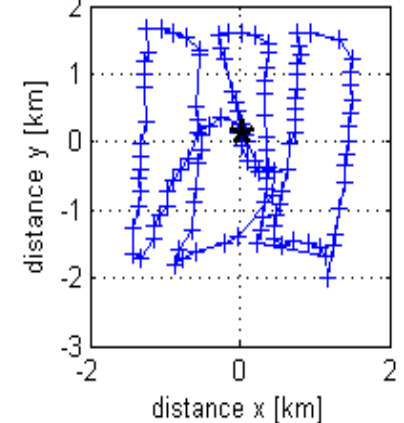


repère lagrangien

vessel track in lagrangian reference

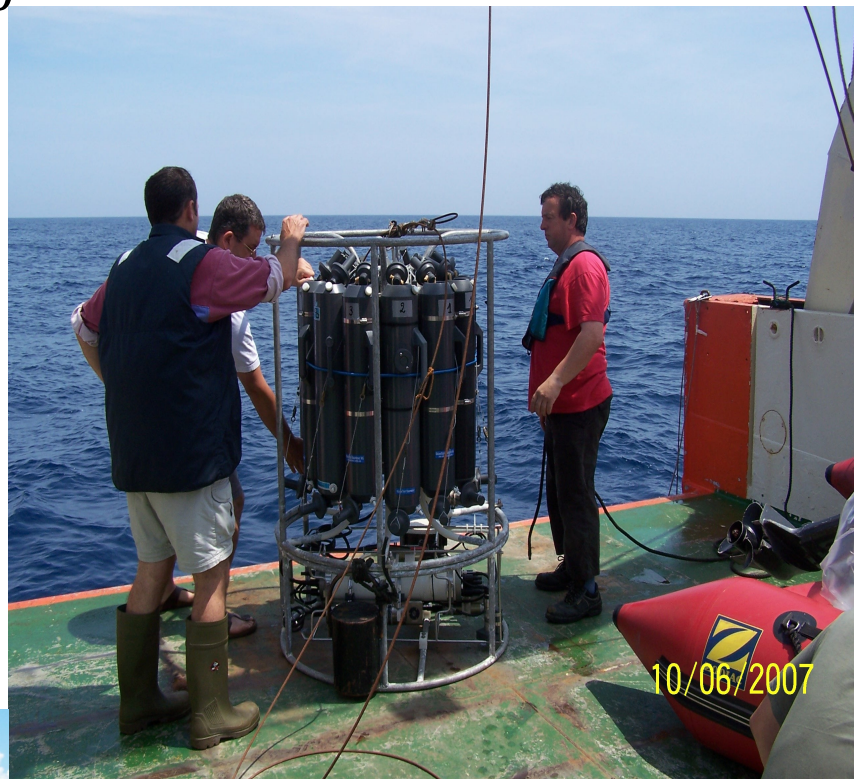


vessel track in lagrangian reference



FINE

LATEX00



POM flow chart

