

of submesoscale horizontal eddy diffusion coefficients across a front

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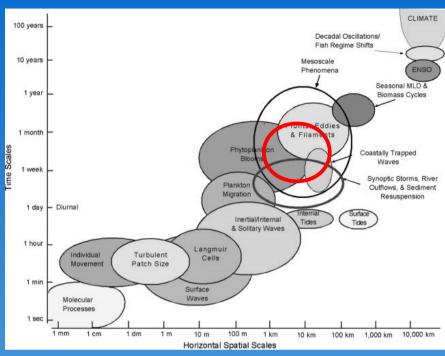








(from http://oceanservice.noaa.gov/)

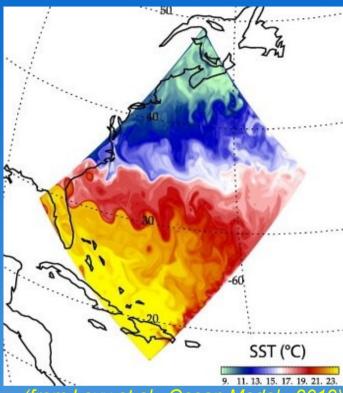


(from Dickey et al., J. Mar. Syst, 2003)

- → Fronts, jets and eddies: Horizontal scales ~(100m 10km)
 Time scales ~(days week)
- → Key role for : Energy transfer
 Horizontal and vertical transport
 Biogeochemical cycles



→ Numerical model studies
 Favored by:
 Advances in computational power
 Development of regional models



(from Levy et al., Ocean Model., 2012,



→ Numerical model studies

Favored by:

Advances in computational power Development of regional models

→ In-situ observations

Challenging due to small spatial and temporal scales



SST (°C)

9. 11. 13. 15. 17. 19. 21. 23.

(from Levy et al., Ocean Model., 2012)

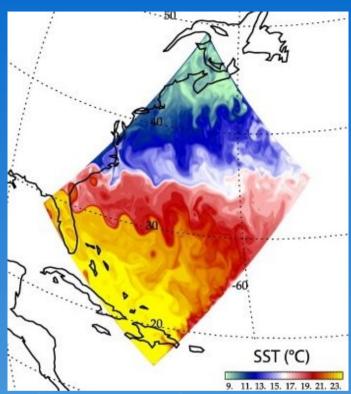
Limited estimates of key physical parameters



- → Numerical model studies
 - Favored by:
 - Advances in computational power Development of regional models

- → In-situ observations
 - Challenging due to small spatial and temporal scales





(from Levy et al., Ocean Model., 2012)

Limited estimates of key physical parameters

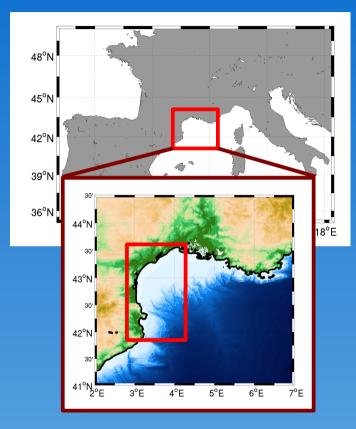
Focus of this study

New approach to estimate horizontal eddy diffusion coefficients (Kh) from in-situ observations

Lagrangian Transport EXperiment



→ Latex10 campaign (September 1-24, 2010)



- Western part of Gulf of Lion (NW Mediterranean)
- Main goals : (Sub)mesoscale dynamics
 Cross-shelf exchanges
- Adaptive sampling strategy based on :
 - Satellite data
 - Ship-based current measurements
 - Iterative Lagrangian drifter releases

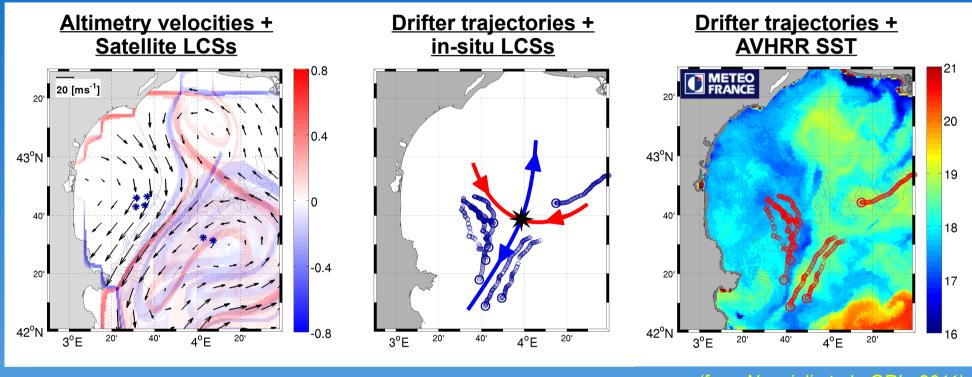


In-situ LCSs



Example:

→ Drifter array deployment "Lyap01" (Sep. 15, 2010)



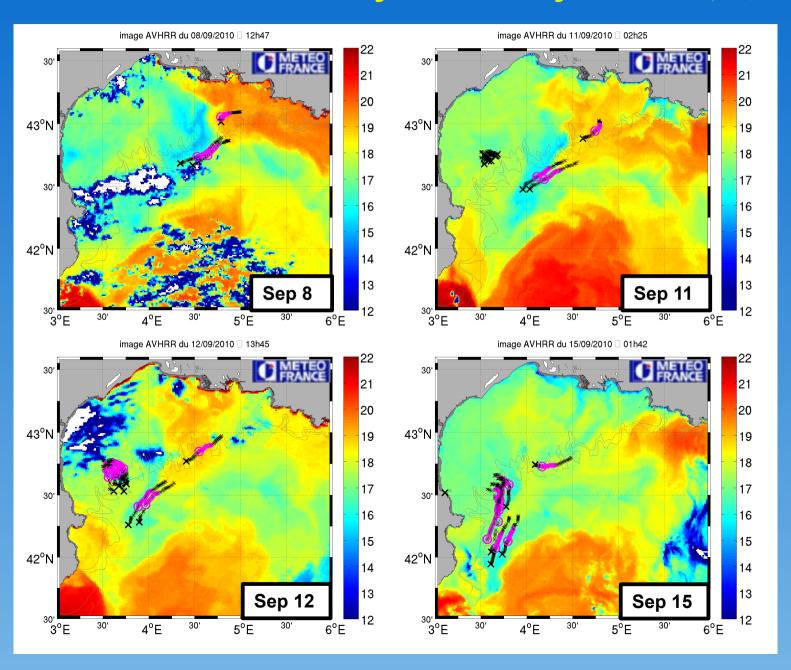
(from Nencioli et al., GRL, 2011)

- Identified in-situ Lagrangian Coherent Structures (LCSs)
- Evidenced inaccuracy of altimetry in coastal region
- LCS associated with a frontal structure

Front development



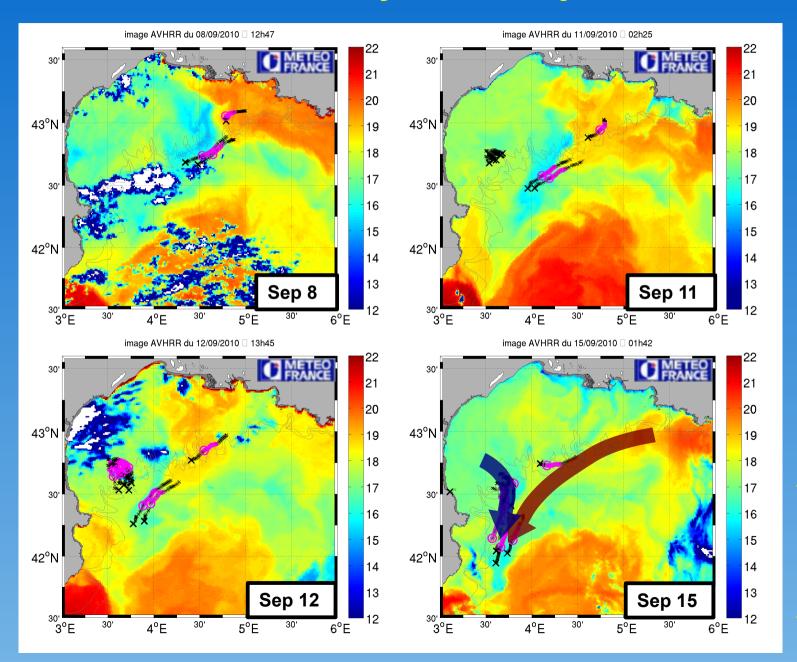
→ AVHRR SST + 3-day drifter trajectories (Sep. 8 to 15, 2010)



Front development



→ AVHRR SST + 3-day drifter trajectories (Sep. 8 to 15, 2010)

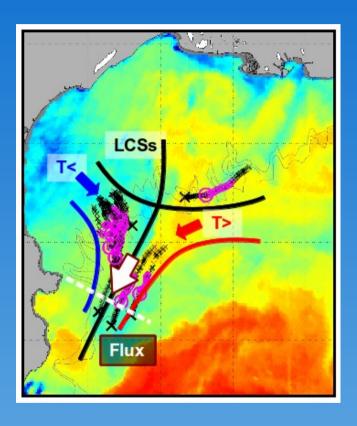


Convergence of warmer (eastern outer shelf) and colder (western inner shelf) water masses

In-situ approach



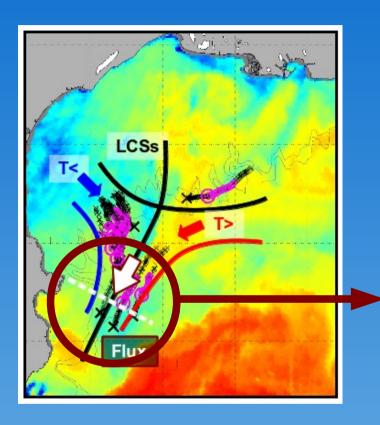
→ Colder and warmer water masses converging along attractive LCS

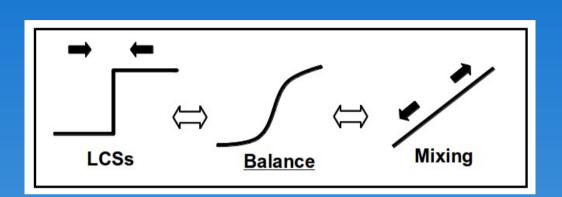


In-situ approach



→ Colder and warmer water masses converging along attractive LCS



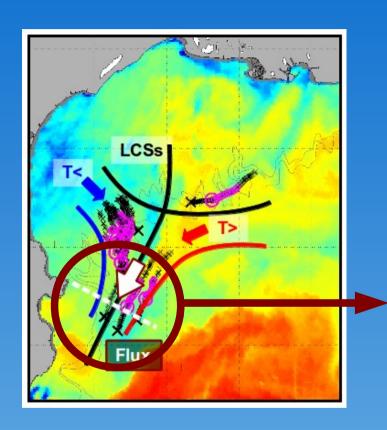


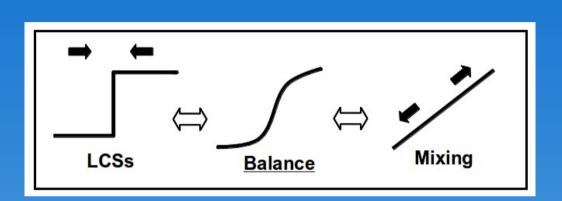
 Shape of T and S profile across the front results from balance between convergence and horizontal mixing

In-situ approach



→ Colder and warmer water masses converging along attractive LCS





 Shape of T and S profile across the front results from balance between convergence and horizontal mixing

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$

 Analytical solution to 1D advection diffusion equation for a tracer T

→ Analogous to Flament et al. 1985, Ledwell et al. 1998 (from satellite)

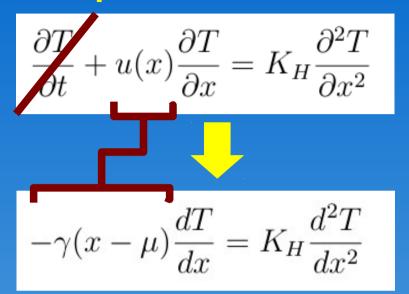


1D equation for a tracer T

$$\frac{\partial T}{\partial t} + u(x)\frac{\partial T}{\partial x} = K_H \frac{\partial^2 T}{\partial x^2}$$



1D equation for a tracer T



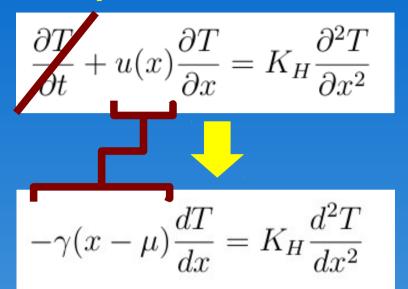
Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction

with



1D equation for a tracer T

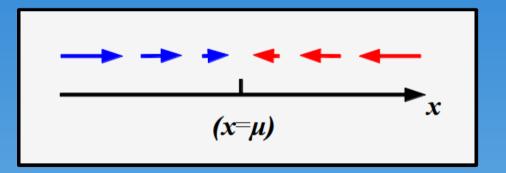


Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction

 γ : Strain rate (Lyapunov exponent)

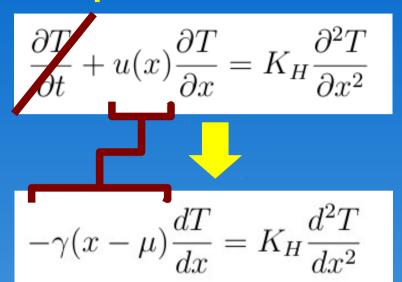
 μ : Position of LCS axis



with



1D equation for a tracer T



Boundary Conditions

$$T(x = -\infty) = T_1;$$

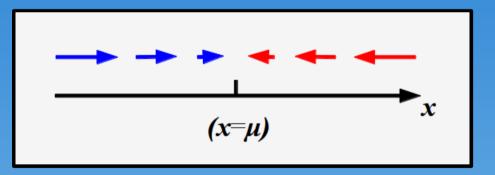
$$T(x = +\infty) = T_2;$$

Assumptions:

- Front is at equilibrium (steady state)
- x is the across LCS direction

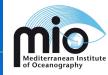
 γ : Strain rate (Lyapunov exponent)

 μ : Position of LCS axis

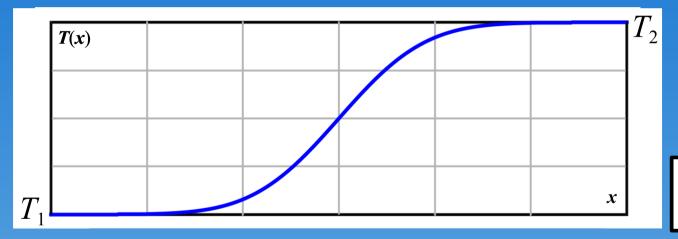


$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$

Analytical solution

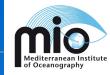


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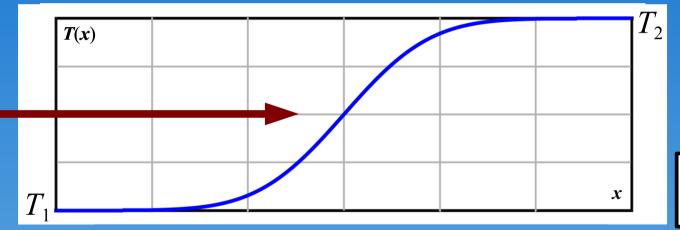
$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} dt$$

Analytical solution



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}}\sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$

C1



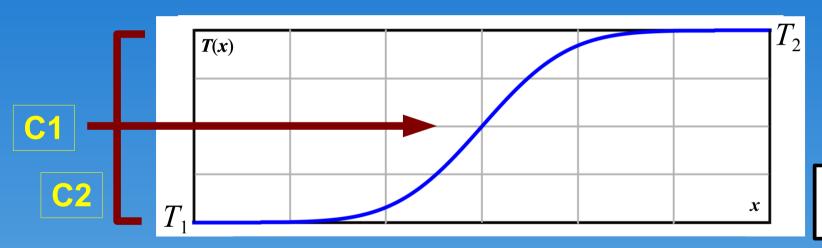
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C1 C2



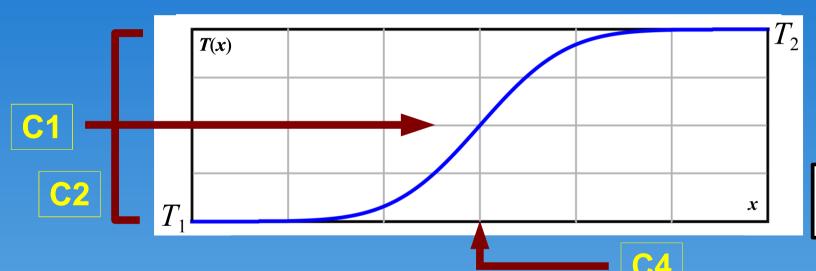
$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} \mathrm{d}t$$

Analytical solution



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$

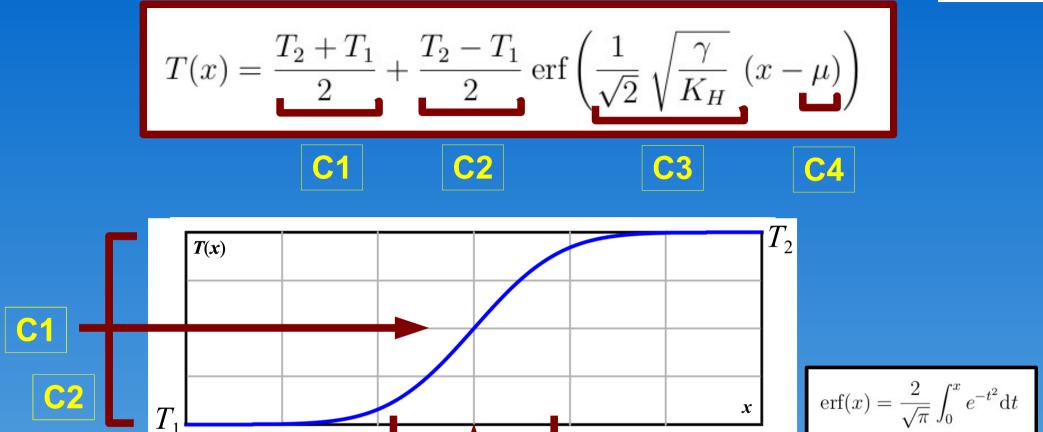
C1 C2 C4



$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} dt$$

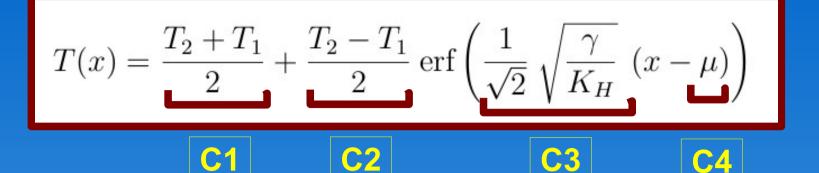
Analytical solution





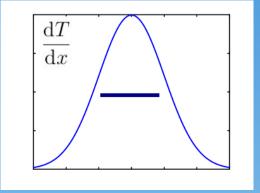
Analytical solution







$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} dt$$



$$W_{front} = 2\sqrt{\frac{K_H}{\gamma}}$$

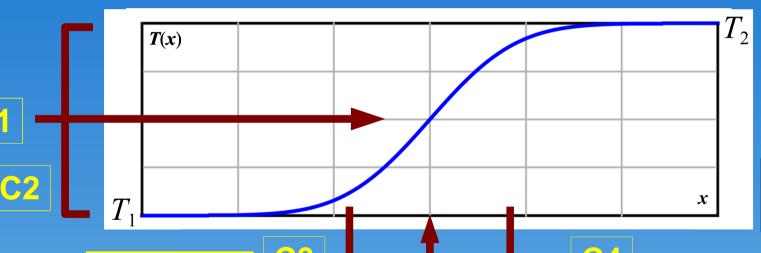
68% of T

Analytical solution



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$

C1 C2 C3 C4



$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} dt$$

 $\frac{\mathrm{d}T}{\mathrm{d}x}$

$$W_{front} = 2\sqrt{\frac{K_H}{\gamma}}$$

68% of T variation

 Coefficients computed by best fitting T and S sections

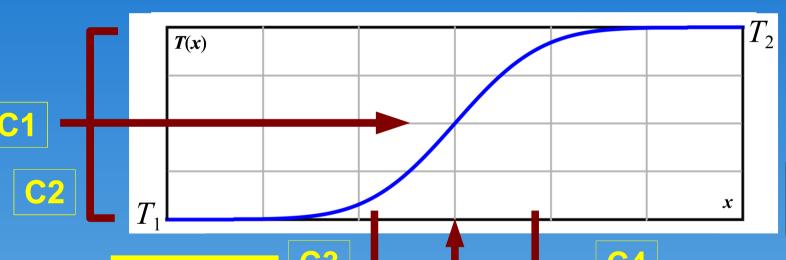
$$K_H = \frac{\gamma}{\left(2 \, \mathrm{C} 3^2\right)}$$

Analytical solution



$$T(x) = \frac{T_2 + T_1}{2} + \frac{T_2 - T_1}{2} \operatorname{erf}\left(\frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}} (x - \mu)\right)$$

C1 C2 C3 C4



$$\operatorname{erf}(x) = \frac{2}{\sqrt{\pi}} \int_0^x e^{-t^2} dt$$

 $\frac{\mathrm{d}T}{\mathrm{d}x}$

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68% of T variation

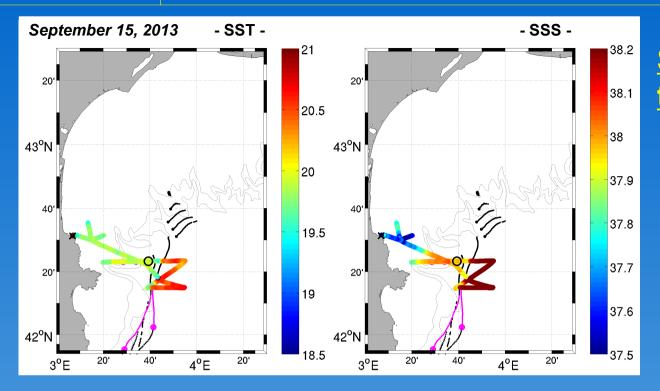
 Coefficients computed by best fitting T and S sections

$$K_H = rac{\gamma}{\left(2 \, \mathrm{C} 3^2\right)}$$
 from drifter dispersion!

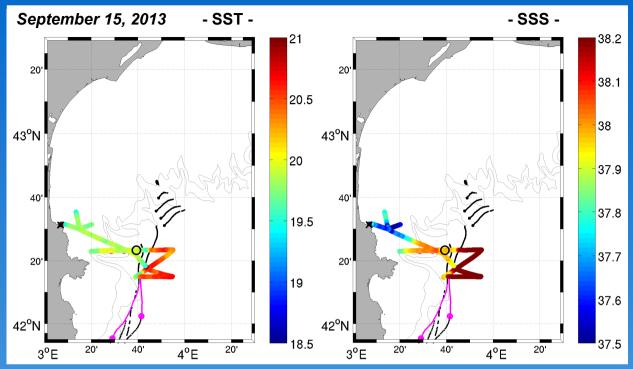
4. Kh

Cross-front transects

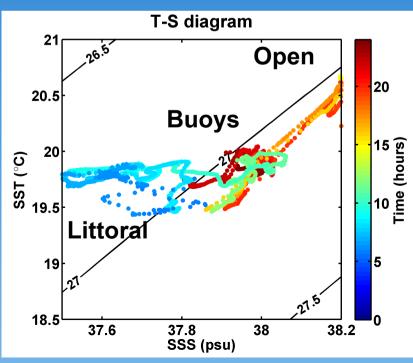




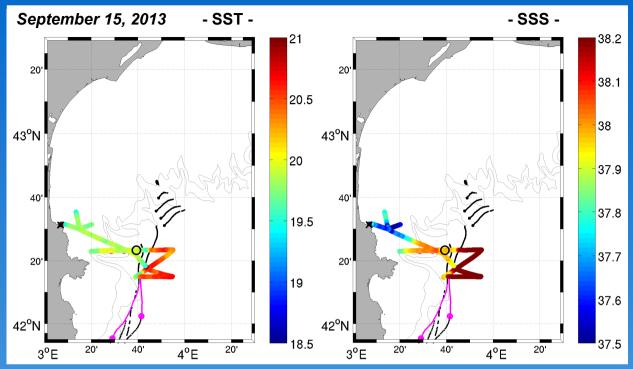




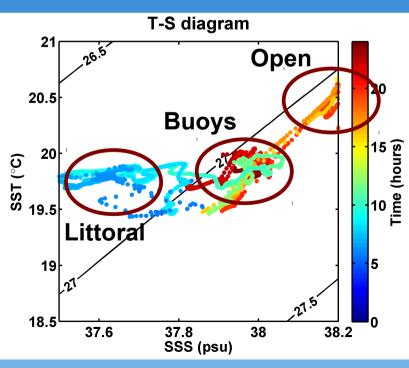




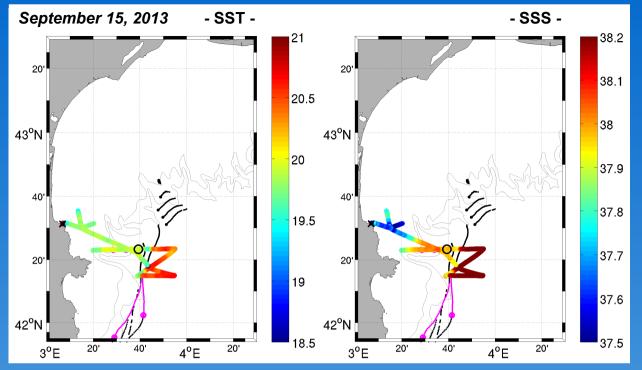


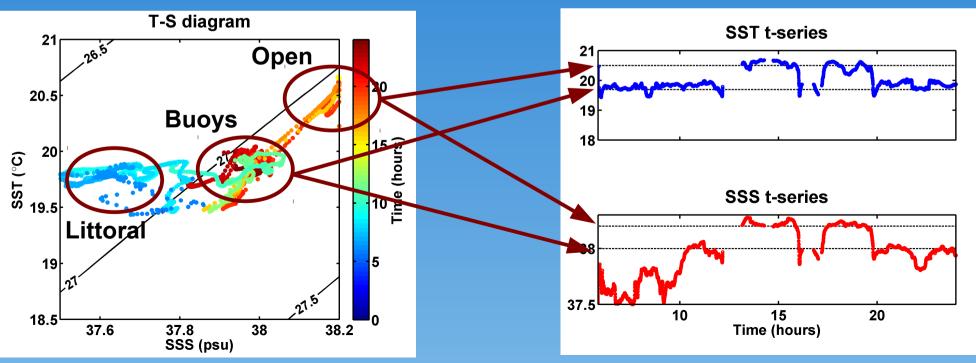




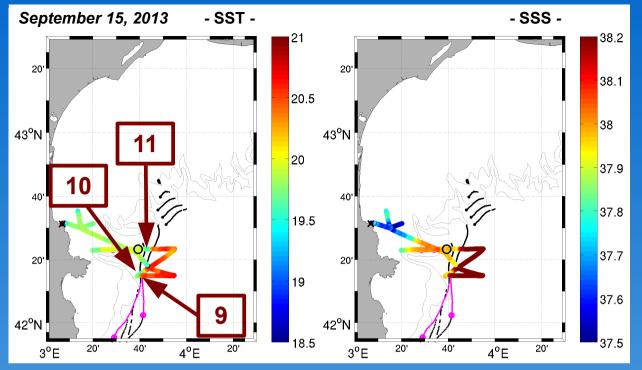


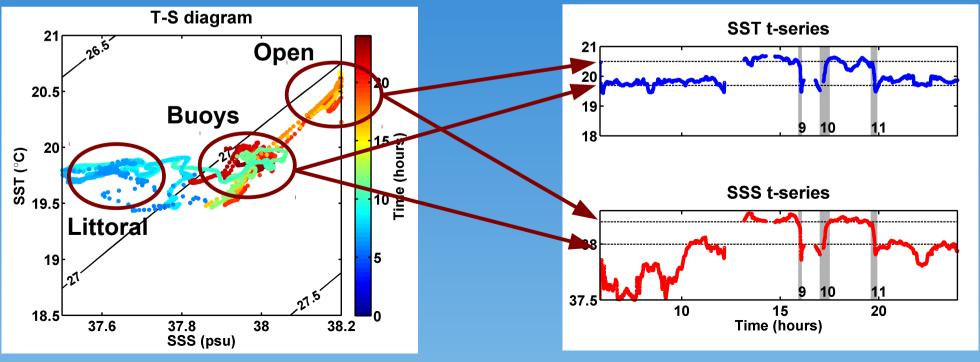




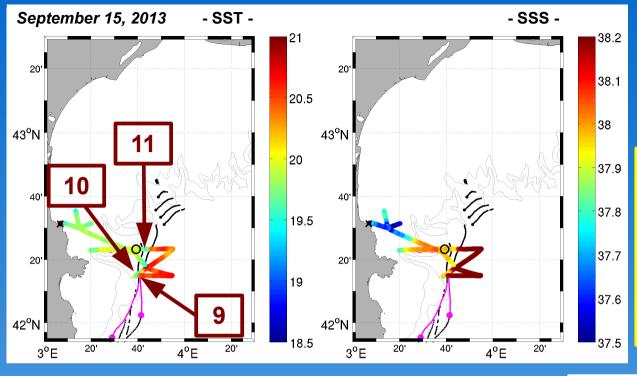




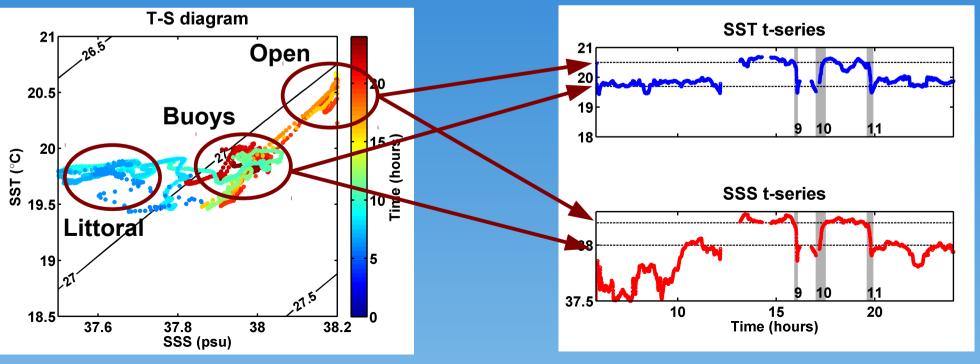






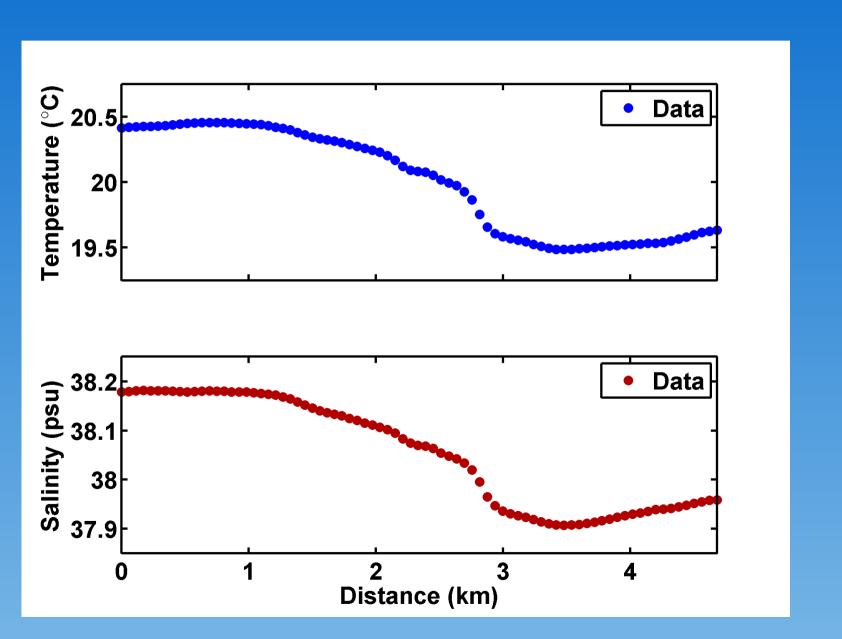


- Identified 30 cross-front transects
- Transects projected on the direction normal to the LCS axis



Mediterranean Institut

Example: Transect 11

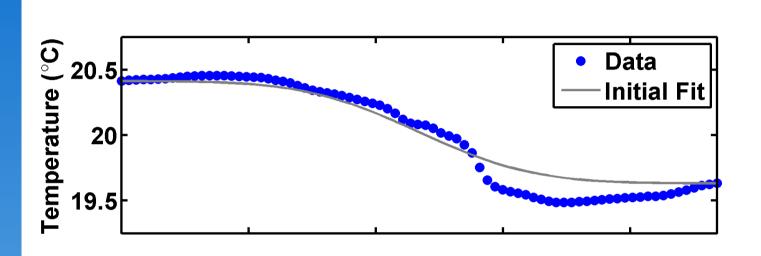


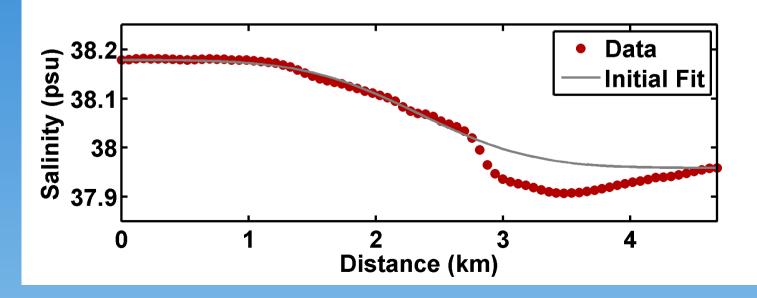
Mediterranean Institute of Oceanography

Example: Transect 11

Front equation

$$T(x) = C_1 + C_2 \operatorname{erf}(C_3 (x - C_4))$$





Initial fit

$$C_1 = \frac{T_2 + T_1}{2}$$

$$C_2 = \frac{T_2 - T_1}{2}$$

$$C_3 = 1$$

$$C_4 = \frac{Dist}{2}$$

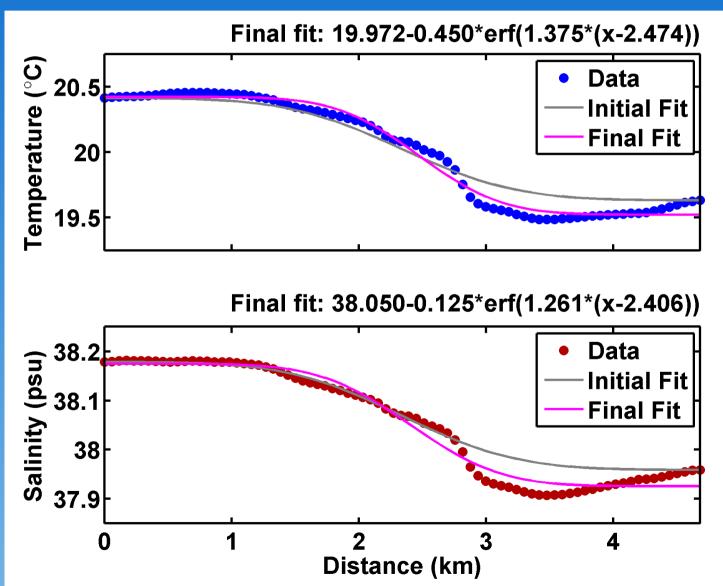


Example: Transect 11

Parameters from least square estimation using Nelder-Mead simplex direct search

Front equation

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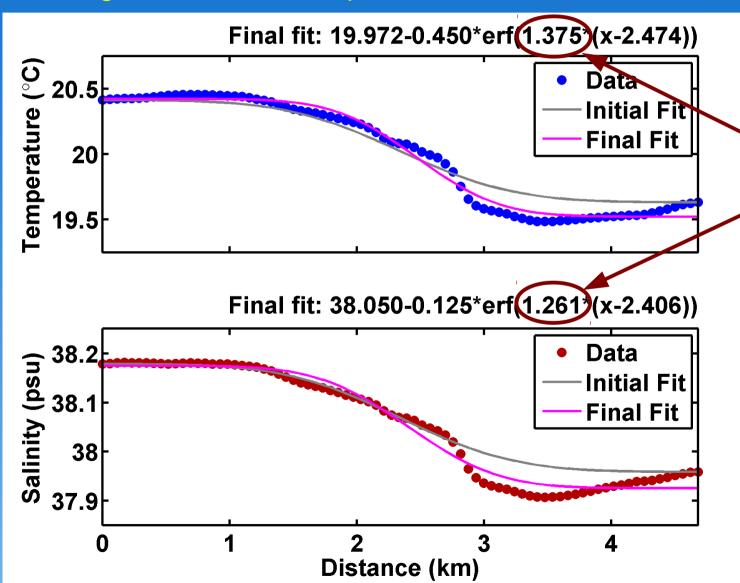


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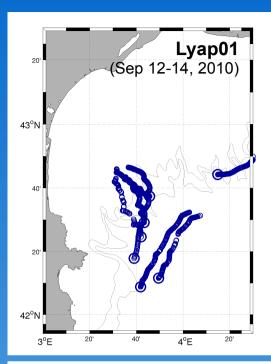
$$C_3 = \frac{1}{\sqrt{2}} \sqrt{\frac{\gamma}{K_H}}$$

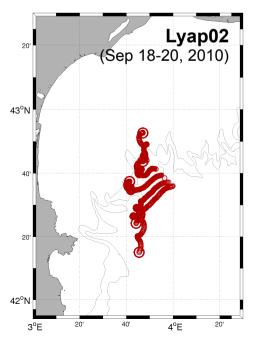
 No fit for 11 out of 30 transects: limits of starting assumptions

i.e. impact of vertical dynamics

Strain rate



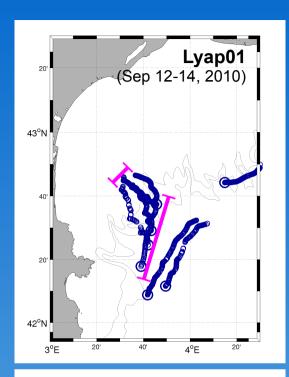


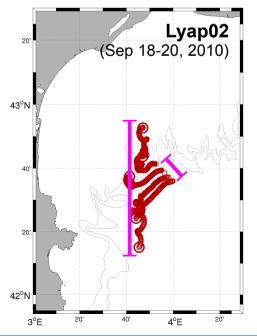


Dispersion patterns of drifter arrays

Strain rate



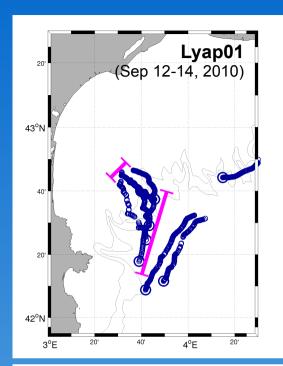


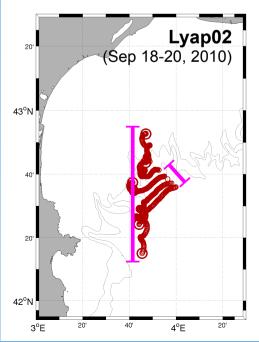


Dispersion patterns of drifter arrays

 For each deployment, computed fastest separation rate between buoy couples (analogous to Lyapunov exponent)

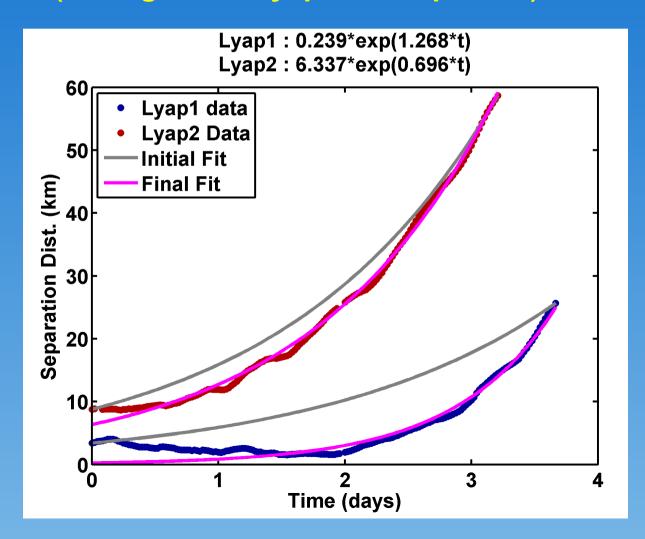




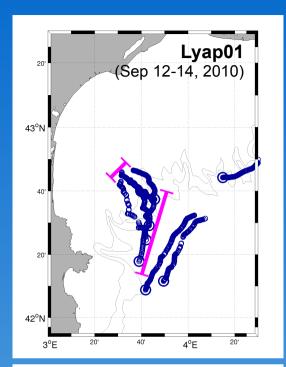


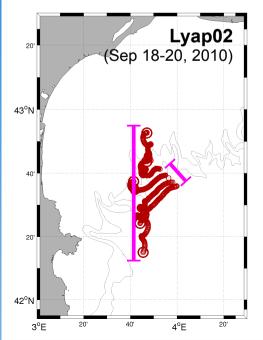
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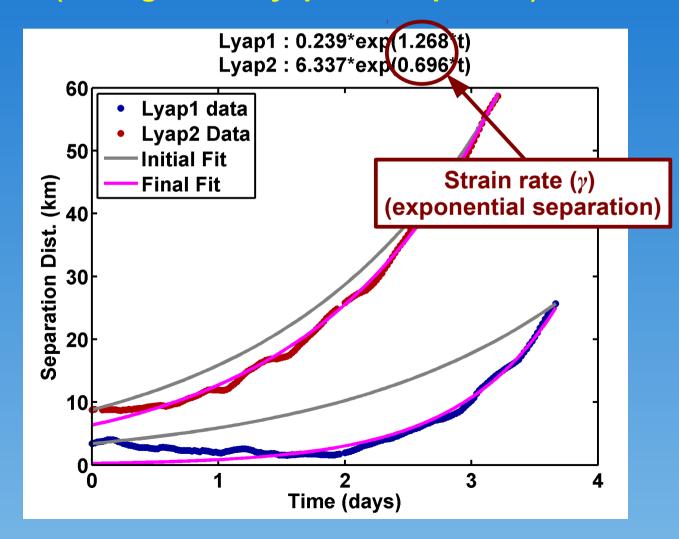






Dispersion patterns of drifter arrays

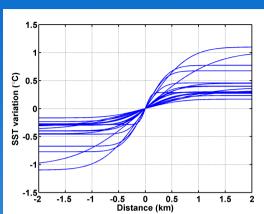
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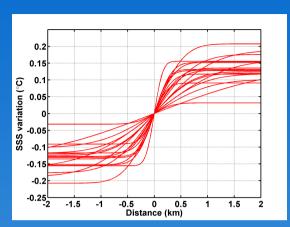
Eddy diffusivity coefficients



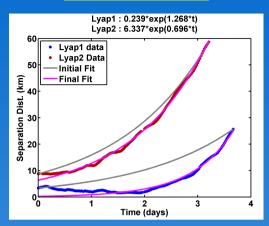
T Front



S Front

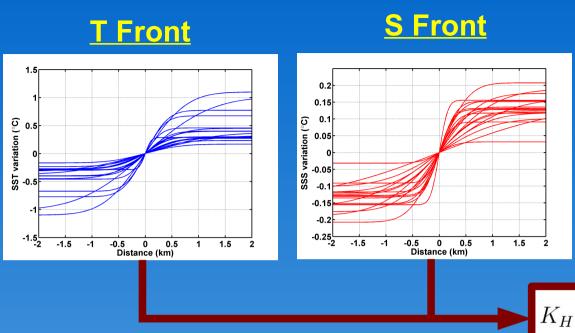


Strain rate

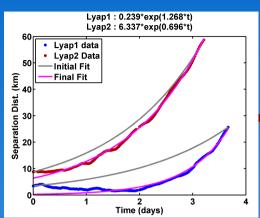


Eddy diffusivity coefficients

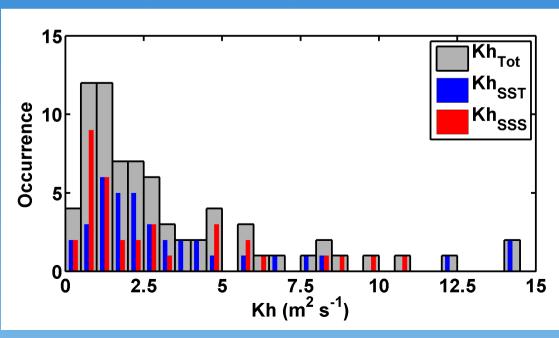


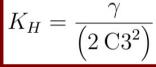


Strain rate



Eddy diffusivity coefficients





variation (°C)

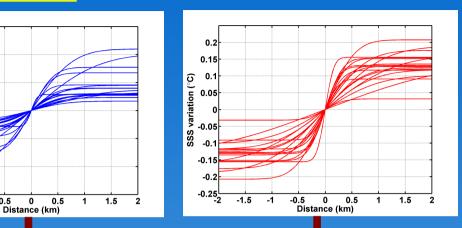
Eddy diffusivity coefficients

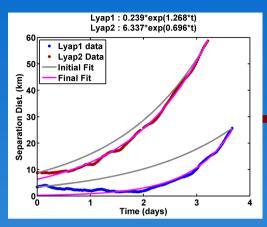






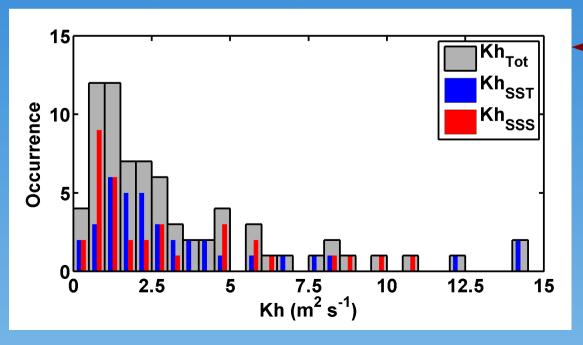
Strain rate





 $K_H = \frac{\gamma}{\left(2 \, \mathrm{C} 3^2\right)}$

Eddy diffusivity coefficients



- Kh_{sst} similar to Kh_{sss}
- 75% of estimates between 0.5 - 5 m s⁻²
- Front widths range from 250 m to 2km
- Long tail

Conclusions



→ New approach relatively simple and cheap (i.e. compared to passive tracer release experiments)

→ In-situ estimates of Kh at the submesoscale in line with values used in high-resolution numerical models

- → Tail of high values affects Kh statistics (mean and standard deviation) => check starting assumptions:
 - Steady state
 - Uniform strain rate
 - Vertical motions
 - 0

Perspectives



- → Further dedicated in-situ experiments
- → Test approach from high-resolution models
- → Extend analysis of Kh over wider regions/the global ocean using remote sensed datasets

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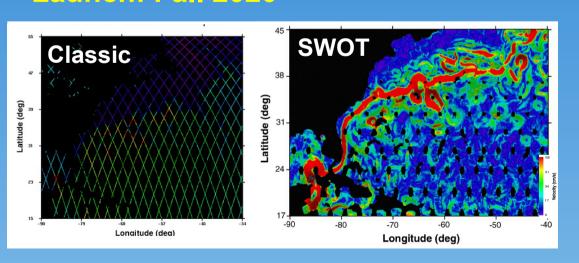




Surface Water and Ocean Topography

NASA - CNES mission

- New generation, high-resolution (1Km) altimeter
- Launch: Fall 2020



References



This work has been developed within the project:



Lyapunov Analysis in the CoaSTal Environment (LACOSTE)

Marie Curie Intra-European Fellowship Call: FP7 - PEOPLE - 2011 - IEF Leading PI: F. Nencioli





F. Nencioli, F. d'Ovidio, A. Doglioli, A. Petrenko Surface coastal circulation patterns by in-situ detection of Lagrangian Coherent Structures. Geophysical Research Letters, 38, L17604, doi:10.1029/2011GL048815

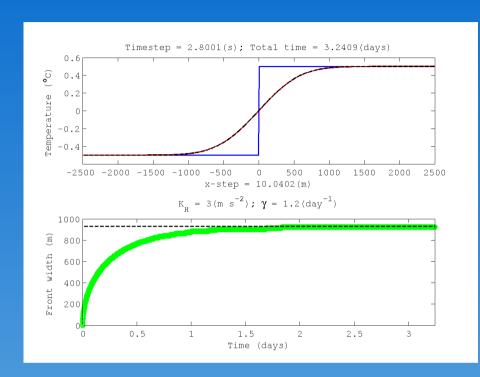
LATEX website: www.com.univ-mrs.fr/LOPB/LATEX

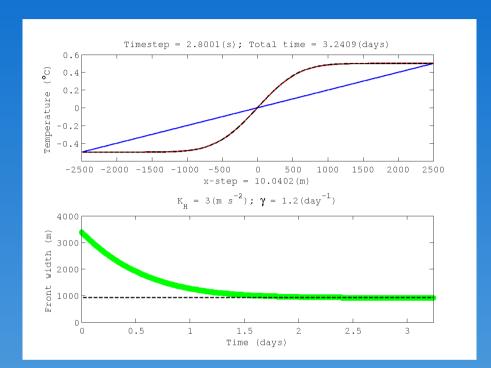
EXTRA SLIDES

Numerical solution



First order upwind scheme

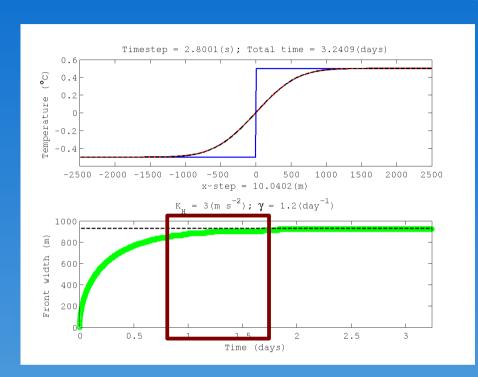


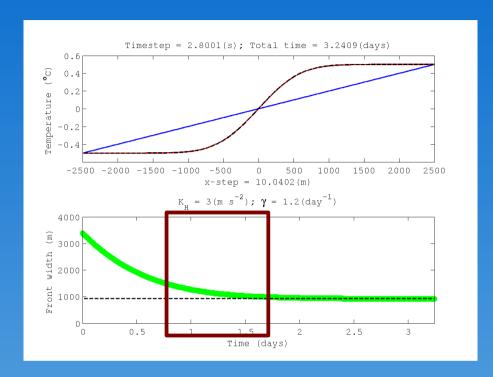


Numerical solution



First order upwind scheme





- Fast adjustment to equilibrium (within 1 day)
- However Kh proportional to square of width
- Even small errors in width could affect estimate of Kh